

Article

Remote Sensing and GIS in Landslide Management: An Example from the Kravarsko Area, Croatia

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Abstract: The Kravarsko area is located in a hilly region of northern Croatia, where numerous landslides endanger and damage houses, roads, water systems, and power lines. Nevertheless, natural hazard management plans are practically non-existent. Therefore, during the initial research, a landslide inventory was developed for the Kravarsko pilot area based on remote sensing data (high-resolution digital elevation models), and some of the landslides were investigated in detail. However, due to the complexity and vulnerability of the area, additional zoning of landslide-susceptible areas was needed. As a result, a slope gradient map, a map of engineering geological units, and a land-cover map were developed as inputs for the landslide susceptibility map. Additionally, based on the available data and a landslide inventory, a terrain stability map was developed for landslide management. Analysis and map development were performed within a geographical information system environment, and the terrain stability map with key infrastructure data was determined to be the “most user-friendly and practically usable” resource for non-expert users in natural hazard management, for example, the local administration. At the same time, the terrain stability map can easily provide practical information for the local community and population about the expected landslide “risk” depending on the location of infrastructure, estates, or objects of interest or for the purposes of future planning.

Keywords: remote sensing (RS); landslide inventory (LI); geographic information system (GIS); landslide susceptibility map (LSM); terrain stability map (TSM); natural hazard management (NHM)



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1. Introduction

Natural hazards cover a wide range of phenomena that might have a negative effect on living and artificial things. Within the category of natural hazards, geohazards can occur as a result of active geologic processes or human activity [1]. Landslides are geohazards that frequently endanger safety and damage property in the ever-changing environment [2,3]. Landslides are usually described as mass movements [4–6] and are divided into categories according to the mechanical behavior of the movement [7–9]. Both natural and human factors contribute to the occurrence of landslides, and understanding these interactions is crucial for effective landslide risk management. Natural causes include seismic activities such as earthquakes [10], heavy rainfall [11], and volcanic eruptions [12], while human activities such as mining [13], construction [14], deforestation [15], and changes in land use [16] can significantly alter the landscape, making it more susceptible to landslides. One of the first steps in landslide research in a specific area is the development of a reliable landslide inventory [17–19]. A reliable landslide inventory of the area of the interest (at the local, regional, or state level) is a basis for successful landslide mitigation processes and assessments of susceptibility, hazards, and risk (which also depend on the other available data) [20–22]. On landslide inventory maps, the spatial distribution of landslides is presented with additional information [17,19,23] while, on landslide susceptibility maps, zones prone to landslides are distinguished [17,21,24,25]. As basic as they may be, landslide inventories and landslide susceptibility maps are of great value in planning and natural hazard (i.e., landslide) management [2,3,17,26,27].

Conventional landslide inventory development includes field mapping and remote sensing, i.e., the analysis of stereoscopic aerial photography [19,28]. This approach has its limitations, as the accessibility of the terrain can vary or vegetation cover can mask features in the photos [28,29]. However, with the rapid development of another remote sensing technique, light detection and ranging (LiDAR), the collection of reliable, accurate, and high-resolution terrain data has been enabled [30,31]. As a result of LiDAR point cloud data analysis, the development of high-resolution digital elevation models (hrDEMs) has been enabled. One such hrDEM with proven applicability in landslide research is the bare-ground model/digital terrain model (DTM) [19,22,32]. The main advantages of the use of this remote sensing product (hrDTM) are that there is no inaccessible terrain area, and land cover can be removed, providing a reliable and accurate basis for landslide feature mapping [31,32]. Nevertheless, in the development of a landslide inventory based on LiDAR data for any area, the field mapping aspect should not be neglected [19,32,33]. Landslide susceptibility mapping involves analyzing and mapping the spatial distribution of factors that contribute to landslide susceptibility, with the main goal being to identify areas that are prone to landslides, providing valuable information for land-use planning, risk assessment, and hazard mitigation [20,24,33]. In the development of landslide susceptibility maps, much depends on the available data and their quality and resolution [20,24,33]. In overcoming this obstacle (data availability and quality), different approaches are used (heuristic, probabilistic, or deterministic) [18,21,34]. Based on a review of articles employing diverse quantitative–statistical, multicriteria decision-making, and machine learning methods for landslide susceptibility mapping, Pourghasemi et al. [35] observed that logistic regression is the predominant technique, followed by frequency ratio and weights of evidence, while the slope gradient is considered the most important conditioning factor in landslide occurrence. In recent years, a variety of machine learning algorithms have emerged as essential tools in the field of landslide susceptibility mapping [36–39]. Regardless of the approach used, the developed landslide susceptibility map has its value, but the “final user” has to be informed about its limitations: the data analyzed, the methodology used, and its scale applicability [20,21].

In the Kravarsko settlement area (research area) in northern Croatia, landslides are common [32,40], and damage to private and public properties occurs (houses and infrastructure) [26,32,40,41]. However, the first acquisition of high-quality LiDAR data for the Kravarsko area occurred just five years ago in the spring of 2018; consequently, a landslide inventory was developed just three years ago in 2020 [32]. During data analysis, field work on the landslide inventory, and specific landslide research in the area, it became obvious that site-specific geological and anthropogenic conditions are important for landslide activation and development in Kravarsko, along with ongoing climate change [26,32,41]. The initial landslide susceptibility map for the Kravarsko area was developed in 2021, but it has not been published [42]. As landslides continue to occur or reactivate in the Kravarsko area and landslide data (the inventory and susceptibility map) provide important information for community mitigation plans and management, the aim of this paper is to present the perennial research results simply, reliably, and in a usable format, “all in one place”, with additional analysis, comments, and recommendations and a new landslide management map, i.e., a terrain stability map including key infrastructure.

2. Materials and Methods

2.1. Study Area

The Kravarsko pilot area (PA \approx 61.7 km²) is in the Vukomeričke Gorice hilly area, ~25 km south of Zagreb in northern Croatia, as shown in Figure 1a. The geological development and structure of the investigated area in Vukomeričke Gorice are described in detail in [32]. Based on those data, it is clear that for PA, there are three main engineering geological units present on the area: alluvial sediments (the youngest Quaternary sediments from the Holocene, represented by gravel, sands, and clays), loess-type sediments (Quaternary sediments from the Pleistocene, represented by silts, sands, and clays), and the

“problematic” informal lithostratigraphic unit—the Vrbova fm. (formation). The Vrbova fm. constitutes Tertiary Pliocene sediments represented by sands, clays, and gravels; in this unit, there are multiple records of landslides on the available 1:100,000- and 1:500,000-scale maps [43,44], as shown in Figure 1b. Regardless, the lack of available detailed landslide data and information for the wider Kravarsko area is surprising, as in 2014, for example, a natural disaster was declared for the area due to landslide phenomena [40]. Within Kravarsko PA, a smaller field validation area (VA $\approx 10.3 \text{ km}^2$, $\approx 17\%$ of the investigated PA) was defined, in accordance with recommendations from the literature [19,45], and the VA also included the Kravarsko settlement area (Figure 1b).

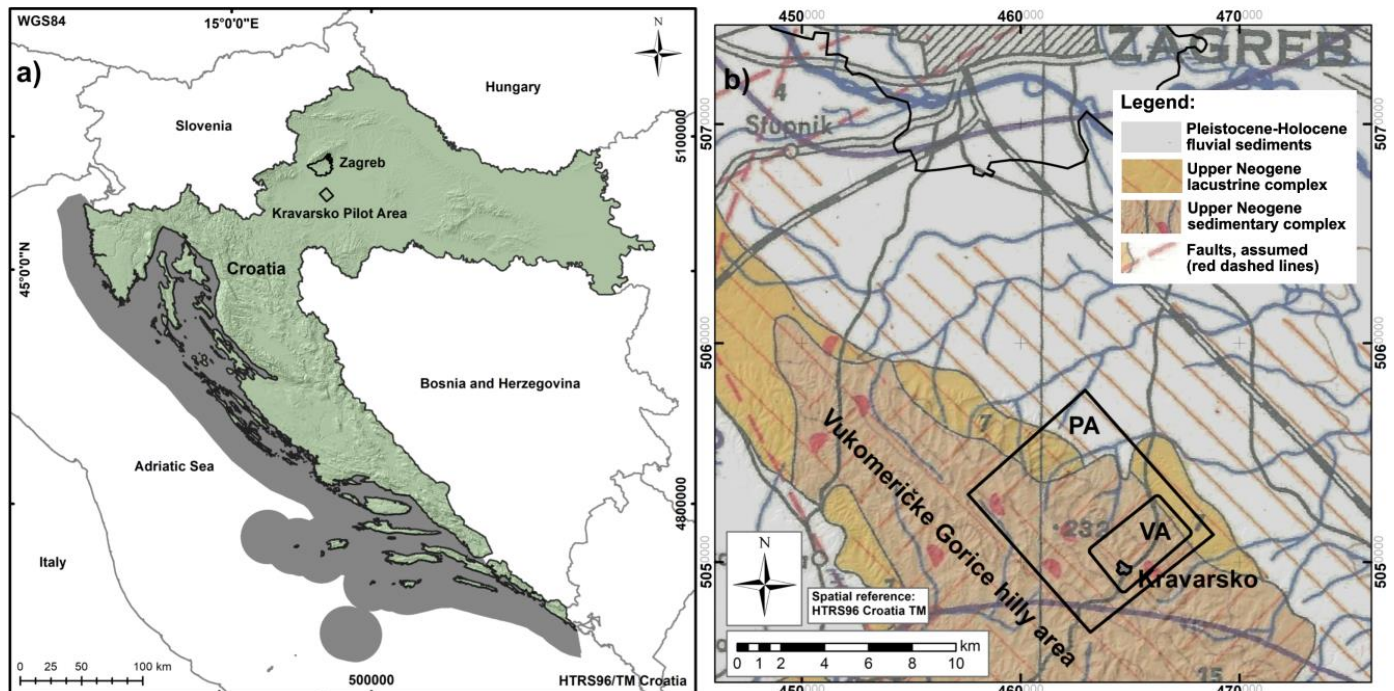


Figure 1. Kravarsko pilot area location: (a) in northern Croatia, south of Zagreb city; (b) in the Vukomeričke Gorice hilly area, which is prone to landslides. In addition to the pilot area (PA $\approx 61.7 \text{ km}^2$), the field validation area (VA $\approx 10.3 \text{ km}^2$) and the Kravarsko settlement area are marked. As a base map, a detail from the engineering geological map of Yugoslavia at the (original) scale of 1:500,000 is used, and even in this map (scale), landslides are identified in the (wider) area (red polygons). On the map, three units are present in the PA: (i) the Upper Neogene sedimentary complex—sandstones, marly clays, marls, and sands prone to erosion and sliding (brown on map); (ii) the Upper Neogene lacustrine complex—sands, gravels, and clays prone to erosion and sliding (yellow on map); and (iii) Pleistocene–Holocene fluvial sediments—sandy gravels that are sporadically clayey, mostly covered with loam, and porous (white on map). Assumed faults (wider area) are marked with red dashed lines [44].

2.2. Site-Specific Geohazard Mapping

Geohazards [1,46–48] and landslides [6–9] can be classified. Based on the available landslide data [26,32,41], the most common geohazards in PA are slides in soil that can be described as rotational, planar, or compound slides in clays and silts or as debris slides in gravels and sands [49] (see Figure 2a–h). From here on, these soil slides and other mass movements in the PA will be referred to simply as landslides.

The available (historical) geological and remote sensing data (historical and new) were analyzed, as described in detail in [26,32,41]. Based on recent remote sensing data (20 points per m^2 airborne LiDAR scan data from early spring 2018), high-resolution DEMs with a cell size of $0.5 \times 0.5 \text{ m}$ were developed: a digital surface model (DSM), a digital terrain model (DTM), and a hillshade digital terrain model (DTMh). The development of a

reliable landslide inventory is possible based on these hrDEMs if distinct landslide features can be mapped on them [19,22,50] while the mapping criteria remain the same [31,32,51].

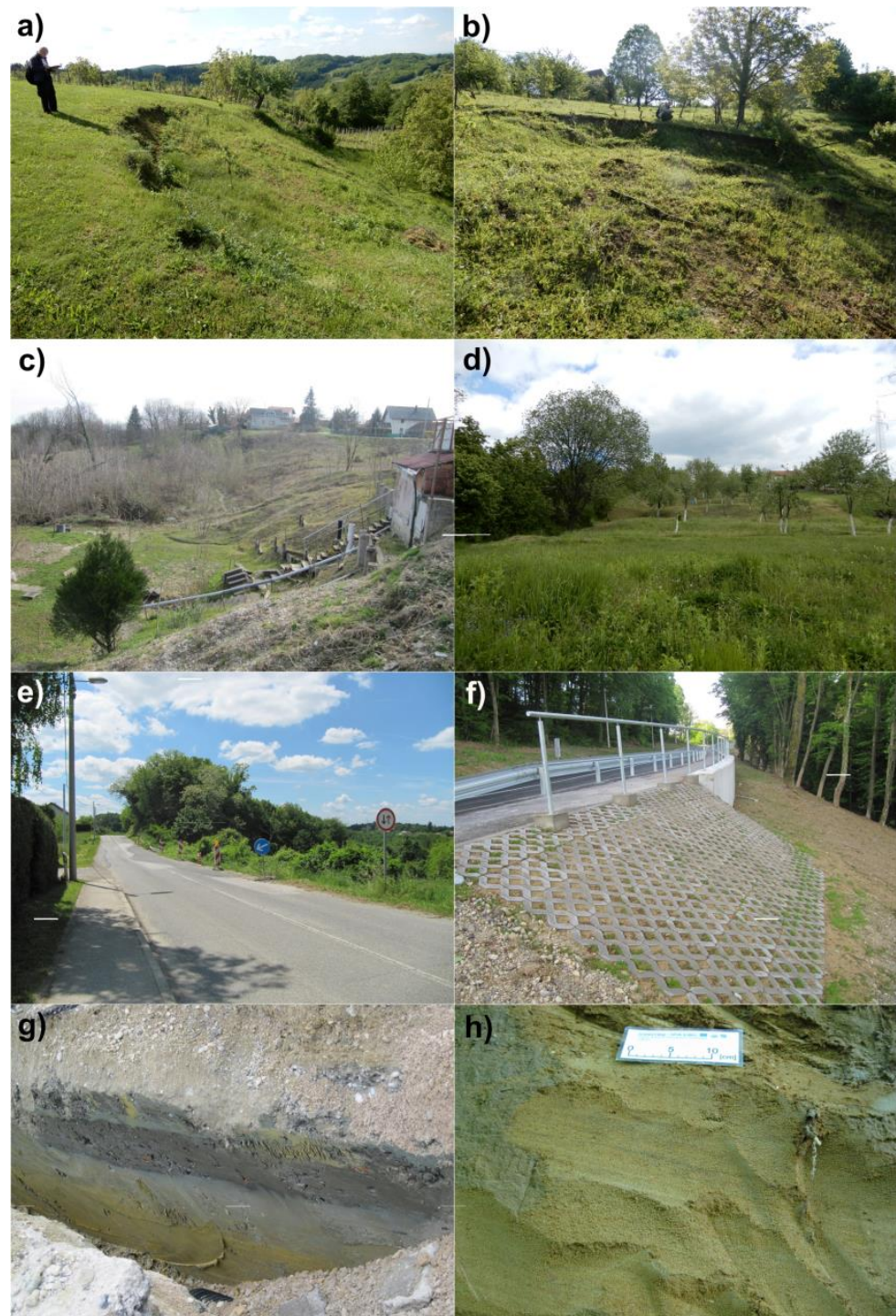


Figure 2. Kravarsko pilot area field examples: (a) landslide head scarp; (b) landslide flank; (c) landslide damaging houses and road; (d) cultivated orchard in an old landslide body; (e) road damaged by landslide; (f) road mitigation measures; (g) Vrbova fm. clays; (h) Vrbova fm. sands.

2.3. Developed Landslide Inventory

For the Kravarsko PA, a landslide inventory was developed based on high-resolution remote sensing data (hrDEMs) [32]. For the remotely mapped landslides (on hrDEMs), a scoring system was established based on the confidence of the identification of the ge-

omorphological indicators and landslide features [32]. Essentially, if landslide features (head scarp, flanks, toe parts, and internal deformations) are present and can be mapped, the score is higher, and the landslide is mapped with greater confidence, as described in detail in [32]. At the same time, a low-score landslide can indicate an older landslide with some of the features degraded and hardly visible, even on hrDEMs. Therefore, the “state” of the landslide feature (visibility, clarity, and the possibility of identification) can also be considered an indicator of landslide activity and “age”: old, dormant, or active [52]. The landslide inventory was developed within the geographical information system environment (ArcGIS) and consists of 1430 mapped landslides for the Kravarsko PA with the following (simplified) classifications: 595 old landslides (42% of the mapped landslides), 648 dormant landslides (45% of the mapped landslides), and 187 active landslides (13% of the mapped landslides) (Figure 3). The total area (sum) of the mapped landslides is 6.5 km² (11% of the PA), with 3.2 km² of old landslides (49% of the mapped landslide area), 2.8 km² of dormant landslides (43% of the mapped landslide area), and 0.5 km² of active landslides (8% of the mapped landslide area). The developed landslide inventory was validated in the field on a smaller validation area (Figure 3), producing the following main conclusions: (i) landslide mapping is easier, quicker, and even more reliable on a detailed hrDEM due to the vegetation cover and gullies present in the field; (ii) shallow and slow slope movements (soil creeps) are difficult to map precisely in the field; and (iii) hrDEMs are a powerful tool in the development of reliable landslide inventories, but field mapping cannot be neglected, as it is essential for the calibration and fine tuning of the “scoring system” in order to reflect the specific aspects of the investigated area (geological, geomorphological, anthropogenic, etc. (more details are provided in [32])).

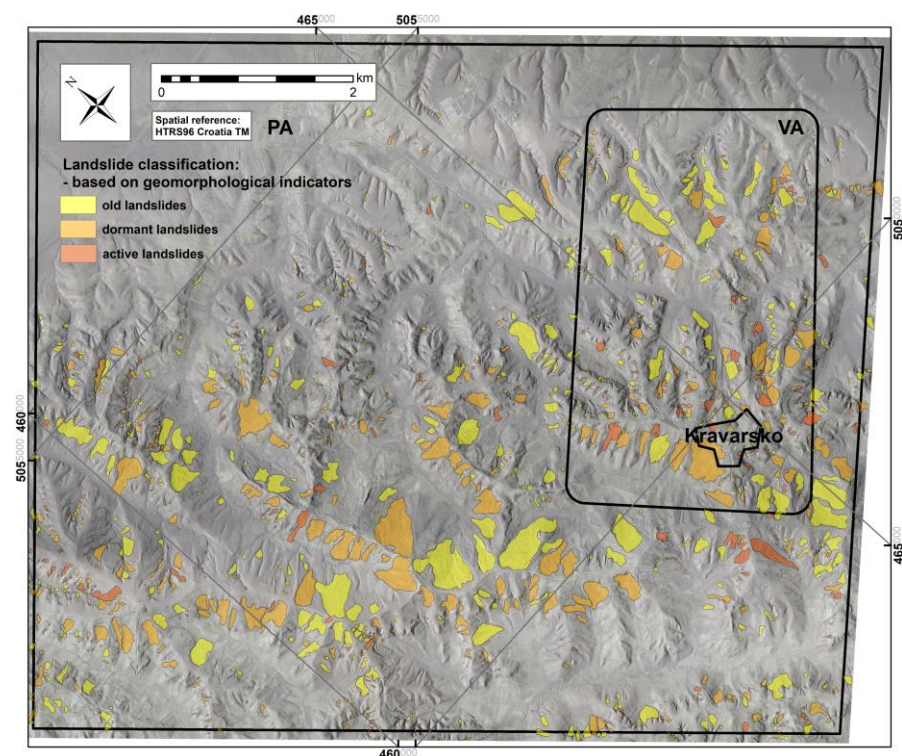


Figure 3. Kravarsko pilot area landslide inventory map with landslide classifications based on geomorphological indicators. In addition to the pilot area (PA \approx 61.7 km²), a field validation area (VA \approx 10.3 km²) and the Kravarsko settlement area are marked with, the high-resolution hillshade digital terrain model (hrDTMh) used as the base layer.

3. Results

For the Kravarsko area, the existing landslide susceptibility maps were on national (Republic of Croatia, [53]) and regional scales (Zagreb County, [54]). These maps are small-scale maps. The landslide susceptibility map for the Kravarsko PA presented herein was

based on a heuristic approach, but the developed landslide inventory for the area was also reviewed, and field validation was performed on a smaller validation area; additionally, a terrain stability map was developed. For the development of the landslide susceptibility map, the following datasets were analyzed: a detailed slope gradient map (derived from detailed LIDAR data: hrDEMs with a 0.5×0.5 m pixel size), an engineering geological unit map based on the available geological data (mostly 1:100,000 scale), and land-cover data based on detailed orthophotos from 2018 (10×10 cm pixel size). All analysis was performed within the GIS environment with 1×1 m pixel-size raster format data for each layer: slope gradient data, engineering geological units, and land-cover data. The number of classes for each input layer set was based on the available dataset's "quality" and "importance". For the terrain stability map, the developed landslide inventory and landslide susceptibility map were analyzed, and the key infrastructure data were reviewed.

3.1. Slope Gradient Data

When developing a landslide susceptibility map, usually, the slope gradient data are regarded as the "most important" [53,55,56]. For the Kravarsko PA, based on the hrDEMs (0.5×0.5 m), it was possible to develop a detailed slope gradient data/slope angle map with 1° distinction (Figure 4). The high-resolution input was reflected in the number of classes (9) used in the initial analysis for this layer, and, at the same time, this layer was regarded as the "most important" in the analysis (with an overall weight of 0.5). For the Kravarsko PA (≈ 61.7 km²), the classification of detailed slope gradient data with percentages and areas is summarized as follows: flat areas (with no inclination, 0°) were present in $\sim 3\%$ of the area (~ 1.9 km² of the PA), almost flat areas (with an inclination in the range of $1\text{--}5^\circ$) were present in $\sim 34\%$ of the area (~ 21.0 km² of the PA), gentle slopes (with an inclination in the range of $6\text{--}15^\circ$) were present in $\sim 45\%$ of the area (~ 27.8 km² of the PA), slopes (with an inclination in the range of $16\text{--}25^\circ$) were present in $\sim 13\%$ of the area (~ 8.0 km² of the PA), inclined slopes (with an inclination in the range of $26\text{--}35^\circ$) were present in $\sim 4\%$ of the area (~ 2.5 km² of the PA), steep slopes (with an inclination in the range of $36\text{--}45^\circ$) were present in $\sim 1\%$ of the area (~ 0.6 km² of the PA), and very steep slopes (with an inclination in the ranges of $46\text{--}60^\circ$, $61\text{--}75^\circ$, and $76\text{--}90^\circ$) were present in less than $\sim 0.2\%$ of the area (~ 0.1 km² of the PA), as shown in Figure 4.

3.2. Engineering Geological Units

In landslide development in some areas (depending on the landslide susceptibility of the area), the geological setting (in the wider sense) plays an important role [57–59]. For the Kravarsko PA, the most detailed available geological data are at the scale of 1:100,000 [31], which is a relatively "coarse" scale compared to the resolution of available hrDEMs. Nevertheless, for the PA, three major engineering geological units (EGUs) can be differentiated based on the available geological data: alluvial sediments, represented by gravels (G), sands (S), and clays (C); loess-type sediments, represented by silts (M), sands (S), and clays (C); and the Vrbova fm., represented by sands (S), clays (C), and gravels (G). The vast majority of the landslides in the PA occur in Vrbova fm. [26,32,40,41], and the developed map of EG units reflects this (Figure 5). As the geological input data (in the wider sense) were relatively coarse, "just" three classes for this layer were used in the analysis. At the same time, however, this layer was regarded as "almost as important as" the slope gradient data layer (with an overall weight of 0.4). For the Kravarsko PA (≈ 61.7 km²), the reclassified engineering geological data with percentages are summarized as follows: alluvial sediments were present in $\sim 3\%$ of the area (~ 1.9 km² of the PA), loess-type sediments were present in $\sim 15\%$ of the area (~ 9.3 km² of the PA), and Vrbova fm. was present in $\sim 82\%$ of the area (~ 50.6 km² of the PA), as detailed in Figure 5.

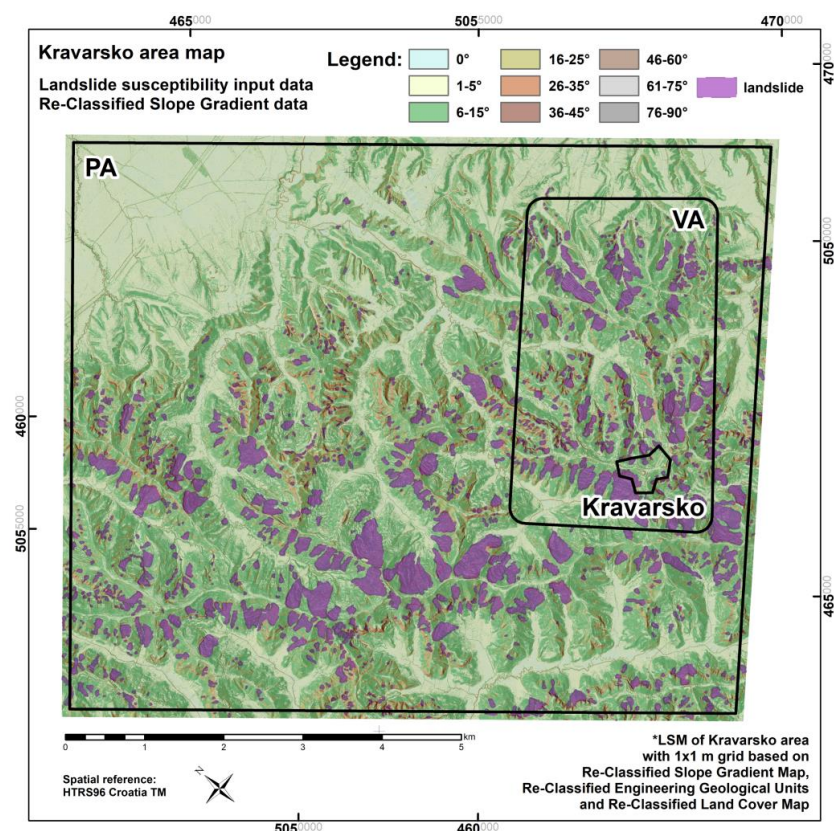


Figure 4. Kravarsko pilot area slope gradient data: slope angle map with nine classes overlaid with landslide inventory data (developed within the ArcGIS environment).

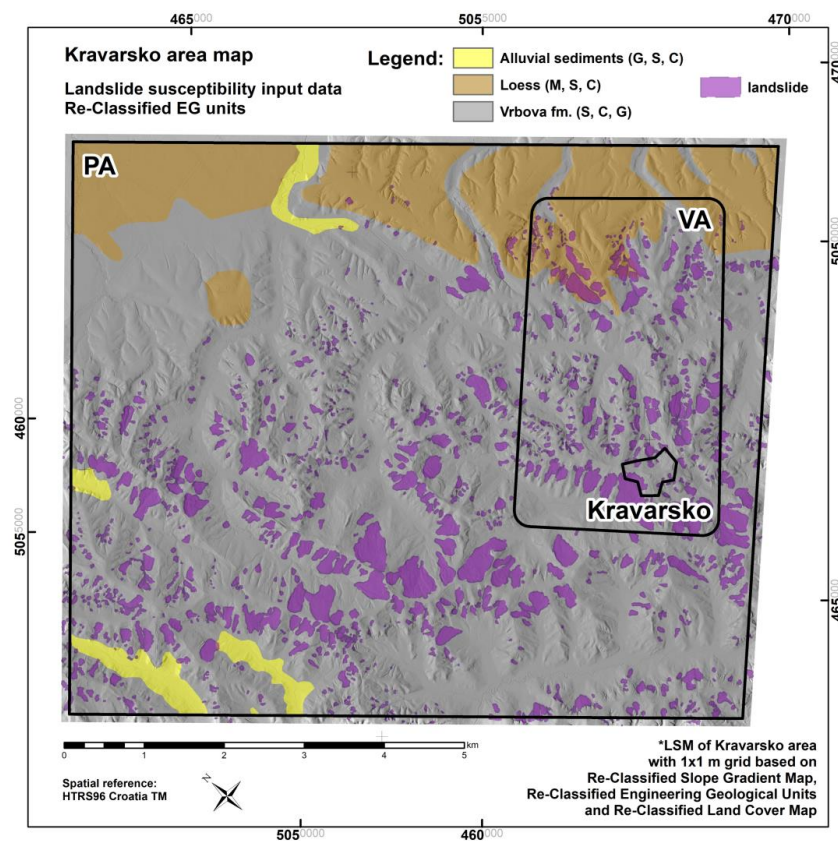


Figure 5. Kravarsko pilot area engineering geological unit map with three classes overlaid with landslide inventory data (developed within the ArcGIS environment).

3.3. Land-Cover Data

In landslide development in some areas (depending on the landslide susceptibility of the area), the land-cover type can play some role [60–62]. For the Kravarsko PA, new land-cover data (map) were developed based on detailed orthophotos from 2018 (10 × 10 cm pixel size), and four classes of land cover were differentiated: surface water areas, forest areas (or areas with “high” vegetation cover), urban areas (or areas with “artificial” cover), and fields (or areas with “low” or no vegetation) (see Figure 6). In our analysis, the land-cover data were regarded as the “least important” layer (with an overall weight of 0.1). For the Kravarsko PA ($\approx 61.7 \text{ km}^2$), the detailed land-cover data reclassified by classes with percentages and areas are summarized as follows: surface water areas were present in less than 0.1% of the area ($\sim 0.1 \text{ km}^2$ of the PA), forest areas were present in $\sim 59\%$ of the area ($\sim 36.4 \text{ km}^2$ of the PA), urban areas were present in $\sim 6\%$ of the area ($\sim 3.7 \text{ km}^2$ of the PA), and field areas were present in $\sim 35\%$ of the area ($\sim 21.6 \text{ km}^2$ of the PA), as shown in Figure 6.

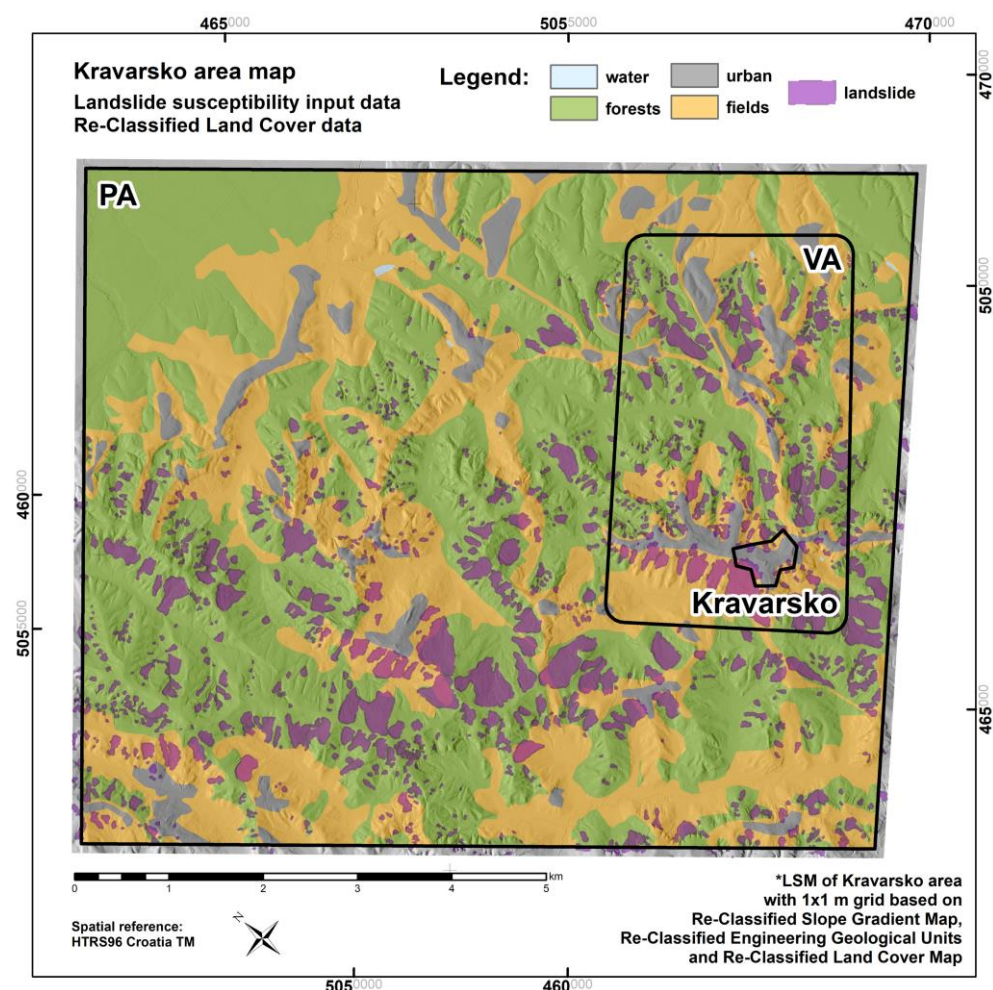


Figure 6. Kravarsko pilot area land-cover data map with four classes overlaid with landslide inventory data (developed within the ArcGIS environment). The land-cover map was developed based on a detailed orthophoto with a 10 × 10 cm pixel size.

3.4. Development of the Kravarsko Landslide Susceptibility Map

Landslide susceptibility maps are classified according to the methodology and data used during their development: they can be quantitative or qualitative; direct or indirect; or heuristic, probabilistic, or deterministic, with differentiated or undifferentiated types of landslide mechanisms [63–67]. For the Kravarsko PA, the landslide susceptibility map was developed using a heuristic approach with the weight factors presented in Table 1.

The heuristic approach was chosen based on the available data and experience, while the weight of the criteria for the layers, as well the overall weight (Table 1), was based on recommendations from the literature [55,62] and expert judgment [53]. It is important to mention that for Croatia, a small-scale LSM was developed in 2015, and those results were also taken into account [53]. Still, as extensive field work was carried out for the Kravarsko PA (described in detail in [32], with the material properties described in detail in [26]), the weights of criteria used herein were modified based on those findings and engineering judgments to more accurately reflect the terrain conditions. A detailed slope gradient map and land-cover data were developed as completely new datasets, while the available geological data were used to develop reclassified engineering geological units. Based on these three “basic” layers (slope angle, engineering geological units, and land cover as 1×1 m pixel-sized raster format data), a landslide susceptibility map was developed. The landslide inventory was used to check the developed landslide susceptibility map, with field validation performed on a smaller validation area (Figure 7).

Table 1. Weights of criteria and classes used for the analysis: layer-specific and overall.

Landslide Susceptibility Map Layer	Classes	Weight of Criteria (Layer)	Weight of Criteria (Overall)
Slope angle (°)	0	1	0.5
	1–5	3	
	6–15	6	
	16–25	9	
	26–35	12	
	36–45	15	
	46–60	17	
	61–75	18	
	76–90	19	
Engineering geological units	Alluvial sediments	5	0.4
	Loess	20	
	Vrbova fm.	75	
Land cover	Water	0	0.1
	Forest	25	
	Urban	35	
	Fields	40	

On the resulting LSM map, five classes were distinguished (Table 2); whereas water areas can be disregarded as “landslide-free areas” ($\sim 0.1\%$ and ~ 0.1 km² of the PA area), the low-LSM zone can be considered a “not or somewhat landslide-prone area” ($\sim 31.4\%$ and ~ 19.1 km² of the PA area). Meanwhile, medium, high, and very-high LSM zones can be considered “landslide-prone areas” ($\sim 68.5\%$ and ~ 42.6 km² of the PA area), as shown in Figure 7. It should be noted that landslides can occur in “not landslide-prone areas”, but their occurrence is not expected to be as frequent and devastating as in “landslide-prone areas”.

Table 2. Landslide susceptibility map classes, percentages, and zones.

Resulting LSM Class	Percentage (%)	Area (km ²)	Resulting LSM Zone
0	0.1	0.1	Water
1	31.4	19.1	Low
2	35.9	22.2	Medium
3	27.2	16.7	High
4	5.4	3.7	Very high

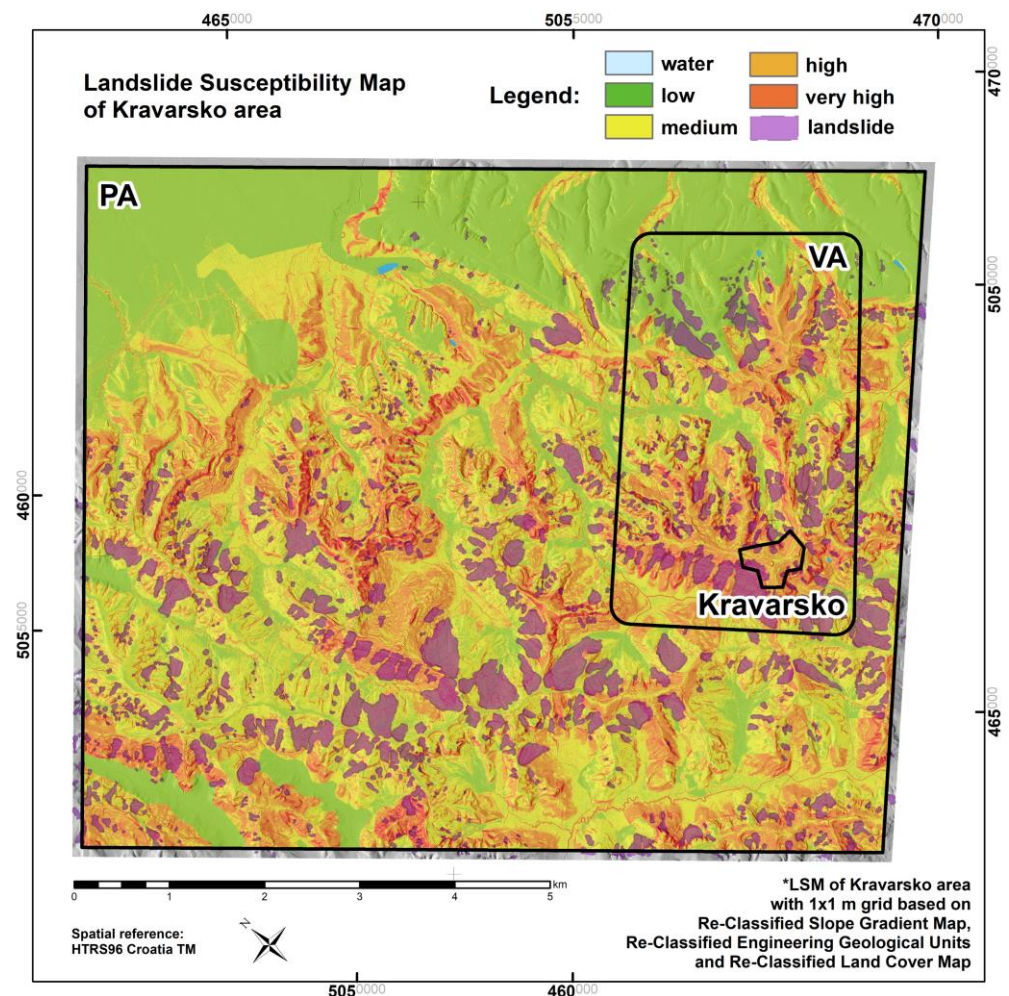


Figure 7. Kravarsko pilot area landslide susceptibility map with five classes overlaid with landslide inventory data (developed within the ArcGIS environment).

3.5. Development of the Kravarsko Area Terrain Stability Map

For the Kravarsko PA, the main goal was to develop a simple and usable map for natural hazard management (NHM) for the local community (i.e., non-expert users) [32,61,68–71]. It was concluded that for this purpose, a somewhat modified landslide susceptibility map (LSM) with additional content of interest (anthropogenic structures) could prove more useful: this would result in a terrain stability map (TSM) with key infrastructure data. The results of the previous research were taken into account [19,25,29], along with the data and analyses presented herein, and the Kravarsko Area Terrain Stability Map was developed, as shown in Figure 8. On this map, the Kravarsko PA is differentiated into three zones of terrain stability: stable zones ($\sim 7.3 \text{ km}^2$ and $\sim 12\%$ of the PA area) are areas without recorded landslides (0 landslides), possibly unstable zones ($\sim 23.0 \text{ km}^2$ and $\sim 37\%$ of the PA area) are areas with some (in general, smaller) recorded landslides (131 landslides from inventory, corresponding to $\sim 9\%$ of landslides and $\sim 0.11 \text{ km}^2$ of the landslide area), and unstable zones ($\sim 31.4 \text{ km}^2$ and $\sim 51\%$ of the PA area) are areas with recorded landslides (1299 landslides from inventory, corresponding to $\sim 91\%$ of landslides and $\sim 6.40 \text{ km}^2$ of landslide area). For this type of zoning, reliable landslide inventory data are vital. In natural hazard management (NHM), anthropogenic structures and their possible endangerment are important aspects of future planning and funding allocation, so the available digital data regarding settlements, roads, power lines, and water systems were taken into account in relation to the defined zones of terrain stability (see Table 3).

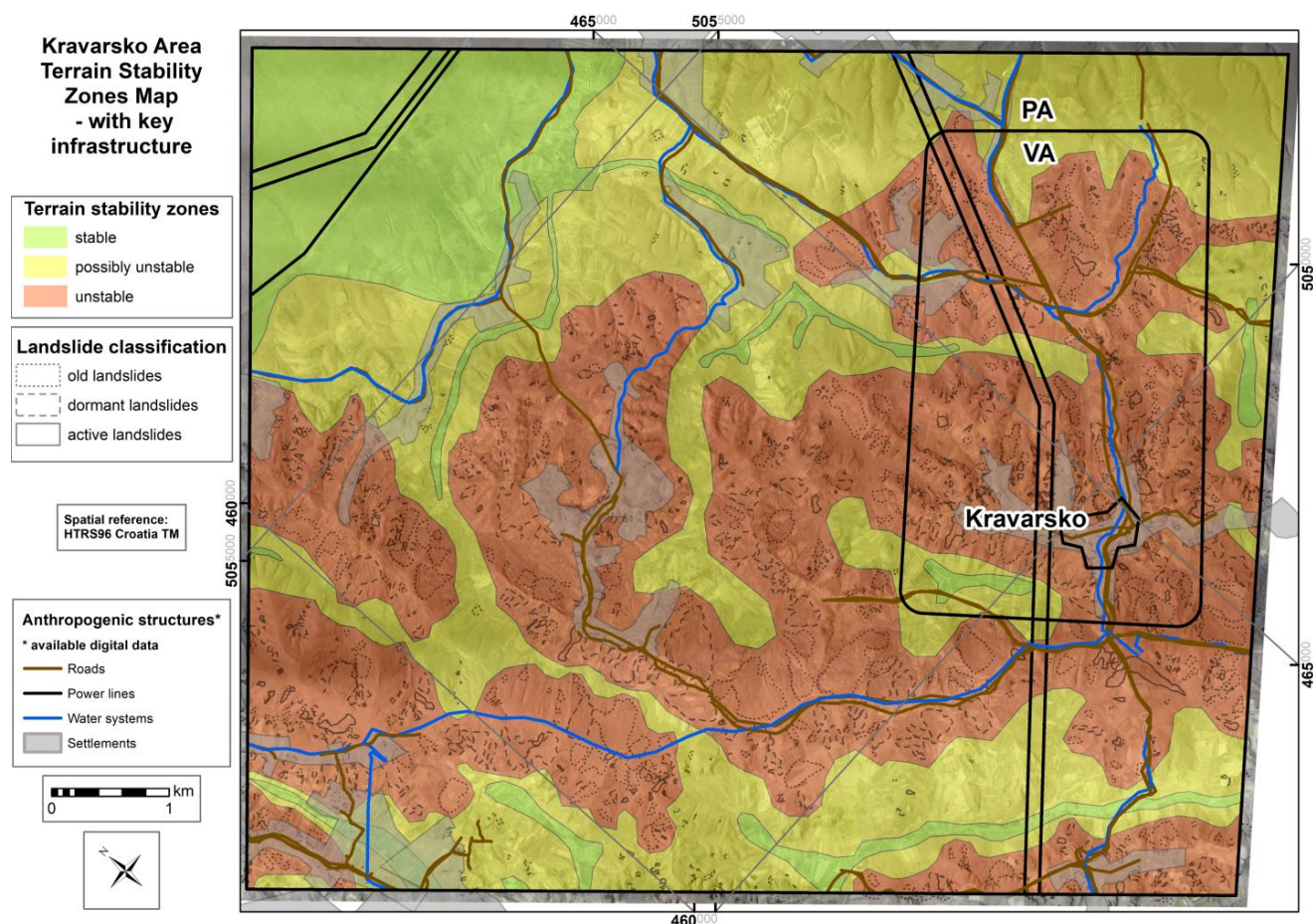


Figure 8. Kravarsko Pilot Area Terrain Stability Zone Map with key infrastructure and three classes of terrain stability overlaid with the landslide inventory data (developed within the ArcGIS environment).

Table 3. Anthropogenic structures: settlements (sett.), roads, water systems (w.s.), power lines (p.l.), and zones of terrain stability (stable, possibly unstable, and unstable zones).

Anthropogenic Structures	Kravarsko PA (~61.7 km ²)	Stable Zone (~7.3 km ² of PA)	Possibly Unstable Zone (~23.0 km ² of PA)	Unstable Zone (~31.4 km ² of PA)
Settlements (km ²)	5.69 (~9% of PA)	0.29 (~5% of sett.)	2.13 (~37% of sett.)	3.27 (~58% of sett.)
Roads (km ²)	85.42 (100%)	4.10 (~5% of roads)	24.24 (~28% of roads)	57.08 (~67% of roads)
Water systems (km ²)	34.65 (100%)	2.38 (~7% of w.s.)	12.91 (~37% of w.s.)	19.36 (~56% of w.s.)
Power lines (km ²)	21.72 (100%)	7.12 (~33% of p.l.)	5.55 (~25% of p.l.)	9.05 (~42% of p.l.)

4. Discussion

In the process of developing the Kravarsko landslide susceptibility map (LSM), the input layers were “treated” with different weight factors and numbers of classes in order to find the optimum balance between data quality and the intention that the LSM reflect the terrain conditions for landslide occurrences as closely as possible. In order to achieve this goal, the results of previous detailed research studies were also taken into account [26,32,41]. The presented LSM for the Kravarsko PA is considered acceptable and usable regarding the input data and field observations [26,32]. However, after detailed landslide inventory analysis, recommendations can still be made about input layers and cases when the landslide susceptibility map is optimal and cases when a “new” map (a terrain stability map with key infrastructure data) is more appropriate for the local community to use in natural hazard management (NHM).

4.1. Comments on Input Layers

For the Kravarsko PA, detailed slope gradient data with a 1° distinction and nine classes were developed; however, the steep and very steep slope classes (with inclinations in the range of $36\text{--}90^\circ$) were present only in $\sim 1.2\%$ of the area ($\sim 0.7\text{ km}^2$ of PA), while flat areas (with no inclination, 0°) were present in $\sim 3\%$ of the area ($\sim 1.9\text{ km}^2$ of PA) (see Figure 4). This was expected, as the “usual and most common” natural terrain has some inclination to it, and it does not comprise flat areas or very steep slopes (as steep slopes “tend” to become gentler due to different natural processes: erosion, mass movements, etc.). This finding was also reflected in the landslide frequency for these areas: $\sim 2\%$ of landslides ($\sim 0.13\text{ km}^2$ of landslide area) were in the steep and very steep slope classes (with inclinations in the range of $36\text{--}90^\circ$), and $\sim 2\%$ of landslides ($\sim 0.13\text{ km}^2$ of the landslide area) were in flat areas (with no inclination, 0°). Based on these data (and field observations), it can be concluded that the “problematic” slope inclination values for the Kravarsko PA are the following ranges: (i) $1\text{--}5^\circ$, almost flat areas covering $\sim 34\%$ and $\sim 21.0\text{ km}^2$ of the PA, with $\sim 16\%$ of landslides and $\sim 1.04\text{ km}^2$ of the landslide area; (ii) $6\text{--}15^\circ$, gentle slope areas covering $\sim 45\%$ and $\sim 27.8\text{ km}^2$ of the PA, with $\sim 53\%$ of landslides and $\sim 3.45\text{ km}^2$ of the landslide area; and (iii) $16\text{--}35^\circ$, slopes and inclined slope areas covering $\sim 17\%$ and $\sim 10.5\text{ km}^2$ of the PA, with $\sim 27\%$ of landslides and $\sim 1.76\text{ km}^2$ of the landslide area. In general, for the Kravarsko PA, in almost flat areas ($1\text{--}5^\circ$), soil creeps (long-lasting and shallow movements) are common; meanwhile, on slopes ($6\text{--}35^\circ$), slides in soil (which can be described as rotational, planar, or compound slides in clays and silts or as debris slides in gravels and sands) are more frequent [32,49].

For the Kravarsko PA, alluvial sediments were present only in $\sim 3\%$ of the area ($\sim 1.9\text{ km}^2$ of the PA), with only $\sim 0.1\%$ of landslides and $\sim 0.01\text{ km}^2$ of the landslide area. Therefore, for the Kravarsko area, alluvial sediments (in general) can be regarded as not being landslide-prone areas. Loess-type sediments were present in $\sim 15\%$ of the area ($\sim 9.3\text{ km}^2$ of the PA), with $\sim 4.3\%$ of landslides and $\sim 0.28\text{ km}^2$ of the landslide area. Therefore, for the Kravarsko area, loess-type sediments (in general) can be regarded as somewhat landslide-prone areas. Vrbova fm. was present in $\sim 82\%$ of the area ($\sim 50.6\text{ km}^2$ of PA), with $\sim 95.6\%$ of landslides and $\sim 6.22\text{ km}^2$ of the landslide area, as shown in Figure 5. In general, for the Kravarsko PA, landslides occur in Vrbova fm. sediments (landslide-prone areas [26,32,40,41]), and they are shallow (those with a deepest sliding surface of up to 5 m in depth are the most frequent) and small to medium in ($100\text{--}20,000\text{ m}^2$) [7].

For the Kravarsko PA, a new land-cover map was developed, where surface water areas were present in less than 0.1% of the area ($\sim 0.1\text{ km}^2$ of PA); these were areas without landslides. Forest areas were present in $\sim 59\%$ of the area ($\sim 36.4\text{ km}^2$ of PA), comprising $\sim 78\%$ of the landslides and $\sim 5.01\text{ km}^2$ of the landslide area. Urban areas were present in $\sim 6\%$ of the area ($\sim 3.7\text{ km}^2$ of PA), with $\sim 1\%$ of the landslides and $\sim 0.07\text{ km}^2$ of the landslide area. Field areas were present in $\sim 35\%$ of the area ($\sim 21.6\text{ km}^2$ of PA), with $\sim 21\%$ of the landslides and $\sim 1.37\text{ km}^2$ of the landslide area (Figure 6). The land-cover data should be critically considered with the following points kept in mind: (i) the majority of the landslides are registered in forest areas due to the available mapping method (hrDTMh), whereby the landslide features are “preserved” (more or less) in their natural environment; (ii) in fields (due to cultivation), landslide features are “quickly” masked/destroyed, and “older” landslides cannot be mapped; and (iii) in urban areas (as for the fields but due to mitigation measures), landslide features are “quickly” masked/destroyed, and “older” landslides cannot be mapped, even though for the Kravarsko PA, the anthropogenic factor (in addition to the geological setting) often plays a key role in the landslide activation process [32,60,61].

4.2. Comments on the Kravarsko Landslide Inventory and Landslide Susceptibility Map

For the Kravarsko PA, a new landslide inventory and landslide susceptibility map were developed. While both of these datasets are valuable, they have their limitations in terms of practical usage. The landslide inventory was developed based on hrDEMs, which

contain precise data. However, the interpretation of the landslide inventory (landslide mapping) is a subjective process [19,31,32]; therefore, periodic LiDAR scanning and updates of the landslide data are recommended. Additionally, periodic data gathering and updates could serve as a basis for landslide hazard map development; this would provide the much-needed time component in landslide datasets [17,20,21]. In LSM for the Kravarsko PA, ~17% of landslides and ~1.11 km² of the landslide area are in low-LSM zones (~31.4% and ~19.1 km² of the PA area), ~43% of landslides and ~2.80 km² of the landslide area are in the medium LSM zone (~35.9% and ~22.3 km² of the PA area), ~32% of landslides and ~2.08 km² of the landslide area are in the high-LSM zone (~27.2% and ~16.7 km² of the PA area), and ~8% of landslides and ~0.52 km² of the landslide area are in the very-high-LSM zone (~5.4% and ~3.7 km² of the PA area), as shown in Figure 7. In general, based on the resulting LSM, ~70% of the Kravarsko PA can be considered a “landslide-prone area”, where ~85% of the landslides occur. At first glance, it is a high-value “landslide-prone area”, but it can be directly correlated with the defined EGUs for the research area: Vrbova fm. landslide-prone sediments cover more than 80% of the investigated area (Figure 5).

4.3. Comments on the Kravarsko Area Terrain Stability Map

From the perspective of optimizing natural hazard management, the Kravarsko area terrain stability map has the following advantages: (i) the crucial management area (areas with landslides) is smaller in terms of TSM than LSM (~50% vs. ~70% of the Kravarsko PA, respectively), and (ii) the at-risk key infrastructure areas are clearly defined with concrete values (Table 3). It is also worth noting that although landslides are concentrated in Vrbova fm., as is the case for settlements, roads, power lines, and water systems, the TSM data allow us to focus on ~50% of the PA in which ~90% of the landslides occur: ~60% of settlements areas, ~70% of roads, ~60% of water systems, and ~40% of power lines in the PA represent management priorities because they are endangered by landslides.

Based on the relatively simple analysis presented here, the area of (intensive) natural hazard management is reduced practically by half (50%), and key infrastructure “maintenance” is reduced by 40–60%. Of course, this does not mean that key infrastructure in possibly unstable zones (or even stable zones) should not be checked or maintained, but the frequency and the cost of such checks can be reduced, as can the eventual mitigation measures needed. That stated, the TSM can directly reflect (positively) budget optimization processes and future (urban) planning in the local community.

4.4. Reflections and Comments on the Presented Perennial Research

4.4.1. Innovation and Novelty of the Research

There are many different types of research conducted and a variety of methods used in the development of LSMs [35–39], but the novelty of the present results is that they do not focus on the LSM but on the development of a reliable LI based on high-resolution RS data in combination with the available digital key infrastructure data in order to provide a simple and usable map for the local community: the TSM. In this sense, the TSM presented herein is a “practical” novelty, and the presented methodology can be used/upscaled for larger areas and regions or at a national level. With adequate data, there are no limitations; only the final TSM scale and “purpose of the use” must be kept in mind and defined accordingly.

4.4.2. Validation and Field Research

The developed TSM is based on RS and field data. For the Kravarsko PA, old and new detailed RS data were analyzed, and the developed LI was validated in an on-the-ground setting for 10% of the research area and 15% of the developed LI [32]. Moreover, extensive field work was carried out (described in detail in [26,32]). The field data collected to determine the on-site material characteristics included 16 shallow boreholes, 43 field pocket penetrometer tests, 29 field shear vane tests, and 676 field points (geological and engineering

geological) for the VA with 113 field points (geological) and ~200 different laboratory tests for the wider area (granulometry, mineralogy, XRD, CaCO₃ content, etc.) [32].

4.4.3. Research Limitations

It must be remembered that the quality of the developed TSM is based on the quality of the input data. Limitations in TSM development can be associated with the quality of the RS data and the available LI and infrastructure data. RS data quality is affected by the season, light conditions (i.e., the length of a day and angle of sun beams), and the density of the vegetation cover. To minimize these limiting conditions for the Kravarsko area, airborne LIDAR scanning was conducted in early spring 2018 with 20 points per m²; at the same time, orthophotos and stereopairs with pixel sizes of 10 × 10 cm were acquired. Based on these data, hrDEMs (0.5 × 0.5 m pixel size) were developed for analysis. The developed LI is based on high-quality hrDEMs analysis; nevertheless, the subjectivity of the expert(s) [31,32] cannot be neglected. The following main problem in LI development remains: how are we to accurately identify “older” landslides when some of their features are masked/degraded/poorly visible or non-visible [31,32]? As for infrastructure data, they should be provided by the local community or regional management, depending on the final scale of the developed TSM.

4.4.4. Broader Implications and Scalability

The methodology of the present study can be applied to larger areas and regions or used at the national level, regardless of the topography and climate. In this regard, the Kravarsko PA (≈61.7 km²) can be considered a smaller test polygon for which the results (the developed LI and TSM) were positive. The next step involves upscaling the presented methodology on a county level, i.e., using it for Zagreb City County (≈641 km², an area 10 times larger).

4.4.5. Research Steps and Technical Explanations

The perennial research presented herein can be described as three-step study with certain prerequisites. These prerequisites are as follows: a defined area of research, hrDEMs of the area, and a developed scoring system for the classification of identified landslides and key digital infrastructure data. The steps are as follows:

- Step 1: The interpretation of hrDEMs and LI development based on the applied scoring system (described in more detail in [32]);
- Step 2: The verification of results in terms of the cabinet (historical data review) and field research (calibration of the mapping and scoring system, if needed) described in more detail in [32];
- Step 3: Product development, i.e., the development of a TSM with key infrastructure (presented in this paper).

All of the research was performed within CAD (AutoCad2023 and StereoCad) and the GIS environment (ArcGIS 10.2.1.). The basic statistical analyses are described in [32], while more detailed statistical analysis is intended for publication along with stakeholder feedback data and new research results.

5. Conclusions

In some areas (for example, the Kravarsko PA), landslides represent a constant threat to safety and property. Landslides cannot be predicted exactly, but the conditions and areas (zones) in which they are activated can be estimated with great confidence, provided there are adequate data available. In landslide analysis, a reliable landslide inventory is of significant value, as it is a basis for developing a viable natural hazard (landslide) management plan for a certain area. In optimized landslide management, the area of interest should be divided into (smaller) zones with the same characteristics; this can be achieved by developing an LSM based on basic data (slope gradient, geology, and land cover) or by developing a TSM, such as that presented here, based on the available relevant

data (inventory and infrastructure). Basic LSMs can provide adequate solutions in cases when landslide inventories are not available, but a TSM with a landslide inventory, and key infrastructure data are on the “user-friendly side” for non-expert users engaged in natural hazard management, for example, local administrations. At the same time, it can provide easily understandable information for the local community. It is reasonable to expect future landslides to occur in the areas where they have occurred in the past or where they are occurring now [72]. The results of the present study can assist in the management of natural hazards and urban planning:

- The LSM presented here was developed for the Kravarsko PA; however, it should be noted that the developed LSM is heavily influenced by terrain slope inclination and geology (EGU) data.
- The developed TSM is also heavily influenced by landslide inventory data.
- The novelty of the results presented herein resides in their focus on the development of a reliable LI based on high-resolution RS data in combination with the available digital key infrastructure data in order to provide a simple and usable map for local community, i.e., the TSM.
- With the TSM, the areas with higher landslide frequency (“risk”) are reduced (optimized). In the TSM, the area of (intensive) natural hazard management is reduced practically by half (50%), and key infrastructure “maintenance” is reduced by 40–60% for the Kravarsko PA.
- We advise that future periodic (landslide) data gathering, the analysis of hazards, and the development of risk maps are still needed for the PA; these measures would improve the present state of natural hazard (landslide) management in the Kravarsko area.
- The presented methodology for TSM development is described according to three steps and defined prerequisites.
- With the provision of adequate data, the TSM can be used/upscaled for larger areas and regions or used at the national level, regardless of the differing topographies and climates of these regions.

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