

Article

Analysis on the Effects of SAR Imaging Parameters and Environmental Conditions on the Standard Deviation of the Co-Polarized Phase Difference Measured over Sea Surface

Andrea Buono ^{1,*},†,‡ , Carina Regina de Macedo ^{1,‡}, Ferdinando Nunziata ^{1,‡} ,
Domenico Velotto ^{2,‡} and Maurizio Migliaccio ^{1,‡}

¹ Dipartimento di Ingegneria, Università degli Studi di Napoli Parthenope, 80143 Napoli NA, Italy; carinaregina.demacedo@uniparthenope.it (C.R.d.M.); ferdinando.nunziata@uniparthenope.it (F.N.); maurizio.migliaccio@uniparthenope.it (M.M.)

² German Aerospace Center (DLR), Remote Sensing Technology Institute, 82234 Weßling, Germany; domenico.velotto@dlr.de

* Correspondence: andrea.buono@uniparthenope.it; Tel.: +39-081-547-6706

† Current address: Centro Direzionale, isola C4, 80143 Napoli, Italy.

‡ These authors contributed equally to this work.

Received: 15 November 2018; Accepted: 19 December 2018; Published: 21 December 2018



Abstract: This study aimed at analyzing the effect of Synthetic Aperture Radar (SAR) imaging parameters and environmental conditions on the standard deviation of the co-polarized phase difference (σ_{φ_C}) evaluated over sea surface. The latter was shown to be an important polarimetric parameter widely used for sea surface target monitoring purposes. A theoretical model, based on the tilted-Bragg scattering, is proposed to predict the behavior of σ_{φ_C} against incidence angle for different roughness conditions. Then, a comprehensive experimental analysis, based on the processing of L-, C- and X-band polarimetric SAR scenes collected over different test areas under low-to-moderate wind conditions and covering a broad range of incidence angle, was carried out to discuss the effects of sensor's and environmental parameters on sea surface σ_{φ_C} . Results show that SAR imaging parameters severely affect σ_{φ_C} , while the impact of meteo-marine conditions, under low-to-moderate wind regime, is almost negligible. Those outcomes have significant relevance to support the design of effective and robust algorithms for marine and maritime applications based on σ_{φ_C} , including the detection of metallic targets (ships and offshore infrastructures as oil/gas platforms, aquacultures, wind farms, etc.) and polluted areas.

Keywords: oceans; Synthetic Aperture Radar; polarimetry; co-polarized phase difference

1. Introduction

The preservation and sustainable management of ocean resources and ecosystems are mandatory goals according to several international strategies, policy programmes and technical reports as the 2030 Agenda, the Blue Growth European Union program and the United Nations report for sustainable development [1–5]. Within this context, remote sensing tools, due to their synoptic view and frequent revisit time, are of paramount importance for a broad range of applications including global weather predictions, storm and hurricane warnings, wave and current forecastings, coastal storm surges, ship routing, commercial fishing and climate change [6].

Microwave imaging of sea surface from space allows the retrieval of several geophysical parameters that, once tailored models are available, can be transformed into added-value products that

include, among others, significant wave height from radar altimeters, wave spectrum from Synthetic Aperture Radars (SARs), sea surface wind field from scatterometers and sea surface temperature and salinity from radiometers [6]. Furthermore, additional information on ocean targets of interest such as sea ice extent, pollutants, metallic infrastructures, icebergs, ships, etc. can be derived.

When dealing with microwave active remote sensing of the oceans, the unprecedented benefits offered by SAR sensors operating in multi-polarization modes is unambiguously recognized for a wide range of marine and maritime applications, including coastline extraction [7–9], metallic target detection [10–12], sea pollution monitoring [13–16] and sea ice observation [17–19]. All the approaches share a similar physical rationale that relies on the exploitation of the different polarimetric properties that characterize the target of interest and the sea clutter. Nonetheless, the performance is significantly affected by the amount of scattering information available that, in turn, depends on [8,10,14,20,21]:

- SAR acquisition parameters, e.g., polarization, Angle Of Incidence (AOI), incident wavelength and Noise-equivalent sigma zero (NESZ);
- target features, e.g., damping properties of the pollutant, ship orientation and ice layer thickness; and
- meteo-marine conditions, e.g., sea state, swell and wave patterns.

Among the different polarimetric SAR architectures, conventional dual-polarimetric SAR imaging modes, i.e., the ones that consist of transmitting a linearly polarized wave (horizontal (H) or vertical (V)) while receiving coherently in an orthogonal linear basis (H–V), are attracting more attention in the perspective of operational services since they offer, for a wide range of ocean applications, a sufficient polarimetric information content over a large swath [7,16,22]. Among the dual-polarimetric SARs, it is worth mentioning the one operated by the German TerraSAR-X (TSX) mission that provides coherent HH–VV SAR measurements routinely. Among the polarimetric features that can be extracted from a coherent dual-polarimetric HH–VV SAR, the standard deviation of the co-polarized phase difference (σ_{φ_C}) has been demonstrated to be a powerful tool for marine and maritime applications. In fact, it is successfully used to perform the observation of metallic targets [23–26], the monitoring of oil pollution [14,16,27,28] and iceberg detection/sea ice classification [29,30].

All those applications share the same physical rationale, i.e., they exploit σ_{φ_C} as a reliable and robust estimator of the correlation between the co-polarized channels. This means that, at least from a theoretical viewpoint, since sea surface scattering is ruled by the Bragg mechanism, the co-polarized channels result in a unitary correlation that makes the co-polarized phase difference statistical distribution resembling a Dirac delta function [31–34]. In real cases, when a low-depolarizing scenario is considered, e.g., a sea surface Bragg scattering, a large correlation between co-polarized channels applies that results in a narrow co-polarized phase difference distribution. Hence, marine targets, e.g., ships, oil slicks, icebergs, etc., that result in departure from Bragg scattering are characterized by a broader co-polarized phase difference distribution. This means that σ_{φ_C} can be successfully exploited to emphasize non-Bragg scattering targets with respect to the Bragg-like sea surface background.

Following this rationale, in the literature, σ_{φ_C} has been mainly investigated in terms of sea/target separability (e.g., mean contrast and average target-to-clutter ratio are usually adopted as figures of merit [17,20,23,26,35]). In [28], σ_{φ_C} values spanning the range ≈ 3 – 18° have been reported, at C-band, over sea surface under low-to-moderate wind conditions over different geographical locations. Dierking and Wesche [29] found that, at C-band, larger σ_{φ_C} values are in place when the measured co-polarized channels approach the sensor's noise floor. In [30], σ_{φ_C} values of about 6° have been measured at C-band with AOI = 40° over sea surface and thin sea ice. In [10], larger σ_{φ_C} values, i.e., about 30° , have been measured over sea surface at X-band. The authors also observed a sensitivity of σ_{φ_C} with both the estimation window's size and the sea state parameters. In [36], the behavior of σ_{φ_C} is investigated with reference to a specific test case, i.e., the Taylor Energy oil seep in the Gulf of Mexico, where a large TSX SAR dataset was considered under limited SAR imaging configuration (X-band, AOI = 26° , 34° and 43°).

Hence, literature studies clearly point out the dependence of σ_{φ_C} on both sensor's and environmental parameters. Nevertheless, the behavior of σ_{φ_C} is analyzed for specific applications only, i.e., for a given target (ships and oil slicks); sensor's configuration, i.e., for a given frequency and/or limited AOI range; and environmental conditions [16,23,29,30,37]. In no case, to the best of our knowledge, the role of sensor's and environmental parameters on σ_{φ_C} is investigated in a systematic way.

In this study, a significant extension of the work carried out in [36] was made. The analysis undertaken in [36] was improved as follows:

- The analysis on the effects of different frequencies, i.e., L- and C-band, and sea surface tilting angle on sea surface σ_{φ_C} was included.
- The analysis on the influence of incidence angle on sea surface σ_{φ_C} was undertaken on a much broader AOI range, i.e., about 20–60°.
- The behavior of sea surface σ_{φ_C} with respect to AOI was investigated by comparing model predictions' with actual SAR measurements over the whole range of Bragg AOIs (≈ 20 –60°).

Hence, a comprehensive analysis on the behavior of σ_{φ_C} over sea surface was provided for the first time. A large polarimetric SAR dataset collected in a wide range of SAR acquisition parameters (frequency, AOI, and NESZ) and meteo-marine conditions (sea surface roughness and wind speed (WS)) was considered to give a more complete understanding on how those parameters affect sea surface σ_{φ_C} . Furthermore, to better interpret the experimental results, a theoretical scattering framework, based on the polarimetric X-Bragg model, is proposed that allows giving a physical understanding of the behavior of sea surface σ_{φ_C} under noise-free conditions. It must be explicitly pointed out that this comprehensive study is of paramount importance to support all the operational techniques based on σ_{φ_C} , e.g., to find the most suitable SAR configuration in relation with the ocean target to be detected.

The remainder of the paper is organized as follows: in Section 2, the theoretical background that lies at the basis of the sensitivity analysis is presented; in Section 3, SAR dataset and ancillary information are provided; in Section 4, the experiments based on the theoretical model provided in Section 2 are presented and discussed; in Section 5, experiments related to the sensitivity analysis undertaken on the actual SAR dataset are presented and discussed; and conclusions are drawn in Section 6.

2. Polarimetric Framework

The co-polarized phase difference was theoretically predicted over sea surface using a polarimetric sea surface scattering model. The latter, based on the X-Bragg scattering model developed in [36], assumed that sea surface is mainly governed by the tilted-Bragg scattering mechanism. This is a reasonable assumption under low-to-moderate wind conditions (i.e., ≈ 2 –14 m/s) and under intermediate AOIs (i.e., ≈ 20 –60°).

From the scattering viewpoint, sea surface is considered as a distributed low-depolarizing scene. Hence, second-order descriptors, i.e., the covariance matrix \mathbf{C} , are needed to describe sea surface polarimetric scattering [38]. In the monostatic backscattering case, invoking reciprocity and the reflection symmetry property, the 3×3 Hermitian and semi-definite positive covariance matrix \mathbf{C} can be defined as follows [38]:

$$\mathbf{C} = \begin{pmatrix} C_{11} & C_{12} & C_{13} \\ C_{12}^* & C_{22} & C_{23} \\ C_{13}^* & C_{23}^* & C_{33} \end{pmatrix} = \begin{pmatrix} \langle |S_{hh}|^2 \rangle & 0 & \langle S_{hh}S_{vv}^* \rangle \\ 0 & 2\langle |S_{hv}|^2 \rangle & 0 \\ \langle S_{vv}S_{hh}^* \rangle & 0 & \langle |S_{vv}|^2 \rangle \end{pmatrix}, \quad (1)$$

where S_{pq} is the complex scattering amplitude with $\{p, q\} = \{h, v\}$, while $\langle \cdot \rangle$, $|\cdot|$ and $*$ stand for ensemble average, modulus and complex conjugate operators, respectively.

According to the X-Bragg scattering theory, the sea surface covariance matrix can be predicted as follows [39–41]:

$$\mathbf{C}_X = \begin{pmatrix} C_{X11} & 0 & C_{X13} \\ 0 & C_{X22} & 0 \\ C_{X13}^* & 0 & C_{X33} \end{pmatrix} = \mathbf{U}^{-1} \begin{pmatrix} |B_h + B_v|^2 & (B_h + B_v)(B_h^* - B_v^*)\text{sinc}(2\beta) & 0 \\ (B_h^* + B_v^*)(B_h - B_v)\text{sinc}(2\beta) & \frac{1}{2}|B_h - B_v|^2(1 + \text{sinc}(4\beta)) & 0 \\ 0 & 0 & \frac{1}{2}|B_h - B_v|^2(1 - \text{sinc}(4\beta)) \end{pmatrix} \mathbf{U} \quad (2)$$

where the subscript “X” stands for X-Bragg model and \mathbf{U} is a con-similarity linear transformation given by [38]:

$$\mathbf{U} = \frac{1}{\sqrt{2}} \begin{pmatrix} 1 & 0 & 1 \\ 1 & 0 & -1 \\ 0 & \sqrt{2} & 0 \end{pmatrix} \quad (3)$$

In Equation (2), β is the surface local tilting angle, while B_h and B_v are the Fresnel complex coefficients for Bragg scattering under horizontal and vertical polarizations, respectively, which are given by [41]:

$$\begin{cases} B_h = \frac{\cos(\text{AOI}) - \sqrt{\epsilon - \sin^2(\text{AOI})}}{\cos(\text{AOI}) + \sqrt{\epsilon - \sin^2(\text{AOI})}} \\ B_v = \frac{(\epsilon - 1) \left(\sin^2(\text{AOI}) - \epsilon(1 + \sin^2(\text{AOI})) \right)}{\left(\epsilon \cos(\text{AOI}) + \sqrt{\epsilon - \sin^2(\text{AOI})} \right)^2} \end{cases} \quad (4)$$

In the X-Bragg sea surface scattering model, β is related to the amount of surface roughness and it rules both the cross-polarized backscattering and the HH-VV coherence, while the Bragg scattering coefficients are related to the incidence angle and to the frequency-dependent relative electric permittivity of seawater, ϵ [41]. β was assumed to be uniformly distributed in the range $[-90^\circ, +90^\circ]$, where $\beta \approx 0^\circ$ describes an almost flat sea surface (negligible cross-polarized backscattering and HH-VV coherence close to 1), while $\beta \approx \pm 90^\circ$ characterizes an extremely rough sea surface (significant cross-polarized backscattering and HH-VV coherence that tends to 0) [41]. In [15,40], it is found that reliable β values fall within the range $[-30^\circ, 30^\circ]$.

Once \mathbf{C} is predicted according to the X-Bragg scattering model, the co-polarized phase difference can be obtained:

$$\angle C_{X13} = \angle \left(\frac{1}{2}|B_h + B_v|^2 - \frac{1}{4}|B_h - B_v|^2(1 - \text{sinc}(4\beta)) - j\Im(B_h^* + B_v^*)(B_h - B_v)\text{sinc}(2\beta) \right) \quad (5)$$

where $\angle(\cdot)$ and $\Im(\cdot)$ mean phase and imaginary part, respectively. Note that Equation (5) allows explicitly pointing out the relationship between the co-polarized phase difference and SAR acquisition parameters (through the Bragg scattering coefficients, i.e., AOI and frequency) and geometric/dielectric properties of sea surface (through β and ϵ). Hence, it is a starting point to understand the effects of such parameters on the co-polarized phase difference distribution.

3. Datasets

In this section, the SAR dataset and the ancillary data are briefly described.

3.1. SAR Dataset

The L-, C- and X-band polarimetric SAR dataset consists of 14 scenes acquired between June 2006 and July 2013. The two L-band Single-look complex (SLC) full-polarimetric Uninhabited Aerial Vehicle Synthetic Aperture Radar (UAVSAR) scenes were acquired with an AOI ranging from about 22° to

65° in the northern part of the Gulf of Mexico. The UAVSAR system operates at center frequency of 1.26 GHz and is characterized by a NESZ of about −53 dB at mid-swath [42].

The four L-band SLC full-polarimetric ALOS PALSAR-1 (AP) scenes were acquired between June 2006 and April 2009 with an AOI from about 22° to 25° in the northern part of the Gulf of Mexico and off the northeastern coast of Brazil. The Japanese AP system operates at center frequency of 1.27 GHz with a NESZ of −29 dB at near range (21.5°) [43].

The four C-band SLC full-polarimetric RADARSAT-2 (RS) scenes were collected between January 2009 and May 2010 with AOI ranging from about 22° to 41° in the northern part of the Gulf of Mexico and off the California coast. The Canadian RS system operates at center frequency of 5.40 GHz with a NESZ of −35 dB [44].

The four X-band SLC dual co-polarimetric HH-VV TSX scenes were collected between December 2011 and July 2012 with an AOI range of 25° to about 35° in the northern part of the Gulf of Mexico. The German TSX system operates at center frequency of 9.60 GHz with a mean NESZ of −22 dB [45].

Figure 1 shows excerpts (1000 × 1000 pixels) of the VV-polarized intensity images of the whole SAR dataset representing homogeneous sea surface areas. Figure 1 is organized in a matrix format, where rows refer to the different SAR sensors (UAVSAR, AP, RS and TSX from the top to the bottom) and columns are arranged according to increasing AOI (from left to right). Note that, even though the same decibel (dB) scale is adopted, i.e., [0 -40], the sea surface patterns and, therefore, the backscattering change significantly according to the SAR imaging parameters and meteo-marine conditions. An overview of the SAR dataset is provided in Tables 1 and 2. Note that the whole SAR dataset is partitioned according to three AOI ranges: low (22–27°), intermediate (31–35°) and high (38–42°) (see Table 2).

Table 1. General information on SAR imagery.

| SAR Sensor | UAVSAR | AP | RS | TSX |
|--------------------------------------|-------------------|-------------------|-------------------|----------------------|
| Frequency (GHz) | 1.26 | 1.27 | 5.40 | 9.60 |
| Imaging mode | Full-polarimetric | Full-polarimetric | Full-polarimetric | Dual co-polarimetric |
| Slant range × azimuth resolution (m) | 1.7 × 1.0 | 9.4 × 3.6 | 4.7 × 5.1 | 1.2 × 6.6 |
| Number of scenes | 2 | 4 | 4 | 4 |
| Nominal NESZ (dB) | −53 | −29 | −35 | −22 |

Table 2. Overview of the dataset.

| SAR Sensor | SAR Data | | | Ancillary Wind Info | |
|------------|----------|------------------|---------------|---------------------|---------------|
| | Scene ID | Acquisition Date | AOI Range (°) | Speed (m/s) | Direction (°) |
| UAVSAR | 1 | 25/07/2013 | 22.0–27.0 | 4.2 | 115.4 |
| | | | 31.0–35.0 | | |
| | | | 38.0–42.0 | | |
| | 2 | 22/06/2010 | 22.0–26.0 | 7.1 | 321.0 |
| AP | 3 | 02/04/2009 | 22.7–25.0 | 3.1 | 232.0 |
| | 4 | 20/06/2006 | 22.7–25.0 | 4.9 | 298.6 |
| | 5 | 14/03/2007 | 22.7–25.0 | 7.4 | 332.2 |
| | 6 | 19/03/2009 | 22.7–25.0 | 8.0 | 243.3 |
| RS | 7 | 31/01/2009 | 22.6–24.2 | 9.0 | 215.1 |
| | 8 | 04/05/2010 | 23.4–25.3 | 4.3 | 129.3 |
| | 9 | 26/09/2009 | 31.3–33.0 | 4.7 | 142.9 |
| | 10 | 01/05/2010 | 39.3–40.7 | 12.0 | 331.3 |
| TSX | 11 | 01/07/2012 | 25.0–26.7 | 4.1 | 1.3 |
| | 12 | 05/12/2011 | 25.0–26.7 | 9.2 | 5.1 |
| | 13 | 17/05/2012 | 33.0–34.5 | 3.8 | 168.6 |
| | 14 | 19/06/2012 | 33.0–34.5 | 8.4 | 305.0 |

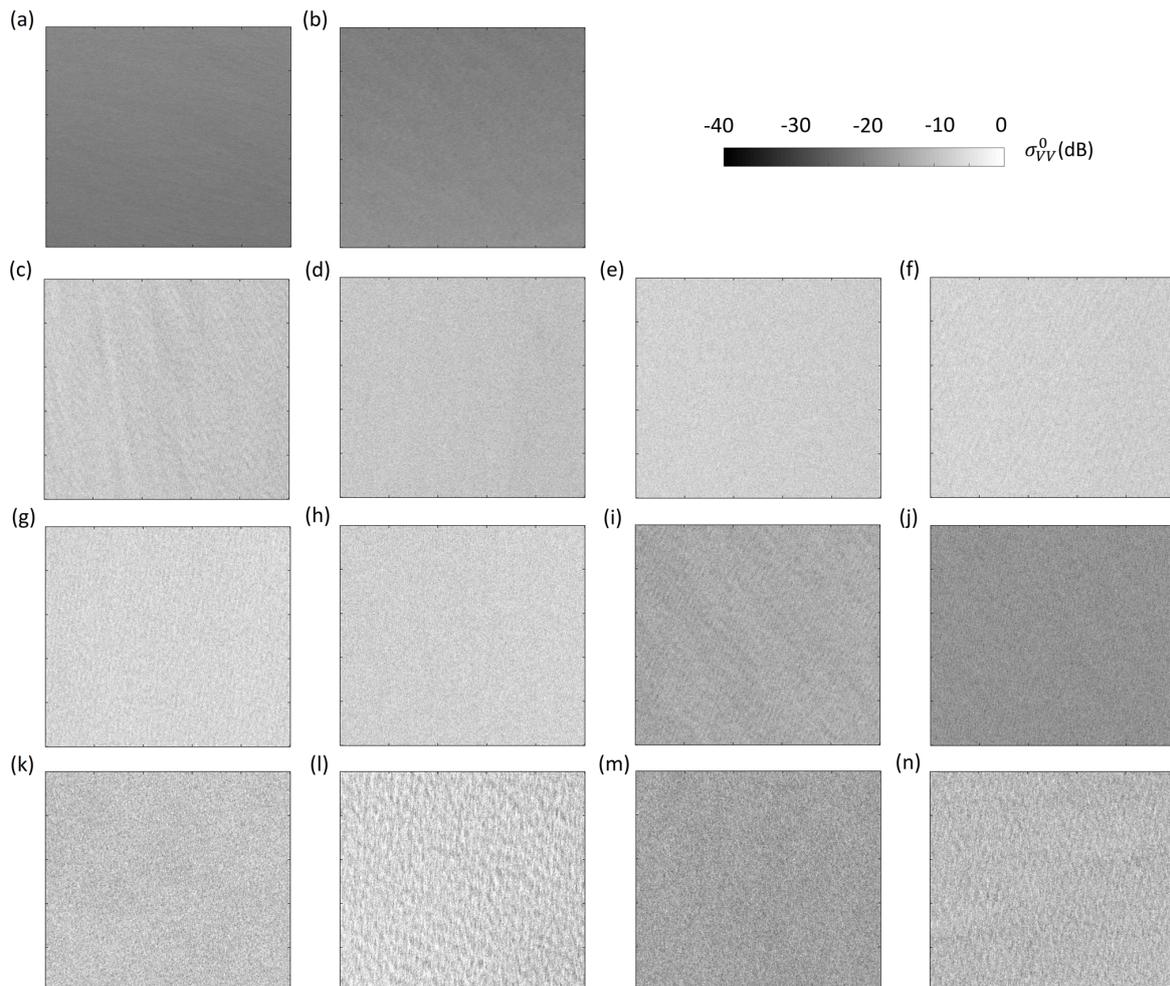


Figure 1. SAR dataset. VV-polarized intensity images (excerpts' size: 1000×1000), in dB scale, over homogeneous sea surface area. Rows refer to the different SAR sensors: (a,b) UAVSAR; (c–f) AP; (g–j) RS; and (k–n) TSX. Columns are organized according to increasing AOI: (a,c,g,k) low, i.e., $22\text{--}27^\circ$; (d,e,h,i,l,m) intermediate, i.e., $31\text{--}35^\circ$; and (b,f,j,n) high, i.e., $38\text{--}42^\circ$.

3.2. Ancillary Wind Field Information

The wind information was provided by the Physical Oceanography Distributed Active Archive Center (PO.DAAC) that makes available science data to a wide user community [46–48]. The wind information was collected from two different sources depending on the SAR imagery acquisition period. The wind field co-located to SAR scenes acquired before 2012 was collected from the Cross-Calibrated Multi-Platform (CCMP) project, while the wind field related to the SAR scenes acquired since 2012 was collected from the Advanced Scatterometer (ASCAT) instrument on MetOp-A (Meteorological Operational Satellites) satellite.

The CCMP product is characterized by a spatial resolution of about 25 km and combines cross-calibrated satellite winds from RMS Remote Sensing Systems derived from microwave radiometers and from scatterometers, in situ measurements and reanalysis data from the European Center for Medium-Range Weather Forecasts (ECMWF) ERA-40 Reanalysis [49]. The wind products obtained from ASCAT are characterized by an effective resolution of 50 km and are delivered by the European Organization for the Exploitation of Meteorological Satellites (EUMETSAT) Ocean and Sea Ice Satellite Application Facility (OSI SAF) through the Royal Netherlands Meteorological Institute (KNMI) [50].

The SAR scenes were all collected under low-to-moderate wind conditions, i.e., from about 3 m/s to 12 m/s. The wind information is summarized in Table 2.

4. Model-Based Experimental Results

In this section, the behavior of σ_{φ_C} against AOI and β is discussed using both the theoretical model presented in Section 2 and actual polarimetric SAR measurements collected by UAVSAR (see Tables 1 and 2). The latter, on the one side, is characterized by a very low NESZ that guarantees high-quality measurements even at larger AOIs, while, on the other side, UAVSAR scenes cover a broad range of AOIs (spanning from about 20° to 60°) and, therefore, they make possible comparing theoretical predictions with actual measurements over the whole range of Bragg AOIs.

To predict realistic σ_{φ_C} values, we used β values estimated from the UAVSAR scenes using the formula proposed in [51]:

$$\text{sinc}(4\beta) = \frac{\text{Tr}(\mathbf{C}) - 2C_{22} - 0.5\Re(C_{13})}{\text{Tr}(\mathbf{C}) - 0.5\Re(C_{13})}, \quad (6)$$

where $\text{Tr}(\cdot)$ is the trace operator and $\Re(\cdot)$ means real part. Hence, for each SAR scene, three equal-sized homogeneous sea surface regions of interest (ROIs) were excerpted that are characterized by low (22–27°), intermediate (31–35°) and high (38–42°) AOI, respectively. Accordingly, to better understand the behavior of β , normalized histograms were computed (see Figure 2). It can be noted that the empirical statistical distribution of β (see blue histograms) resembles a Gaussian bell for any AOI, as witnessed by the fitted distribution (see red curves) that satisfies the chi-square test with 0.05 confidence interval.

Once β was obtained, to predict σ_{φ_C} , a Monte Carlo approach based on 1000 independent simulations for each AOI was implemented.

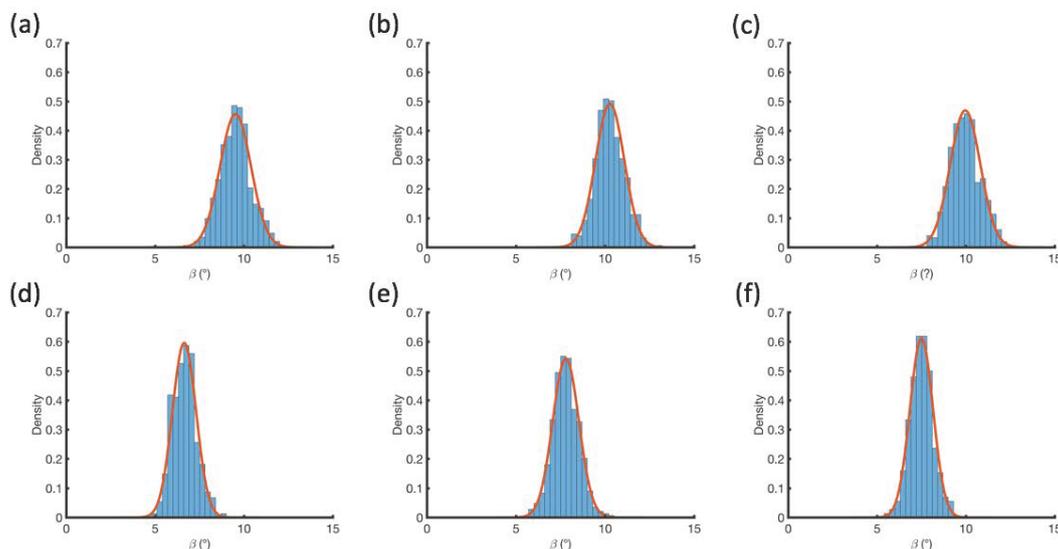


Figure 2. Normalized β histograms and fitted Gaussian distributions referring to: (a–c) UAVSAR scene ID 1 acquired at low (22–27°), intermediate (31–35°) and high (38–42°) AOIs, respectively; and (d–f) the same for UAVSAR scene ID 2.

For each simulation, β was randomly selected according to a Gaussian distribution whose mean and standard deviation values were obtained from the UAVSAR ROIs, i.e., 9.9° and 0.8° for the UAVSAR scene ID 1 and 7.3° and 0.7° for the UAVSAR scene ID 2 (see Table 3).

Simulation results are depicted in Figure 3a, where the continuous and dashed lines are related to β extracted from UAVSAR scenes ID 1 and ID 2, respectively. It can be noted that, in both cases, σ_{φ_C} increases with AOI, with σ_{φ_C} values ranging from about 0.01° at near range to about 0.15° at far range.

σ_{φ_C} values measured from actual UAVSAR scenes using a 9×9 moving window are depicted in Figure 3b, where continuous and dashed lines stand for UAVSAR scene ID 1 and ID 2, respectively. It can be noted that both measured and simulated σ_{φ_C} increase with the incidence angle. This witnesses that the X-Bragg model succeeds in predicting the actual behavior of σ_{φ_C} with respect to AOI. However, the simulated and measured σ_{φ_C} values are completely different, i.e., the simulated σ_{φ_C} values are about one order of magnitude smaller (measured σ_{φ_C} values range from about 0.5° at near range to about 4° at far range). Nonetheless, this is not a surprising result. The theoretical scattering model predicts a noise-free behavior referred to a low-depolarizing Bragg scattering surface, i.e., the co-polarized phase difference distribution should resemble, ideally, a Dirac delta function (i.e., $\sigma_{\varphi_C} = 0^\circ$), while UAVSAR measurements, although very accurate, are noisy and refer to a real scattering surface, i.e., sensor and environmental parameters cause decorrelation between the co-polarized channels that results in a broader co-polarized phase difference distribution (i.e., larger σ_{φ_C} values). Note that a similar trend for σ_{φ_C} is experimentally observed in [42]. Notwithstanding that, the global σ_{φ_C} behavior with respect to AOI is well-described by the X-Bragg scattering model.

Figure 3 also shows that the model succeeds in predicting σ_{φ_C} values related to the UAVSAR scene ID 1, which are slightly larger than the corresponding ones related to the UAVSAR scene ID 2. This is because actual β values were used in the simulations.

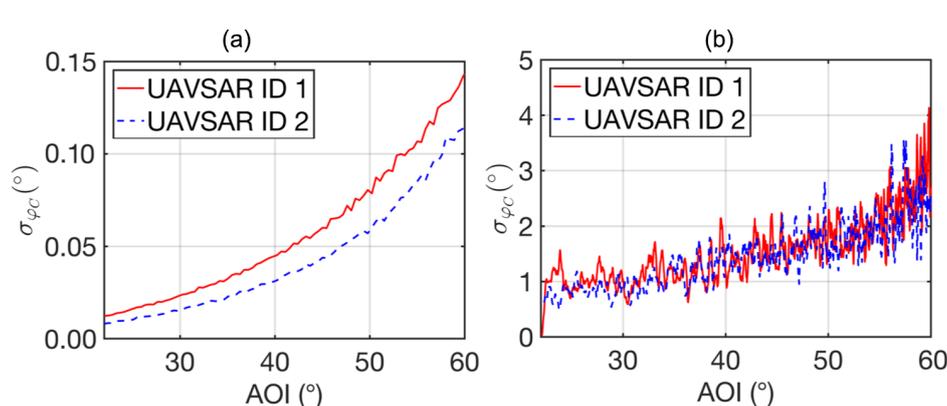


Figure 3. Simulated (a); and measured (b) behavior of σ_{φ_C} with respect to AOI relevant to UAVSAR scenes ID 1 (red line) and 2 (blue dashed line).

5. Experimental Results

The sensitivity of σ_{φ_C} with respect to acquisition parameters, i.e., AOI, NESZ and incident wavelength, and meteo-marine conditions, i.e., WS and β , was analyzed using actual polarimetric SAR data (Tables 1 and 2). When dealing with SAR imaging parameters, hardware and technical constraints suggest that there is a close relationship among NESZ, AOI and wavelength [45,52]. In addition, from a scattering viewpoint, when a Bragg surface backscattering is considered, the larger is the AOI, the lower is the signal-to-noise ratio (SNR), i.e., the mean ratio between co-polarized sea surface backscattering and the nominal NESZ [36,42]. When dealing with meteo-marine parameters, both β and near-surface wind speed are related to geometrical sea surface characteristics, i.e., sea surface roughness. Hence, a preliminary analysis devoted to understanding the relationship between β and WS is due.

To accomplish this task, the normalized β histograms were evaluated for each ROI belonging to AP and RS SAR scenes for low, intermediate and high AOIs (see Figure 4), where the mean WS values obtained from the external sources described in Section 3.2 were also annotated. First, it can be noted that β normalized histograms are broader (about 2–3 times) than the corresponding UAVSAR ones (see Figure 2). This is most likely due to the higher NESZ that characterizes AP and RS rather than rougher sea state conditions (see Table 3). In addition, no clear relationship between WS and β was observed. This suggests that, although both β and WS are related to the sea surface roughness, there is

no clear link between them. Consequently, the sensitivity of σ_{φ_C} with respect to WS and β deserves to be analyzed separately.

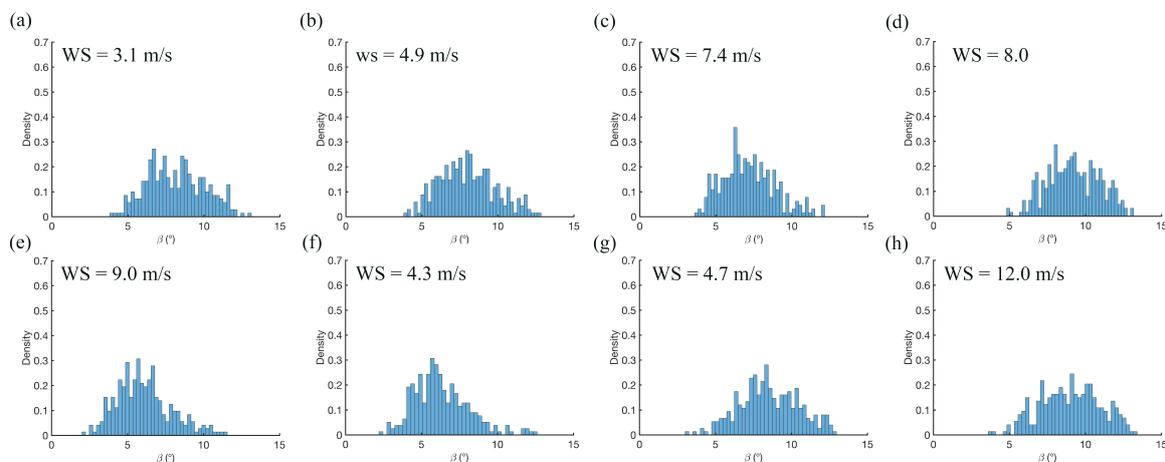


Figure 4. Normalized β histograms, evaluated over the homogeneous sea surface ROIs. The figure is organized in a matrix format. Rows refer to: AP (a–d); and RS (e–h) SAR sensors. Columns refer to the AOI: (a,e) low (22–27°); (b,c,f,g) intermediate (31–35°); and (d,h) high (38–42°). The corresponding mean WS values are also annotated.

Hence, the whole SAR dataset was processed according to the methodology described in Section 4 for the UAVSAR case. The only difference relies on the fact that AP, RS and TSX SAR scenes are characterized by a significantly narrower AOI range if compared to UAVSAR (less than 3° from near to far range, see Table 2) and, therefore, only one ROI was selected for each scene. σ_{φ_C} maps, obtained according to Migliaccio et al. [28], are shown in Figure 5, where it can be noted how σ_{φ_C} is affected by both SAR imaging parameters and environmental conditions. In fact, σ_{φ_C} values range from about 0° to more than 15° along the whole dataset. The lowest σ_{φ_C} values are shown in Figure 5a,b, corresponding to UAVSAR acquisitions (L-band, NESZ = −53 dB), while the largest values are related to TSX acquisitions (X-band, NESZ = −22 dB) (see Figure 5m,n). In addition, since columns in Figure 5 are organized according to increasing AOI (as in Figure 1), it can be noted that σ_{φ_C} increases with AOIs, as suggested by the model-based analysis presented in Section 4.

The visual interpretation was confirmed by the quantitative analysis undertaken on 1000 independent samples, randomly selected from the ROIs, whose σ_{φ_C} mean and standard deviation values were evaluated together with the average SNR (see Table 3). In Table 3, mean and standard deviation values of the β parameter estimated from the whole dataset according to Equation (6) are also listed. It must be noted that β values related to TSX SAR scenes are not available since full-polarimetric information is needed to estimate β . The following subsections focus on the quantitative analysis of the effects produced by sensor and scene parameters on sea surface σ_{φ_C} , separately.

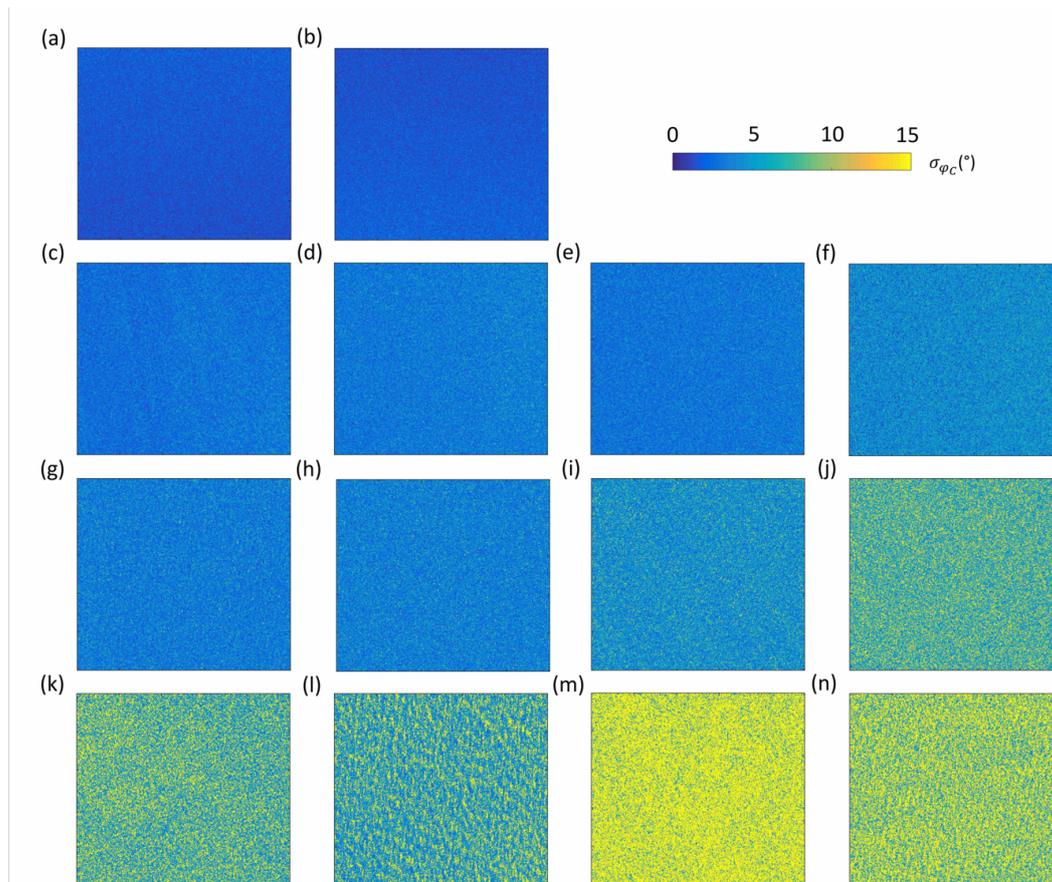


Figure 5. σ_{φ_C} ($^{\circ}$) images evaluated over the ROIs shown in Figure 1. Rows refer to the different SAR sensors: (a,b) UAVSAR; (c–f) AP; (g–j) RS; and (k–n) TSX. Columns are organized according to increasing AOI: (a,c,g,k) low, i.e., 22–27; (d,e,h,i,l,m) intermediate, i.e., 31–35; and (b,f,j,n) high, i.e., 38–42.

5.1. Sensitivity Analysis: Meteo-Marine Parameters

The influence of WS and β on σ_{φ_C} was analyzed. When dealing with the effects of WS on σ_{φ_C} , according to Table 2, the AP SAR imagery were considered since the four SAR scenes were acquired under very similar SAR imaging parameters, i.e., the same AOI ($\approx 24^{\circ}$) and incident wavelength (L-band), and by almost the same estimated mean SNR, i.e., ≈ 19 – 22 dB for both channels (see Table 3). Results listed in Table 3 clearly point out that, although the AP SAR scenes were collected under different wind conditions, i.e., ≈ 3 m/s and 8 m/s for SAR scene ID 3 and 6, respectively, they are characterized by very similar σ_{φ_C} mean values, i.e., $\approx 3.1^{\circ}$. In addition, WS does not affect significantly the variability of σ_{φ_C} —see the standard deviation of σ_{φ_C} in Table 3, which is close to $\approx 1.3^{\circ}$ for all AP SAR scenes. The same comments apply for SAR scene ID 4 and 5. Hence, it can be concluded that, under low-to-moderate wind conditions, σ_{φ_C} is almost unaffected by WS.

To discuss the effects of β on σ_{φ_C} , it is worth analyzing results that refer to AP and RS SAR imagery—see ID 5–8 in Table 3, since they refer to SAR scenes collected under similar acquisition parameters, i.e., L-band, AOI of $\approx 24^{\circ}$ and average SNR of about 21 dB (C-band, AOI of $\approx 24^{\circ}$ and average SNR of about 28 dB). It can be noted that, from the results in Table 3, SAR scenes ID 5 and 6 are characterized by almost the same σ_{φ_C} values, i.e., $3.1^{\circ} \pm 1.3^{\circ}$, although SAR scene ID 6 is characterized by a β value that is about 30% larger than the one related to the SAR scene ID 5. The same comments apply for SAR scene ID 7 and 8, which are characterized by almost the same σ_{φ_C} values, i.e., $4.0^{\circ} \pm 2.0^{\circ}$ and $3.7^{\circ} \pm 1.7^{\circ}$, respectively, although the β parameter estimated from SAR scene ID 7 is about 25% larger than the one evaluated from SAR scene ID 8.

Hence, this experimental analysis shows that there is no clear trend between β and σ_{φ_C} . In particular, β does not play a dominant role in broadening/shrinking the co-polarized phase difference distribution.

Table 3. Values of σ_{φ_C} , β and HH and VV SNR evaluated within ROIs selected over a homogeneous sea surface area. σ_{φ_C} and β are presented as mean \pm standard deviation value.

| SAR Sensor | ID | AOI (°) | σ_{φ_C} (°) | WS (m/s) | β (°) | SNR (dB) | |
|------------|----|-----------|--------------------------|----------|----------------|----------|------|
| | | | | | | HH | VV |
| UAVSAR | 1 | 22.0–27.0 | 1.0 ± 0.4 | 4.2 | 9.6 ± 0.8 | 39.0 | 38.7 |
| | | 31.0–35.0 | 1.1 ± 0.5 | | 10.2 ± 0.9 | 35.6 | 35.9 |
| | | 38.0–42.0 | 1.4 ± 0.5 | | 10.0 ± 0.7 | 30.1 | 32.1 |
| | 2 | 22.0–26.0 | 0.7 ± 0.3 | 7.1 | 6.6 ± 0.7 | 38 | 40.3 |
| | | 31.0–35.0 | 1.0 ± 0.4 | | 7.7 ± 0.7 | 35.7 | 38.7 |
| | | 38.0–42.0 | 1.3 ± 0.5 | | 7.5 ± 0.6 | 30.2 | 34.9 |
| AP | 3 | 22.7–25.0 | 3.0 ± 1.3 | 3.1 | 8.3 ± 2.0 | 18.7 | 20.3 |
| | 4 | 22.7–25.0 | 3.2 ± 1.3 | 4.9 | 8.2 ± 2.0 | 19.3 | 20.3 |
| | 5 | 22.7–25.0 | 3.2 ± 1.4 | 7.4 | 7.3 ± 1.7 | 21.2 | 22.0 |
| | 6 | 22.7–25.0 | 3.0 ± 1.2 | 8.0 | 9.5 ± 2.3 | 20.8 | 21.5 |
| RS | 7 | 22.6–24.2 | 4.0 ± 2.0 | 9.0 | 5.9 ± 1.7 | 27.8 | 28.2 |
| | 8 | 23.4–25.3 | 3.7 ± 1.7 | 4.3 | 6.7 ± 1.9 | 27.4 | 27.8 |
| | 9 | 31.3–33.0 | 5.0 ± 2.3 | 4.7 | 8.4 ± 2.6 | 19.2 | 21.4 |
| | 10 | 39.3–40.7 | 8.0 ± 4.0 | 12.0 | 9.6 ± 3.2 | 14.5 | 18.2 |
| TSX | 11 | 25.0–26.7 | 8.8 ± 6.4 | 4.1 | – | 9.2 | 9.8 |
| | 12 | 25.0–26.7 | 7.4 ± 5.7 | 9.2 | – | 10.2 | 10.4 |
| | 13 | 33.0–34.5 | 18.4 ± 16.5 | 3.8 | – | 2.7 | 3.8 |
| | 14 | 33.0–34.5 | 12.6 ± 10.0 | 8.4 | – | 5.8 | 7.9 |

5.2. Sensitivity Analysis: SAR Imaging Parameters

The influence of noise floor, incidence angle and wavelength on σ_{φ_C} was analyzed. As explicitly pointed out above, it is not straightforward to isolate their individual contribution since they are inter-connected.

To analyze the effects of NESZ on σ_{φ_C} , SAR scenes collected by the L-band UAVSAR and AP platforms were considered since they are characterized by completely different NESZ values, i.e., -53 dB and -29 dB, respectively (see Table 1), while sharing the same operating frequency. In addition, the AOI range that characterizes AP SAR scenes was included into the one provided by UAVSAR. Hence, those datasets allow a fair analysis of the impact of noise floor in broadening/shrinking the co-polarized phase difference distribution. UAVSAR scenes ID 1 and ID 2 and AP SAR scenes ID 3–6 are characterized by a SNR equal to ≈ 39 dB and ≈ 21 dB, respectively (see Table 3). Their mean σ_{φ_C} values are 0.85° and 3.1° , respectively. Hence, one can note that, when halving the SNR, σ_{φ_C} increases of about four times. In addition, results listed in Table 3 show that a lower SNR corresponds to a larger σ_{φ_C} variability, whose standard deviation increases from 0.35° (UAVSAR scenes ID 1–2, low AOI) to 1.3° (AP SAR scenes ID 3–6), i.e., about four times.

When dealing with the effects of AOI on σ_{φ_C} , first results presented and discussed in Section 4, obtained considering UAVSAR imagery, suggest that AOI significantly affects σ_{φ_C} (see Figure 3). It was found that UAVSAR σ_{φ_C} increases of about 23% (29%) when moving from low to intermediate (intermediate to high) AOI. Nonetheless, to further confirm those results, RS and TSX SAR datasets were analyzed according to Table 2. In particular, RS (TSX) SAR scene ID 9 (11) and SAR scene ID 10 (14) were considered for lower and higher AOI, respectively. It must be underlined that both couples of SAR images were acquired under the same incident wavelength and they are characterized by almost the same SNR. When dealing with RS SAR scenes, an increase of about 25% in AOI results in an increase of about 60% in the σ_{φ_C} mean value, i.e., from 5° to 8° when AOI moves from about 32° to

about 40° (see Table 3). When dealing with TSX SAR scenes, the σ_{φ_C} mean value grows from about 9° to about 13° ($\approx 43\%$) when the AOI increases from about 26° to about 34° (31%) (see Table 3). The variability of σ_{φ_C} also increases when AOI increases.

When dealing with the effects of incident wavelength on σ_{φ_C} , according to Table 1, a fair analysis cannot be undertaken since each SAR platform is characterized by its NESZ. However, the L-band AP and C-band RS SAR imagery are considered since SAR scenes are characterized by the closest NESZ values within the dataset (≈ 6 dB difference). In addition, they were observed under almost the same AOI (see SAR scene ID 3–8 in Table 3). It can be observed, in Table 3, that, when moving from L- to C-band SAR images, σ_{φ_C} values increase, on average, of $\approx 24\%$ (from 3.1° to 3.85°), even though it must be pointed out that the mean SNR that characterizes L-band SAR measurements, i.e., 20.5 dB, is about 36% lower than the corresponding C-band one, i.e., 27.8 dB. Hence, although the larger mean SNR provided by C-band RS SAR measurements with respect to L-band ones would result in significantly lower σ_{φ_C} values (as discussed previously), an increasing trend of σ_{φ_C} values was observed when moving from L- to C-band SAR imagery, witnessing the key role played by the incident wavelength on the co-polarized phase difference distribution. In addition, the σ_{φ_C} variability, i.e., the standard deviation of σ_{φ_C} , also increases from L- to C-band of $\approx 42\%$. Results relevant to TSX SAR imagery listed in Table 3 also confirm the increasing trend of σ_{φ_C} when decreasing the incident wavelength. It can be observed how, under almost the same AOI, σ_{φ_C} values approximately doubled when moving from C- to X-band (see SAR scenes ID 7–8 and 11–12, and SAR scenes ID 9 and 14). Nonetheless, a completely different behavior applies for TSX SAR scene ID 13 that is characterized by the largest σ_{φ_C} mean and standard deviation values within the whole SAR dataset, i.e., $18.4^\circ \pm 16.5^\circ$. This is likely due to the fact that SAR scene ID 13 is severely corrupted by noise due to the high TSX NESZ (see Table 1): the average SNR lies very close to 3 dB, which is a threshold value usually adopted to judge the reliability of SAR measurements [24,42]).

6. Conclusions

A theoretical and experimental study aimed at investigating the sensitivity of σ_{φ_C} with respect to SAR acquisition parameters (NESZ, AOI, and incident wavelength) and meteo-marine conditions (WS and β) was performed. The X-Bragg polarimetric scattering model was adopted as a reference scattering framework to predict the behavior of σ_{φ_C} over sea surface, while experiments on actual measurements were accomplished considering a polarimetric SAR dataset that consists of 14 scenes collected over sea surface under different imaging configurations and environmental conditions. The main outcomes of this study are summarized as follows:

- The X-Bragg sea surface scattering model allows predicting the increasing trend of σ_{φ_C} with respect to AOI over sea surface along the whole range of Bragg scattering incidence angles, i.e., $\approx 20\text{--}60^\circ$.
- Under low-to-moderate sea state conditions, SAR imaging parameters have a stronger effect on σ_{φ_C} than meteo-marine parameters, which play a negligible role.
- Among SAR imaging parameters, incident wavelength and NESZ result in the most pronounced effect on sea surface σ_{φ_C} .

These outcomes can altogether support the design of polarimetric SAR architectures/algorithms that aim at enhancing the contrast between a given marine target of interest (ships, surfactants, icebergs, etc.) and sea clutter. Future works may include the extension of such sensitivity analysis on a larger SAR dataset (i.e., to include higher wind regimes) and the application of the proposed approach to find the most suitable SAR configuration to observe reference targets, e.g., ships.

Author Contributions: Conceptualization, A.B. and M.M.; Methodology, A.B. and F.N.; Software, A.B. and C.R.d.M.; Validation, C.R.d.M.; Formal Analysis, A.B. and F.N.; Investigation, A.B., C.R.d.M., F.N. and Domenico Velotto; Data Curation, C.R.d.M. and D.V.; Writing—Original Draft Preparation, A.B. and C.R.d.M.; Writing—Review and Editing, A.B. and F.N.; and Supervision, F.N., D.V. and M.M.

Funding: This study was partly funded by the Università degli Studi di Napoli Parthenope, project ID DING 202 and by the European Space Agency under the Dragon 4 project ID 32235.

Acknowledgments: The authors would like to thank the German Aerospace Center (DLR) that provided the TerraSAR-X SAR data under the AO OCE1045, NASA JPL that provided free of charge the UAVSAR SAR data and the wind field information under the PO.DAAC archive, the Japanese Space Agency (JAXA) that provided the ALOS PALSAR-1 SAR data under the RA-6 project ID 3064, and the Canadian Space Agency (CSA) that provided the RADARSAT-2 SAR data under the SOAR-EU project.

Conflicts of Interest: The authors declare no conflict of interest. The founding sponsors had no role in the design of the study; in the collection, analyses, or interpretation of data; in the writing of the manuscript, and in the decision to publish the results.

Abbreviations

The following abbreviations are used in this manuscript:

| | |
|----------|---|
| AOI | Angle of incidence |
| AP | ALOS PALSAR |
| ASCAT | Advanced Scatterometer |
| CCMP | Cross calibrated multi platform |
| dB | Decibel |
| ECMWF | European Center for Medium-Range Weather Forecasts |
| EUMETSAT | European Organization for the Exploitation of Meteorological Satellites |
| HH | Horizontal transmit-horizontal receive |
| KNMI | Royal Netherlands Meteorological Institute |
| ID | Identifier |
| MetOp | Meteorological operational satellite |
| NASA | National Aeronautics and Space Administration |
| NESZ | Noise-equivalent sigma zero |
| OSI SAF | Ocean and Sea Ice Satellite Application Facility |
| PO.DAAC | Physical Oceanography Distributed Active Archive Center |
| RMS | Remote Sensing Systems |
| ROI | Region of interest |
| RS | RADARSAT-2 |
| SAR | Synthetic aperture radar |
| SLC | Single-look complex |
| SNR | Signal-to-noise ratio |
| TSX | TerraSAR-X |
| VV | Vertical transmit-vertical receive |
| WS | Wind speed |

References

1. Costanza, R. The ecological, economic, and social importance of the oceans. *Ecol. Econ.* **1999**, *31*, 199–213. [[CrossRef](#)]
2. Visbeck, M. Ocean science research is key for a sustainable future. *Nat. Commun.* **2018**, *9*, 690. [[CrossRef](#)] [[PubMed](#)]
3. Fanning, L.; Mahon, R.; Baldwin, K.; Douglas, S. Transboundary waters assessment Programme (TWAP) assessment of governance arrangements for the ocean, Volume 1: Transboundary large marine ecosystems. *IOC Tech. Ser.* **2015**, *119*, 91.
4. European Commission. *Report on the Blue Growth Strategy Towards More Sustainable Growth and Jobs in the Blue Economy*; European Commission: Brussels, Belgium, 2017.
5. Nations, U. *The Sustainable Development Goals Report 2016*; United Nations: New York, NY, USA, 2016.
6. National Academies of Sciences, Engineering, and Medicine and others. Active earth remote sensing for ocean applications. In *A Strategy for Active Remote Sensing Amid Increased Demand for Radio Spectrum*; National Academies Press: Washington, DC, USA, 2015.

7. Ding, X.; Nunziata, F.; Li, X.; Migliaccio, M. Performance analysis and validation of waterline extraction approaches using single-and dual-polarimetric SAR data. *IEEE J. Sel. Top. Appl. Earth Observ. Remote Sens.* **2015**, *8*, 1019–1027. [[CrossRef](#)]
8. Nunziata, F.; Buono, A.; Migliaccio, M.; Benassai, G. Dual-polarimetric C-and X-band SAR data for coastline extraction. *IEEE J. Sel. Top. Appl. Earth Observ. Remote Sens.* **2016**, *9*, 4921–4928. [[CrossRef](#)]
9. Yu, Y.; Acton, S.T. Automated delineation of coastline from polarimetric SAR imagery. *Int. J. Remote Sens.* **2004**, *25*, 3423–3438. [[CrossRef](#)]
10. Velotto, D.; Nunziata, F.; Migliaccio, M.; Lehner, S. Dual-polarimetric TerraSAR-X SAR data for target at sea observation. *IEEE Geosci. Remote Sens. Lett.* **2013**, *10*, 1114–1118. [[CrossRef](#)]
11. Nunziata, F.; Migliaccio, M.; Brown, C.E. Reflection symmetry for polarimetric observation of man-made metallic targets at sea. *IEEE J. Ocean. Eng.* **2012**, *37*, 384–394. [[CrossRef](#)]
12. Li, H.; Perrie, W.; He, Y.; Lehner, S.; Brusch, S. Target detection on the ocean with the relative phase of compact polarimetry SAR. *IEEE Trans. Geosci. Remote Sens.* **2013**, *51*, 3299–3305. [[CrossRef](#)]
13. Migliaccio, M.; Nunziata, F. On the exploitation of polarimetric SAR data to map damping properties of the Deepwater Horizon oil spill. *Int. J. Remote Sens.* **2014**, *35*, 3499–3519. [[CrossRef](#)]
14. Migliaccio, M.; Nunziata, F.; Buono, A. SAR polarimetry for sea oil slick observation. *Int. J. Remote Sens.* **2015**, *36*, 3243–3273. [[CrossRef](#)]
15. Buono, A.; Nunziata, F.; Migliaccio, M.; Li, X. Polarimetric analysis of compact-polarimetry SAR architectures for sea oil slick observation. *IEEE Trans. Geosci. Remote Sens.* **2016**, *54*, 5862–5874. [[CrossRef](#)]
16. Velotto, D.; Migliaccio, M.; Nunziata, F.; Lehner, S. Dual-polarized TerraSAR-X data for oil-spill observation. *IEEE Trans. Geosci. Remote Sens.* **2011**, *49*, 4751–4762. [[CrossRef](#)]
17. Marino, A.; Dierking, W.; Wesche, C. A depolarization ratio anomaly detector to identify icebergs in sea ice using dual-polarization SAR images. *IEEE Trans. Geosci. Remote Sens.* **2016**, *54*, 5602–5615. [[CrossRef](#)]
18. Dabboor, M.; Geldsetzer, T. Towards sea ice classification using simulated RADARSAT Constellation Mission compact polarimetric SAR imagery. *Remote Sens. Environ.* **2014**, *140*, 189–195. [[CrossRef](#)]
19. Gill, J.P.; Yackel, J.J. Evaluation of C-band SAR polarimetric parameters for discrimination of first-year sea ice types. *Can. J. Remote Sens.* **2012**, *38*, 306–323. [[CrossRef](#)]
20. Touzi, R.; Charbonneau, F.; Hawkins, R.; Vachon, P. Ship detection and characterization using polarimetric SAR. *Can. J. Remote Sens.* **2004**, *30*, 552–559. [[CrossRef](#)]
21. Nunziata, F.; Buono, A.; Migliaccio, M. COSMO-SkyMed Synthetic Aperture Radar Data to Observe the Deepwater Horizon Oil Spill. *Sustainability* **2018**, *10*, 3599. [[CrossRef](#)]
22. Kudryavtsev, V.N.; Chapron, B.; Myasoedov, A.G.; Collard, F.; Johannessen, J.A. On dual co-polarized SAR measurements of the ocean surface. *IEEE Geosci. Remote Sens. Lett.* **2013**, *10*, 761–765. [[CrossRef](#)]
23. Liu, C.; Vachon, P.; Geling, G. Improved ship detection with airborne polarimetric SAR data. *Can. J. Remote Sens.* **2005**, *31*, 122–131. [[CrossRef](#)]
24. Minchew, B. Determining the mixing of oil and sea water using polarimetric synthetic aperture radar. *Geophys. Res. Lett.* **2012**, *39*. [[CrossRef](#)]
25. Marino, A.; Velotto, D.; Nunziata, F. Offshore Metallic Platforms Observation Using Dual-Polarimetric TS-X/TD-X Satellite Imagery: A Case Study in the Gulf of Mexico. *IEEE J. Sel. Top. Appl. Earth Obs. Remote Sens.* **2017**, *10*, 4376–4386. [[CrossRef](#)]
26. Ryu, E.; Yang, C.S.; Ouchi, K. Comparison of Ship Detection Accuracy Based on Image Contrast with Different Combinations HH-and VV-polarization TerraSAR-X SAR Images. *IEICE Tech. Rep.* **2017**, *117*, 13–16.
27. Salberg, A.B.; Rudjord, Ø.; Solberg, A.H.S. Oil spill detection in hybrid-polarimetric SAR images. *IEEE Trans. Geosci. Remote Sens.* **2014**, *52*, 6521–6533. [[CrossRef](#)]
28. Migliaccio, M.; Nunziata, F.; Gambardella, A. On the co-polarized phase difference for oil spill observation. *Int. J. Remote Sens.* **2009**, *30*, 1587–1602. [[CrossRef](#)]
29. Dierking, W.; Wesche, C. C-Band radar polarimetry—Useful for detection of icebergs in sea ice? *IEEE Trans. Geosci. Remote Sens.* **2014**, *52*, 25–37. [[CrossRef](#)]
30. Moen, M.A.; Doulgeris, A.P.; Anfinsen, S.N.; Renner, A.H.; Hughes, N.; Gerland, S.; Eltoft, T. Comparison of feature based segmentation of full polarimetric SAR satellite sea ice images with manually drawn ice charts. *Cryosphere* **2013**, *7*, 1693–1705. [[CrossRef](#)]
31. Joughin, I.R.; Winebrenner, D.P.; Percival, D.B. Probability density functions for multilook polarimetric signatures. *IEEE Trans. Geosci. Remote Sens.* **1994**, *32*, 562–574. [[CrossRef](#)]

32. Schuler, D.L.; Lee, J.S.; Hoppel, K.W. Polarimetric SAR image signatures of the ocean and Gulf Stream features. *IEEE Trans. Geosci. Remote Sens.* **1993**, *31*, 1210–1221. [[CrossRef](#)]
33. Lee, J.S.; Hoppel, K.W.; Mango, S.A.; Miller, A.R. Intensity and phase statistics of multilook polarimetric and interferometric SAR imagery. *IEEE Trans. Geosci. Remote Sens.* **1994**, *32*, 1017–1028.
34. Guissard, A. Phase calibration of polarimetric radars from slightly rough surfaces. *IEEE Trans. Geosci. Remote Sens.* **1994**, *32*, 712–715. [[CrossRef](#)]
35. Nunziata, F.; Gambardella, A.; Migliaccio, M. A unitary Mueller-based view of polarimetric SAR oil slick observation. *Int. J. Remote Sens.* **2012**, *33*, 6403–6425. [[CrossRef](#)]
36. Buono, A.; Nunziata, F.; de Macedo, C.R.; Velotto, D.; Migliaccio, M. A sensitivity analysis of the standard deviation of the co-polarized phase difference for sea oil slick observation. *IEEE Trans. Geosci. Remote Sens.* **2018**, doi:10.1109/TGRS.2018.2870738. [[CrossRef](#)]
37. Nunziata, F.; de Macedo, C.R.; Buono, A.; Velotto, D.; Migliaccio, M. On the analysis of a time series of X-band TerraSAR-X SAR imagery over oil seepages. *Int. J. Remote Sens.* **2018**, in print. [[CrossRef](#)]
38. Lee, J.S.; Pottier, E. *Polarimetric Radar Imaging: From Basics to Applications*; CRC Press: Boca Raton, FL, USA, 2009.
39. Yin, J.; Yang, J.; Zhou, Z.S.; Song, J. The extended Bragg scattering model-based method for ship and oil-spill observation using compact polarimetric SAR. *IEEE J. Sel. Top. Appl. Earth Observ. Remote Sens.* **2015**, *8*, 3760–3772. [[CrossRef](#)]
40. Buono, A.; Nunziata, F.; Migliaccio, M. Analysis of Full and Compact Polarimetric SAR Features over the Sea Surface. *IEEE Geosci. Remote Sens. Lett.* **2016**, *13*, 1527–1531. [[CrossRef](#)]
41. Hajnsek, I.; Pottier, E.; Cloude, S.R. Inversion of surface parameters from polarimetric SAR. *IEEE Trans. Geosci. Remote Sens.* **2003**, *41*, 727–744. [[CrossRef](#)]
42. Minchew, B.; Jones, C.E.; Holt, B. Polarimetric analysis of backscatter from the Deepwater Horizon oil spill using L-band synthetic aperture radar. *IEEE Trans. Geosci. Remote Sens.* **2012**, *50*, 3812–3830. [[CrossRef](#)]
43. EORC JAXA: ALOS Data Users Handbook Rev. C. 2008. Available online: https://www.eorc.jaxa.jp/ALOS/en/doc/fdata/ALOS_HB_RevC_EN.pdf (accessed on 5 October 2018).
44. Slade, B. *RADARSAT-2 Product Description*; Tech. Rep. RN-SP-S2-1238; MDA Ltd.: Richmond, BC, Canada, 2009.
45. Eineder, M.; Fritz, T. *TerraSAR-X Basic Product Specification Document*; Technical Report, TX-GS-DD-3302; 2010. Available online: https://www.dlr.de/dlr/Portaldata/1/Resources/documents/TX-GS-DD-3302_Basic-Product-Specification-Document_1_7.pdf (accessed on 18 October 2018).
46. Carvalho, D.; Rocha, A.; Gómez-Gesteira, M.; Alvarez, I.; Santos, C.S. Comparison between CCMP, QuikSCAT and buoy winds along the Iberian Peninsula coast. *Remote Sens. Environ.* **2013**, *137*, 173–183. [[CrossRef](#)]
47. Gao, G.; Wang, X.; Niu, M. Statistical modeling of the reflection symmetry metric for sea clutter in dual-polarimetric SAR data. *IEEE J. Ocean. Eng.* **2016**, *41*, 339–345.
48. Migliaccio, M.; Nunziata, F.; Montuori, A.; Paes, R.L. Single-look complex COSMO-SkyMed SAR data to observe metallic targets at sea. *IEEE J. Sel. Top. Appl. Earth Observ. Remote Sens.* **2012**, *5*, 893–901. [[CrossRef](#)]
49. Atlas, R.; Hoffman, R.N.; Ardizzone, J.; Leidner, S.M.; Jusem, J.C.; Smith, D.K.; Gombos, D. A cross-calibrated, multiplatform ocean surface wind velocity product for meteorological and oceanographic applications. *Bull. Am. Met. Soc.* **2011**, *92*, 157–174. [[CrossRef](#)]
50. Verhoef, A.; Stoffelen, A. ASCAT Wind Product User Manual Version 1.15. Document external project: SAF/OSI/CDOP/KNMI/TEC/MA/126; 2018, EUMETSAT. Available online: http://projects.knmi.nl/publications/fulltexts/ss3_pm_ascat_1.15.pdf (accessed on 18 October 2018).
51. Mattia, F.; Le Toan, T.; Souyris, J.C.; De Carolis, C.; Floury, N.; Posa, F.; Pasquariello, N. The effect of surface roughness on multifrequency polarimetric SAR data. *IEEE Trans. Geosci. Remote Sens.* **1997**, *35*, 954–966. [[CrossRef](#)]
52. Younis, M.; Huber, S.; Patyuchenko, A.; Bordoni, F.; Krieger, G. Performance Comparison of Reflector- and Planar-Antenna Based Digital Beam-Forming SAR. *IEEE J. Antennas Propag.* **2009**, *2009*, 614931. [[CrossRef](#)]

