



Clay Mineralogy and Geochemistry of the Pockmarked Surface Sediments from the Southwestern Xisha Uplift, South China Sea: Implications for Weathering and Provenance

Mei Zhang ^{1,2,*}, Hongfeng Lu ³, Qing Chen ³, Gayan Bandara ^{4,5}, Hui Zhang ¹, Chuanxiu Luo ^{4,5} and Nengyou Wu ⁶

- Key Laboratory of Gas Hydrate, Guangzhou Institute of Energy Conversion, Chinese Academy of Sciences, Guangzhou 510640, China; zhanghui@ms.giec.ac.cn
- ² Key Laboratory of Metallogenic Prediction of Nonferrous Metals and Geological Environment Monitoring, Central South University, Ministry of Education, Changsha 410083, China
- ³ Guangzhou Marine Geological Survey, Guangzhou 510760, China; luhongfeng@hydz.cn (H.L.); chenqing.0913@163.com (Q.C.)
- ⁴ Key Laboratory of Ocean and Marginal Sea Geology, South China Sea Institute of Oceanology, Chinese Academy of Sciences, Guangzhou 510301, China; wgssenarathna12@mails.ucas.ac.cn (G.B.); xiu104@scsio.ac.cn (C.L.)
- ⁵ University of Chinese Academy of Sciences, Beijing 100049, China
- ⁶ The Key Laboratory of Gas Hydrate, Ministry of Land and Resources, Qingdao Institute of Marine Geology, Qingdao 266071, China; wuny@ms.giec.ac.cn
- * Correspondence: zhangmei@ms.giec.ac.cn; Tel.: +86-020-87076246

Abstract: In the northern South China Sea, pockmarks are widely distributed on the seabed offshore on the southwestern Xisha Uplift. The mineralogy and geochemistry of the clay minerals and surface sediments from the pockmark field were identified using X-ray diffraction (XRD) analysis and X-ray fluorescence (XRF) analysis to trace the provenance, weathering, and sediment transportation system in the area. The clay minerals are primarily comprised of illite, smectite, kaolinite, and chlorite, showing a distribution of average weight percentages of 35%, 35%, 18%, and 13%, respectively. Based on the surrounding fluvial drainage basins and various transport mechanisms (current or monsoon), illite and chlorite primarily originate from rivers in Taiwan and the Mekong and Red Rivers. Kaolinite primarily originates from the Pearl River, and smectite derived from the Luzon arc system is primarily transported by surface currents with significant influence from the Kuroshio intrusion.

Keywords: surface sediments; clay minerals; sediment provenance; sediment transport; Xisha Uplift

1. Introduction

Understanding processes involved in modern sedimentary environments is a prerequisite in paleo-oceanographic reconstruction [1–3]. Climatic conditions and lithology of the source rock are the key factors generally controlling the geo-chemical and the mineralogical fingerprints of the surface sediments. Clay minerals, which are widely distributed in the marine sediments, are vital contents, as well as the content assemblages and concentration of chemical elements, which provide useful tools for deciphering provenance, weathering intensity, transport patterns, and paleo-climatic changes (e.g., [4–14]).

Clay minerals, such as illite and chlorite, are common weathering products of igneous and metamorphic rocks. These minerals are typically found in high-latitude marine sediments after mobilization by physical weathering. Kaolinite forms under warm and humid conditions by the intensive chemical weathering of feldspar in tropical soils and is often denoted as a low-latitude mineral [15,16]. The alteration of volcanic rock-derived smectite can be a good indicator of volcanic sediment sources.

Terrigenous sediment loads are received into the northern South China Sea (SCS) annually, from the Red River (130 metric tons (Mt)) into the northwest, Pearl River (69 Mt)



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into the north, rivers in Taiwan (187.5Mt) into the northeast, and some rivers on Luzon Island (>8.2 Mt) into the east [5,17]. However, the detrital sediments in the region have an unknown provenance due to this complexity in the source of the detrital sediments and circulation patterns in the northern SCS. In previous studies, especially in the northeastern region of the SCS, the Pearl River and volcanoes were considered as the primary sedimentary sources [18–21]. Nevertheless, recent investigations on elemental geochemistry and clay mineralogy have focused on three major origins: Taiwan, the Pearl River, and Luzon, where the contribution to the northern slope of the SCS was semi-quantitatively assessed [22,23].

A number of sedimentological and geochemical approaches have been used to analyze of the source of SCS sediments, such as measurements of grain size distribution [24–26], the occurrence of clay and heavy minerals (e.g., [5,22,23,27]) strontium and neodymium isotopes (e.g., [10,28,29]), and major and trace element geochemistry [5].

The SCS is a semi-enclosed basin and known as one of the largest marginal seas in the world, and is located between the Pacific Ocean and the Asian continent. Previous studies mainly focused on the northeastern region of the SCS, due to the complexity of the origin of detrital sediment in the SCS, the interaction and patterns of sea surface currents in combination with the East Asian monsoon winds, and the subsurface intrusion of Kuroshio and deep waters through the Luzon Strait from the western Pacific [22,30–32]. There is a lack of studies on the provenance and geochemistry of sediments on the Xisha Uplift. The aim of this study is to adequatelycorrelate the distribution characteristics of the clay minerals, major elements, and element geochemistry of surface sediment with the relationships between provenance and the sediment transportation in the Xisha Uplift, southwestern SCS.

2. Geological Setting

The study area is located in the southwestern Xisha Uplift between the Guangle Uplift and Zhongjian Island, northern SCS (Figure 1), where the water depth ranges from 800-1300m. The evolution of the sedimentary filling can be categorized into the Eocene-Oligocene rift and a Neocene–Quaternary post-rift subsidence. The deposition of the total thickness of the sedimentary layers occurred during the Eocene–Oligocene rift stage, which could be the major high-quality source rocks in this study area. The Neocene-Quaternary post-rift subsidence stage can be further categorized into thermal subsidence and accelerated subsidence sub-stages. The study is characterized by high sedimentation rates (up to 1.2 mm/yr) and high geothermal gradients (39-41 °C/km) [33-35]. A high percentage of the pockmarks within the study area are classified as mega pockmarks (1000–2500 min diameter and 60–140 min depth). The distribution of these pockmarks is linked with the underlying fluid migration structures, such as gas chimneys, polygonal faults, faults, unconformities, and paleo-channels [36] Moreover, the phase boundaries between free gas and solid gas hydrate for this area have been described with the distinct bottom-simulating reflectors (BSRs) [37]. Furthermore, compelling evidence on gas hydrate dissociation, such as decreased chlorinity and increased pore water δ^{18} O, has been reported (in core C14) in the pockmark field [38,39]. Meanwhile, high geothermal gradients in the region accelerated the maturity of the source rock, and drifted towards the seabed, which refers to the composition change of the sediments.



Figure 1. Map of the study area and sampling sites (modified from Sun et al., 2011).

3. Materials and Methods

A total number of 8 surface sediment samples were collected using a box sampler from the southwestern region of the Xisha Uplift, northern region of SCS, cruised by the South China Sea Institute of Oceanology, CAS on the Shiyan-1 in May 2012. The samples were collected either from or near the giant pockmarks in the southern region of the Qiongdongnan (QDN) basin (details are shown in Figure 1). The characterization of the sediments was done by grayish-green foraminiferous silty clay.

The X-ray powder diffraction (XRD) method was used for the measurement of the mineralogical composition in the bulk sediments. For each sample, 5 g of sediment sample were taken and oven-dried for 24 h at 45 °C, and ground to \leq 200 mesh size in an agate mortar. For the XRD analysis, approximately 1 g of powder was used. The results were semi-quantitatively elucidated using Jade 6.0 software, based on the relationships between the mineral diffraction peak heights and integral area.

Precisely weighed 4g of powder were subsequently added into the center of a polyoxyethylene abrasive apparatus and made into circular samples (4 cm diameter and 8 mm thickness) with the apparatus pressurized for 20 s at 30 tm⁻² pressure. Major oxides and a few trace elements were determined by a Philips Panalytical Magix PW2403 X-ray fluorescence (XRF) spectrometer (Holland) at standard room temperature and pressure (approximately 20 °C, 85 kPa).

Carbonate and organic matter were removed using 10% H_2O_2 and 0.5N HCl, respectively. Then clay minerals (<2 μ m) were separated into 1000 mL glassware, according to Stoke's settling velocity principle [40]. XRD runs were performed three times, followed by air-drying, ethylene glycol (55 °C for 9 h), and heating at 550 °C for 2 h. Based on Liu et al. [5], clay mineral abundances were semi-quantitatively calculated.

The Axios X-ray fluorescence spectrometer (XRF) was used for measuring of the percentages of the major elements, including Al₂O₃, SiO₂, Fe₂O₃, MgO, CaO, K₂O, MnO, Na₂O, P₂O₅, and TiO₂. Loss on ignition (LOI) was determined by referring to a method proposed by Heiri et al. [41].

Normalization of the concentration of selected elements was done by considering the standard concentration of elements in the upper continental crust (UCC; [42,43]). The chemical index of alteration (CIA) was used for the estimation of the intensity of chemical weathering [44]:

$$CIA = Al_2O_3 / (Al_2O_3 + Na_2O + K_2O + CaO_*) \times 100$$

where CaO* is only found in silicates.

In this study, there were two correction methods used, because precise correction of CaO bound in carbonate and phosphate minerals is difficult [45]. The CIA index was calculated from chemical analyses following acid treatment to remove carbonates and phosphates.

4. Results

4.1. Principal Mineral Composition in Surface Sediments/Mineralogy

The studied samples were mostly silt, with abundant foraminifera. XRD analysis of bulk sediment revealed that minerals, including clay minerals (34.8–42.9%), calcite (20.2–24.1%), quartz (15.9–18.4%), minor plagioclase, dolomite, siderite and pyroxene, pyrite, gypsum, anhydrite, k-feldspar, and amphiboles were also detected in the samples (Figure 2, Table 1).



Figure 2. X-ray diffraction diagram of surface sediments in the Xisha Uplift (s9). M: muscovite, II: illite, Cli: clinopyroxenite, Gy: gypsum, Q: quartz, C: calcite, Sid: siderite, Dolo: dolomite, Ha: halite, Py: pyrite.

Table 1. Mineralogical assemblages in bulk sediments of the southwestern Xisha Uplift, South China Sea (SCS).

Samples	Water Depth	Quartz	K-Feldspar	Plagioclase	Calcite	Dolomite	Siderite	Pyrite	Amphiboles
s10	854	17.2	0.7	5	24.1	4	4.8	1.1	0.2
s11	806	18.4	2.4	5.1	23.2	2.6		1.3	0.5
s15	918	15.9	0.5	4	22.4	7.3	6.1		
s18	999	16	0.5	4.6	21.3	7.4	5.8	1	0.3
s13	864	16.9	0.3	4.4	23.4	7.2	5.7		
s9	811	17.5	0.6	5.2	23.6	4	1.3	0.8	0.3
s19	1099	16.2	0.6	4.4	20.2	5.6	5.5	0.8	0.2
s20	1307	18	0.8	5	20.2	4.3	1.5	1.5	
Samples	Water Depth	Gypsum	Anhydrite	Pyroxene	Clay minerals	Smectite	Illite	Kaolinite	Chlorite
s10	854		2.6	3.8	36.5	33	36	12	19
s11	806	0.2	2	3.5	40.8	32	37	13	18
s15	918		3.6	3.9	36.3	24	41	13	22
s18	999		2.6	3.3	37.2	23	40	16	21
s13	864	0.1	3.3	3.9	34.8	38	34	11	17
s9	811	0.1	2.7	4	39.9	42	37	9	12
s19	1099		2.9	3.9	39.7	44	29	11	16
s20	1307		2.2	3.6	42.9	41	28	15	16

Clay minerals in sediments carried by rivers and small mountainous rivers comprised ~80% of total SCS surface sediments [46,47]. Biogenic carbonates with negligible biogenic silicates and volcanic materials are primarily observed as other components of the sediments. In this study, the clay mineral components of eight samples are characterized by moderate smectite (24–44%, average 35%) and illite (29–40%, average 35%), with a lesser abundance of chlorite (16–22%, average 18%) and kaolinite (8–16%, average 13%) (Table 1; Figure 3).



Figure 3. Clay mineral composition of Xisha Uplift surface sediments (<2 microns). S: smectite, It: illite, Kao: kaolinite, C: chlorite.

4.2. Geochemistry

The composition of the major elements in the bulk sediments are characterized by high SiO₂, CaO, and Al₂O₃ (total average content is 66%) and low K₂O, Fe₂O₃, Na₂O, MgO, P₂O₅, TiO₂, and MnO (total average content is 11%) (Table 2). Major elements increase when SiO₂ increases, however, Ca decreases with SiO₂, except for Na and Mn (Figure 3). Xisha sediments enriched in CaO and depleted in Si, Na, Al, Fe, Mg, and K correspond to the UCC standard.

Table 2. Major element composition (%) and chemical index of alteration (CIA) of surface sediment samples in the southwestern Xisha Uplift, SCS.

	TFe ₂ O ₃	K ₂ O	MgO	MnO	Na ₂ O	P_2O_5	SiO ₂	TiO ₂	Al ₂ O ₃	CaO	LOI	CIA
s9	4.24	2.08	1.95	0.35	2.27	0.13	37.29	0.52	11.49	18.08	21.25	54
s10	4.06	1.97	1.90	0.12	2.47	0.12	35.02	0.48	10.65	20.26	22.97	51
s11	2.91	1.33	1.37	0.22	1.83	0.09	23.51	0.32	7.24	31.61	29.94	49
s13	4.15	2.02	1.96	0.09	2.53	0.12	36.71	0.51	10.81	18.91	22.51	51
s15	4.41	2.14	2.09	0.17	2.86	0.13	39.16	0.55	11.58	16.48	20.71	50
s18	3.49	1.75	1.73	0.29	2.52	0.11	29.85	0.41	9.41	24.45	26.37	48
s19	4.33	2.17	1.97	1.04	1.95	0.13	37.96	0.52	11.81	16.78	20.71	57
s20	4.86	2.46	2.29	0.96	2.97	0.13	44.34	0.62	13.32	10.91	17.41	52

5. Discussion

5.1. Sediment Provenance

Previous studies have indicated that different sediment sources may originate different clay minerals in the SCS. Terrigenous sediments transported to the northern SCS, through by fluvial input from southern China, and loaded by the large Pearl River and Taiwan and Luzon Islands via small mountainous rivers [5,21,23,24,48]. Smectite in marine sediments, related to volcanic materials, volcanic activity, or the alteration of volcanic materials by hydrothermal, weathering, or halmyrotic processes [21,49,50]. There are three areas with high smectite content, considered as significant sources in the SCS. The first is around the

Philippines (west of Luzon Island), where 15–50% account for volcanic materials in the silt fraction of sediments [4]. The second area is the east of the Indochina Peninsula, where smectite originates from the Mekong Basin [51] and/or the Red River [52]. The third area is adjacent to the Wan'an Shoal, enriched with volcanic glass [53]. In the surface sediments of the studied area, smectite is the dominant clay minerals, with the average contents of 34%, which means the three high-smectite sources are possible.

Illite and chlorite in marine sediments are some major components which are derived from the continent [48]. The six predominant areas of high illite content in the SCS may be the possible sources for the Xisha Uplift: rivers in Taiwan (48–66%, average 56%), Pearl River (21–51%, average 35%), Red River (31–58%, average 44%), Mekong River (33–42%, average 37%), north Borneo (44–75%, 56%) and north Palawan (38–52%, average 42%) [27].

The formation of kaolinite is mainly considered in tropical conditions from welldeveloped ferrallitic soils in a plain environment with active hydrolysis processes [54,55]. The Pearl River, Hainan Island, and the southern Sunda Shelf are the kaolinite-rich areas. Liu [1] analyzed the drainage basin surface sediments of the Pearl River, the results indicated that kaolinite is common in clay mineral assemblages, while chlorite and illite are less abundant, and smectite is scarce. Possible sources for kaolinite and smectite on the Indonesian and Malaysian islands are abundant andesitic volcanic rocks and older granitic intrusives [48,56–58]. Sediments from the Taiwanese rivers primarily consist of illite and cholorite, with scarce smectite and kaolinite [5]. According to previous studies, the Taiwan-sourced sediments can be transported westward via the Kuroshio Current along the northern slope [58].

Luzon sediments, which are combined with weathering of volcanic materials, are generally characterized by a high percent value (average 46%) of smectite, which generally transports to the northwest of the island via the Kuroshio Current. Taiwan-sourced sediments are characterized by high percentages of illite (69%) and chlorite (30%) due to a relatively higher rate of physical weathering and relatively moderate chemical weathering processes [59] and, moreover, sediments which are characterized by a high percentage value of kaolinite in the Pearl River [1].

We carried out a cluster analysis and developed a ternary diagram of smectite– (illite+chlorite)–kaolinite in the studied area and the adjacent regions, to better constrain the provenance of clay minerals sampled from the Xisha Uplift. The distributions of the clay mineral assemblages in the surface sediments of the Xisha Uplift are shown in Figure 4. Clay minerals in the sediments plotted and fell between the data from the Luzon River and the Mekong River. Further, the clay mineral assemblages showed higher smectite and illite values. Results of the data show that the clay mineral components acquired greater contributions from the Red River and rivers in Taiwan than from the Pearl River. Therefore, illite is primarily supplied to this region by the Red River and rivers in Taiwan, however, the study should extend towards a further analysis in the future.



Figure 4. Ternary diagram of clay minerals of surface sediments from the study area compared with those from the adjacent regions.

Data of Taiwan rivers [22] and Pearl, Red, and Mekong Rivers [1], Luzon rivers [52], north Borneo rivers [58], and Hainan rivers [27] are plotted for comparison.

Recent studies indicated that weathered materials of both the Mekong and Red Rivers are directly involved in forming illite in the southwestern Xisha Uplift sediments, and smectite is primarily derived from the Luzon River [22]. Similar trends of the variation can be seen in analogous minerals like illite and chlorite and these are depleted in correlation with smectite composition (Figure 5), suggesting that the Red River and Mekong River are not the principal source of illite and chlorite. We concluded that the illite originates primarily from Taiwan.



Figure 5. Variation diagrams of clay minerals of surface sediments in the Xisha Uplift.

5.2. Quantifying Weathering Intensity in the Xisha Uplift

Clay minerals appeared as the weathered products of parent rocks on the Earth's surface. The composition of the clay minerals provide important information regarding weathering types and intensity, which are primarily determined by various climatic conditions (rainfall, temperature), rock composition, and tectonic activities [54,57,60–64]. Chemical weathering through the hydrolysis of minerals alters the parent rock composition. This process involves the production of typical weathering products like kaolinite and smectite. Generally, kaolinite is found in monosialitic soils, which are characterized by the complete removal of mobile cations and an extreme hydrolysis process. Furthermore, this is primarily controlled by the continental hydrolysis of the parent aluminosilicate rocks in warm and humid climatic conditions [60]. Kaolinite is abundant around the Pearl River, the Malay Peninsula, and Hainan [27,58,63], which suggested a warm and humid climate. Smectite is conventionally related to the chemical weathering of volcanic rocks under hot and humid climatic conditions [65,66]. A rapid rate of volcanic rock weathering is often associated with a high abundance of smectite, which easily forms on basic materials, such as Fe-Mg species [60]. For instance, volcanic rocks are predominant in Luzon and in the Mekong River areas with abundant smectite. Both illite and chlorite are primary minerals, which are strongly correlated with the physical erosion of bedrocks, with relatively dry

and cold climatic conditions. To identify and evaluate the weathering intensity, we use the CIA and the elemental ratios calculated with respect to the least mobile element Al, which according to the mobile elements (e.g., K, Na, and Ca), was depleted in the parent rocks, but the least mobile elements (e.g., Al, Fe, and Ti) were enriched in the weathering products during chemical weathering [44].

The sediments are derived from heterogeneous sources, and the state of chemical weathering of the rocks is quantified by the CIA values with reference to the loss of mobile elements, such as Na, Ca, and K. The results demonstrate that CIA values for bulk sediments are 48–57, which indicates low chemical weathering intensity, compared with the terrestrial materials from the South China coast. Comparing previously published CIA values with this study showed that chemical weathering occurred at the same intensity.

Here, the elemental ratios were calculated with respect to Al, and used to identify and assess the mobility of major elements. The content ratio of element X and Al₂O₃ in river samples divided by the ratio of the elemental content of UCC provides the following elemental ratio [67]: elemental ratio (X) = $(X/Al_2O_3)_{rivers}/(X/Al_2O_3)_{UCC}$. Hence, the elemental ratio indicates the relative enrichment or depletion of the element, i.e., >1 indicates enrichment, <1 indicates depletion, and 1 indicates no change in the relative abundance of the element. The elemental ratios calculated from average major element concentrations normalized to UCC are shown in Figure 6.



Figure 6. Elemental ratios of surface sediments in the Xisha Uplift normalized to the upper continental crust (UCC) (Taylor and McLennan, 1985) with respect to Al₂O₃.

With the predominance of smectite clay minerals are formed by chemical weathering, which is carried out by the leaching of Na and K and later of Fe and Mn. The combination of both results of the CIA and elemental ratios showed a low–moderate degree of chemical weathering of surface sediments in all analyzed sections of the Xisha Uplift.

5.3. Transport Mechanism of Clay Minerals in the Northern SCS

Recent research works focused on the clay minerals of surface sediments and core sediments to discover the sources, transport pathways, and paleo-environmental conditions (e.g., [4,20–23,26,55,68–73]). The dispersal and transport mechanisms of the clay minerals were deduced from the provenance supply and ocean circulation patterns. Most ocean circulation studies indicated that sediment transport in the SCS is controlled by both the surface and deep water currents (DWC) [27,48]. For this study, the distribution of clay minerals is closely related with the current patterns. The relatively high smectite contents prevail in the Xisha Uplift, highly correlated with the flow shift from the west-

ward SCS Kuroshio branch to the northeastward SCS warm current [27], which generates numerous mesoscale westward propagating eddies to transport the smectite from the Luzon Strait westward to the Xisha Uplift, before the deposition on the slope and outer shelf [63].Therefore, the surface currents may be primarily involved in the transportation of smectite. The high illite and chlorite content in the Xisha Uplift indicates the provenance from the Red River, the Mekong River, and rivers in Taiwan. The illite content in both Red and Mekong Rivers is higher or equal to that in the Xisha Uplift [1]. Moreover, in the Red and Mekong River basins, humid and cold climatic conditions increase the physical erosion and decrease the hydrolytic weathering processes of metamorphic and granitic parent rocks. This weathering process is associated with the production of a high illite content [1]. The high physical weathering rate in Taiwan is consistent with the present illite chemistry index and illite crystallinity data with complete Fe–Mg-rich illite.

6. Conclusions

Surface sediments from the Xisha Uplift in the SCS were measured for clay minerals, primary elements and minerals for tracing sediment provenances, weathering intensity, and transport. Our results emphasize that the clay mineral assemblages of the study areas primarily consist of smectite (24–44%, average 35%) and illite (29–40%, average 35%), with a lesser abundance of chlorite (16–22%, average 18%) and scarce kaolinite (8–16%, average 13%). In conclusion, sediments in the study areas are derived primarily from large rivers in the northern SCS (the Mekong and Red Rivers and Luzon River). We suggest that the DWC and surface monsoon circulation are the primary transport dynamics for sediments in the Xisha Uplift by combining the clay mineral compositions of the surrounding basins and surface sediments in the northern SCS.

Author Contributions: H.Z., H.L., Q.C., C.L. and N.W. performed the expedition with fieldwork and sediment sampling and test; M.Z. wrote the paper. M.Z. and G.B. revised the paper. All authors have read and agreed to the published version of the manuscript.

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Data Availability Statement: Data available on request.

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