

Article



# Investigating the Potential Role of Geological Context on Groundwater Quality: A Case Study of the Grenville and St. Lawrence Platform Geological Provinces in Quebec, Canada

Roxane Tremblay \*<sup>D</sup>, Julien Walter <sup>D</sup>, Romain Chesnaux and Lamine Boumaiza <sup>D</sup>

Centre d'Études sur les Ressources Minérales, «Risk Resources Water» (R2eau) Research Group, Département des Sciences Appliquées, Université du Québec à Chicoutimi, Saguenay, Québec, QC G7H 2B1, Canada; julien\_walter@uqac.ca (J.W.); romain\_chesnaux@uqac.ca (R.C.); lamine.boumaiza@uqac.ca (L.B.) \* Correspondence: tremblayroxane3@gmail.com

Abstract: The hydrogeochemical study of the Lanaudière and Eastern Mauricie regions (Canada) demonstrates that trace elements appear to be better tracers of geological influence on groundwater chemistry than major elements. Isotopic ratios and the similar chemical composition of groundwater suggest that the physicochemical parameters of groundwater have a greater effect on hydrogeochemical mechanisms than the immediate geological environment The results allow us to propose a conceptual model of groundwater geochemical evolution with the aim to guide the protection and sustainable management of regional groundwater resources in the Lanaudière and Eastern Mauricie regions. These regions were selected because of their location at the boundary of the Grenville and St. Lawrence Platform geological provinces, representing two distinct geological contexts (Precambrian crystalline rocks and Paleozoic sedimentary rocks). Regional-scale hydrogeochemical and isotopic groundwater characterization was carried out to identify the role of the differences in regional geology on groundwater quality. Our analyses included major and trace elements, stable isotopes, and multivariate statistics. Similar processes are at the origin of dissolved major chemical elements and suggest that soluble minerals common to both geological provinces control groundwater chemistry. If differences exist, they are due to the hydrogeological conditions of the samples, such as residence time or groundwater entrapment at the time of the postglacial marine incursion of the Champlain Sea, rather than the geological context. Some differences, sometimes significant, were observed for some minor elements ( $F^-$ ,  $Mn^{2+}$ ,  $H_2S$ ), which implies a more comprehensive knowledge of the chemistry of the stratigraphic units within the Lanaudière and Eastern Mauricie aquifers.

**Keywords:** conceptual model; groundwater statistics; hydrochemistry; sedimentary rocks; crystalline rocks; Québec

# 1. Introduction

In Quebec, Canada, groundwater represents an economically exploitable source of potable water because of its quality and proximity to sites of its consumption [1]. For several years, hydrogeological data, including borehole logs and groundwater chemical analyses, have been collected by public and private organizations in Quebec. To produce a complete database accessible to hydrogeologists and water managers, the Quebec government implemented the Groundwater Knowledge Acquisition Program in 2008 (Programme d'Acquisition de Connaissances sur les Eaux Souterraines, PACES) under the direction of the Ministère de l'Environnement et de la Lutte contre les Changements Climatiques (MELCC) [1–3]. By the end of this project, approximately two-thirds of the administrative territories of southern Quebec will be characterized hydrogeologically with the aim of rendering the relevant data available for sustainable groundwater resource protection and management. As part of the PACES project, the Centre d'Études sur les Ressources Minérales of the Université du Québec à Chicoutimi (CERM-UQAC) was tasked to compile



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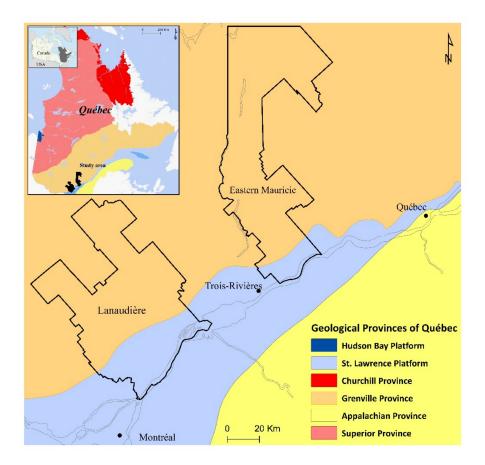
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existing hydrogeological information and then expand groundwater knowledge of the Lanaudière and Eastern Mauricie administrative regions of eastern Quebec (Figure 1).

**Figure 1.** Geological provinces within Quebec (inset) and the location of the study areas of the Lanaudière and Eastern Mauricie administrative regions.

Groundwater is an important geological agent due to (i) its interaction with the ambient environment and (ii) its spatial flowing distribution [4]. Chebotarev and Toth, for example, observed an increase in total dissolved solids (TDS) from groundwater inflow to outflow areas [4,5]. Because the lithologies of aquifers and the groundwater residence time are related to the geological context, groundwater chemistry is expected to differ between geological provinces. Six geological provinces underlie Quebec, i.e., the Hudson Bay and St. Lawrence platforms, as well as the Churchill, Grenville, Appalachian, and Superior provinces (Figure 1). Regional hydrogeochemical characterizations were conducted for the St. Lawrence Platform, Grenville, and Appalachian provinces as part of the PACES program [6–12].

Beaudry et al. [6], Lacasse [10], and Cloutier et al. [12] showed that the Champlain Sea incursion about ca. 11,000 to 13,000 years BP continues to affect groundwater quality in confined aquifers within the St. Lawrence Platform and Appalachian Province. These authors identified two main processes affecting groundwater geochemistry within the St. Lawrence Platform. The first is carbonate dissolution resulting in high concentrations of Ca<sup>2+</sup> and Mg<sup>2+</sup>. The second involves the Ca<sup>2+</sup>/Na<sup>+</sup> cation exchange and mixing with the trapped fossil pore seawater originating from the incursion of the Champlain Sea in southern Quebec; the result is highly saline groundwater (Na–Cl in composition). Multiple investigations of the groundwater chemistry of the Canadian Shield belonging to the Grenville Province identified silicate weathering and Ca<sup>2+</sup>/Na<sup>+</sup> cation exchange as the main processes governing regional groundwater chemistry [7–9,13,14]. Montcoudiol et al. demonstrated that water from the postglacial Champlain Sea continued to influence

groundwater quality at the junction between the St. Lawrence Platform and the Grenville Province [9]. The water–rock interaction between the Grenvillian rocks and groundwater is likely the main process responsible for the elevated concentrations in barium, boron, manganese, and fluorine in deep rock aquifers [8,13].

Most of the above-cited studies have established a conceptual model of the chemical evolution of groundwater of individual geological provinces, although the authors did not evaluate differences in geology among the study regions. The main objective of this study is to investigate the processes and mechanisms governing differences in groundwater chemical composition between two geological provinces, the St. Lawrence Platform, which is composed of sedimentary rocks, and the Grenville Province, which is composed of igneous and metamorphic rocks. The specific objectives are to (1) investigate different processes at the origin of groundwater chemistry and verify whether they are distinct to each geological province and (2) identify those natural contaminants that are spatially correlated with the respective geological provinces.

## 2. Materials and Methods

# 2.1. Study Area

The 11,845 km<sup>2</sup> study area includes the Lanaudière and Eastern Mauricie regions, both covering portions of the St. Lawrence Platform and Grenville Province (Figure 1). The southern portion of the regions is dominated by agricultural uses, whereas the northern regions are predominantly forested areas. The urban agglomerations are scattered and consist of mainly residential and commercial areas [15,16]. The Lanaudière region has 55 administrative municipalities spread over an area of 5962 km<sup>2</sup> with a population (2021) of 346,418 [17]. Eastern Mauricie has 12,387 inhabitants (2021) spread over an area of 5883 km<sup>2</sup>. The latter region is divided into 22 administrative municipalities [18].

Approximately 80% of the study area lies within the Grenville Province (Proterozoic), and 20% lies within the St. Lawrence Platform (Cambrian to Ordovician). The Grenville Province is composed mainly of gneiss (tonalitic and trondhjemitic), paragneiss, and marble. Some zones contain gabbros and metagabbros, whereas other zones in the province are composed of granite and pyroxeniferous monzonites [19–22]. The topography of the Grenville Province includes hills, plateaus, and low-lying areas. The average elevation is 390 m, and the highest point is located around Saint-Donat (884 m). The hydrographic network of the Grenville area is controlled by the pattern of hills being generally intersected by narrow valleys oriented N–NW/S–SE.

The St. Lawrence Platform corresponds to a succession of sedimentary units. The sediments were deposited initially in horizontal accreted units during the Appalachian Orogeny [23,24]. These units include sandstone conglomerates, dolomites, limestone, and shale; they form a complete sedimentary sequence representing a full transgression–regression cycle [23,24]. The St. Lawrence Platform corresponds to a low-lying area with an average elevation of 37 m, which slopes gradually toward the St. Lawrence River.

Quaternary materials covering both the Grenville Province and the St. Lawrence Platform were deposited during and immediately following the retreat of the Laurentide ice sheet. Till covers the Grenville Province and is composed of fine to coarse particles produced by erosion of the bedrock. At around 11,000 to 13,000 years BP, the Champlain Sea entered the isostatically depressed St. Lawrence Valley. This marine incursion induced a saltwater recharge under glaciomarine deposits and produced a marine clay aquitard through the deposition of fine silts and marine clays. Fluvioglacial sediments were deposited locally above the glaciomarine deposits and are often found in the St. Lawrence Valley [25].

The regional climate is characterized by six months of cold winters followed by short summers. Average winter temperatures range from -15 to -5 °C and vary between 10 and 19 °C in summer. Rain and snowfall combine for annual total precipitation of 825 mm/year for the Lanaudière and 752 mm/year for the Eastern Mauricie regions [26].

## 2.2. *Methodology*

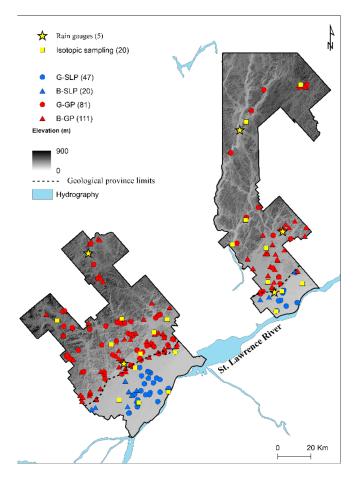
## 2.2.1. Groundwater Samples and Laboratory Analyses

The sampling campaign was conducted during the summer of 2019. A total of 314 groundwater samples were collected, including 231 from the Grenville Province and 83 from the St. Lawrence Platform. Groundwater samples were collected from private drinking wells. Groundwater volume, representing two to four times the volume contained in the well column, had to be purged prior to groundwater sampling to remove standing water [27]. Physicochemical parameters of groundwater—temperature (T; °C), electrical conductivity (EC;  $\mu$ S/cm), pH, redox potential (Eh; mV), dissolved oxygen (DO; %), and total dissolved solids (TDS; mg/L)—were monitored in situ during the purging process using a calibrated HANNA portable analyzer. Monitoring continued until three consecutive readings stabilized within  $\pm 10\%$ . Groundwater samples were then collected using the appropriate bottles (Table 1). Groundwater samples were filtered in the field through a 0.45 µm membrane filter prior to the analysis of metallic ions. Two or three drops of ultrapure nitric acid (HNO3) were added to the groundwater samples used for analyzing the major cation and trace element concentrations. Groundwater samples were kept at four degrees Celsius until they were sent to a commercial laboratory (Bureau Veritas Group) for the chemical analyses of inorganic constituents; these analyses included major and minor elements, metals, nutrients, ammonium, and sulfides. Total alkalinity was determined in the field by titration (HANNA-HI3811).

Bottles	Parameters	Preservatives	<b>Filters</b> NA		
Polyethylene 500 mL	Suspended solids Total solids TDS	NA			
Polyethylene 250 mL	Al, Sb, Ag, As, Ba, Be, Bi, B, Ca, Cr, Co, Cu, Fe, Li, Mg, Mn, Mo, Ni, K, Pb, Se, Si, Na, Sr, Sn, Ti, U, V, Zn	0.5 mL 69% HNO <sub>3</sub>	0.45 µm membrane		
Polyethylene 250 mL	Total alkalinity, NO <sub>2</sub> and NO <sub>3</sub> , Br, Cl, F, SO <sub>4</sub>	NA	NA		
Polyethylene 250 mL	Sulfur	0.5 mL 5 N NaOH	NA		
Polyethylene 250 mL	Inorganic phosphorus	0.5 mL 50% H <sub>2</sub> SO <sub>4</sub>	NA		
Glass 60 mL	Ammoniacal nitrogen	0.17 mL sulfuric acid	NA		
Polyethylene 30 mL	$\delta^{18}$ O, $\delta^{2}$ H	NA	NA		

Table 1. Chemical elements analyzed by the accredited laboratory, Bureau Veritas Group.

Four precipitation gauges were installed within the Grenville Province and one precipitation gauge on the St. Lawrence Platform (Figure 2). Precipitation gauges are designed to collect both rain and snow for stable isotope analyses ( $\delta^2$ H and  $\delta^{18}$ O). The gauge precipitation was collected monthly over 12 months (1 July 2019 to 1 June 2020). We collected 20 groundwater samples in summer 2020 for  $\delta^2$ H and  $\delta^{18}$ O analyses. The  $\delta^2$ H and  $\delta^{18}$ O precipitation gauge data were used to trace the local meteoritic water line (LMWL). The  $\delta^2$ H vs.  $\delta^{18}$ O ratios were determined for each hydrogeological process and served to validate the impact of the geological environment on isotopic chemistry. Complete physicochemical data for the groundwater and precipitations are presented in Table 1. The location of groundwater samples collected for this study is shown in Figure 2.



**Figure 2.** Location of the sampling sites classified according to their respective aquifer type and geological province. The St. Lawrence Platform has 67 sampling sites distributed in granular (G-SLP) and bedrock aquifers (B-SLP), whereas Grenville Province has 192 sampling sites distributed in granular (G-GP) and bedrock aquifers (B-GP). The five rainwater gauges are represented by yellow stars. The 20 sampling sites for isotopic analyses are represented by yellow squares.

#### 2.2.2. Preparation of the Regional Database

The analytical results were validated by testing the ionic balance, for which a value of  $\pm 10\%$  is considered acceptable [28]. Accordingly, 25 samples were rejected because their ion balance values fell outside the recommended range. Consequently, 289 out of 314 collected samples were retained for statistical analysis. Then, the regional database was separated into four data sets according to geological province and aquifer type (granular deposits and bedrock aquifers within the Grenville Province and granular deposits and bedrock aquifers within the St. Lawrence Platform). The database included GPS coordinates and geochemical results in meq/L and mmol/L. Samples with an unknown aquifer were removed from the database. Therefore, a total of 259 samples were retained: 81 samples in granular deposit aquifers within the Grenville Province (G-GP), 111 samples in bedrock aquifers within the Grenville Province (B-GP), 47 samples in granular deposit aquifers within St. the Lawrence Platform (G-SLP), and 20 samples in bedrock aquifers within St. the Lawrence Platform (B-SLP). Figure 2 shows the sample locations of the 259 retained groundwater samples, the five rainwater gauges, and the 20 samples collected for stable isotope analyses.

## 2.2.3. Univariate and Multivariate Statistical Analyses

Univariate statistical analysis provided a summary of groundwater chemistry by describing the variables by parameters of central tendency (i.e., mean, mode, and median),

position (i.e., quartiles), and dispersion (i.e., standard deviation and variance) [29]. To compare groundwater chemistry between the samples, we performed univariate statistics of the major and minor elements and the in situ parameters belonging to each data set.

Multivariate statistical methods commonly used in hydrogeology studies include the correlation matrix and principal component analysis (PCA) [30]. These statistics are based on a unique central value corresponding to the mean, median, and mode of the statistical distribution; therefore, data must respect a normal distribution. Minor and trace elements do not follow a normal distribution, however, because most samples have either very low or below-detection-limit concentrations [31]. In this study, only silica (SiO<sub>2</sub>) followed a normal distribution. The other elements were log-transformed (Table 1). To ensure that each element in Table 1 had equal weight in the statistical analyses, we standardized the values by calculating their standard scores ( $Z_i$ ):

$$Z_{i} = \frac{x_{i} - \bar{x}}{s} \tag{1}$$

where  $Z_i$  is the standard score of sample i,  $x_i$  is the value of sample i,  $\bar{x}$  is the mean, and s is the standard deviation [32]. Elements used in multivariate statistical methods should have a detection ratio greater than 75% [14]. A total of 10 elements (Ca<sup>2+</sup>, Mg<sup>2+</sup>, Na<sup>+</sup>, K<sup>+</sup>, Ba<sup>2+</sup>, SiO<sub>2</sub>, Sr<sup>2+</sup>, HCO<sub>3</sub><sup>-</sup>, Cl<sup>-</sup>, SO<sub>4</sub><sup>2-</sup>) respected the 75% threshold and were used in the multivariate analysis.

The correlation matrix is a bivariate statistical approach that is considered to be the initial step of multivariate analysis [33]. The significance of the correlation coefficient depends on the number of samples (N) in each correlated set (Equation (2)). The correlation is significant when

$$\frac{\left|\mathbf{r}_{xy}\right|}{\sqrt{1-r_{xy}^{2}}} > \mathbf{s}(\alpha/2) \tag{2}$$

where  $r_{xy}$  is the correlation coefficient, N is the sample size, and  $s(\alpha/2)$  is the *t*-score for a significance level  $\alpha$  [29].

Principal component analysis (PCA) extracts the main axes of variation within the data set, with the first component (first axis) corresponding to the dimension that explains the greatest amount of variance within the original data, the second component explaining the second-greatest amount of variance, and so forth [34]. According to the Kaiser criterion, components with an eigenvalue greater than one can be deemed reliable for use in interpreting the output [35]. Typically, the first two or three principal components sufficiently explain variation within a data set. The correlation matrix and PCA were applied to each data set using Rstudio software [36]. R-mode technique and Varimax normalized rotation were applied to maximize the variance of the first three principal axes [31]. Elements used in the correlation matrix and the PCA included major ( $Ca^{2+}$ ,  $Mg^{2+}$ ,  $Na^+$ ,  $K^+$ ,  $HCO_3^-$ ,  $Cl^-$ , and  $SO_4^{2-}$ ) and minor ( $Ba^{2+}$ ,  $SiO_2$ ,  $Sr^{2+}$ ) elements.

## 2.2.4. Saturation Index

The saturation index (SI) compares the logarithm of the solubility product of the mineral ( $K_{sp}$ ) and ion activation product (IAP). The solubility is the amount of solute that can be dissolved, and the IAP involves the measured activities of the product [27]. The SI is expressed as:

$$SI = \log \frac{IAP}{K_{sp}}$$
(3)

At equilibrium, a SI value of 0 indicates that the solution is saturated with the mineral. For SI values less than 0, there is a dissolution process reflected by undersaturated solutions, and for SI greater than 0, there is a precipitation process reflected by an oversaturated solution with respect to the considered mineral [27]. The SI of the samples is calculated using PHREEQC [37].

# 3. Results

# 3.1. Statistical Analysis

Table 2 presents the median concentrations of major and minor elements and the in situ parameters. Groundwater type is defined by a combination of the major cation (Ca<sup>2+</sup>, Mg<sup>2+</sup>, Na<sup>+</sup>, K<sup>+</sup>) and anions (HCO<sub>3</sub><sup>-</sup>, Cl<sup>-</sup>, and SO<sub>4</sub><sup>2-</sup>). In both the St. Lawrence Platform and Grenville Province, the groundwater type in descending order of dominance is Ca-HCO<sub>3</sub>, Na–HCO<sub>3</sub>, and Na–Cl. For the granular aquifers, most groundwater samples in G-GP are Ca–HCO<sub>3</sub> (77% of samples), whereas a slightly lower percentage of samples in the G-SLP (62%) are Ca–HCO<sub>3</sub>. G-GP samples show the lowest median concentrations for all major and minor elements, whereas the G-SLP has the most elevated median concentration of calcium.

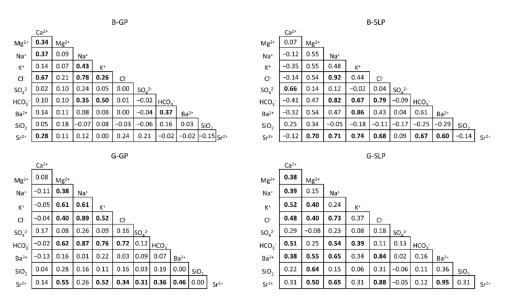
**Table 2.** The median concentration of major and minor elements and in situ parameters. Major and minor elements with the highest median concentrations are in bold.

	G-GP	B-GP	G-SLP	B-SLP St. Lawrence Platform		
	Grenville Province	Grenville Province	St. Lawrence Platform			
Aquifer type	Granular	Bedrock	Granular	Bedrock		
n	81	111	47	20		
	Ca-HCO <sub>3</sub> (62)	Ca-HCO <sub>3</sub> (59)	Ca-HCO <sub>3</sub> (29)	Ca-HCO <sub>3</sub> (30)		
	Na-HCO <sub>3</sub> (10)	Na-HCO <sub>3</sub> (32)	Na-HCO <sub>3</sub> (8)	Na-HCO <sub>3</sub> (24)		
Water type ( <i>n</i> )	Ca-Cl (5)/Na-Cl (4)	Ca-Cl (4)/Na-Cl (13)	CaCl (1)/NaCl (7)	Ca-Cl (5)/Na-Cl (17		
	(Ca-Na)SO <sub>4</sub> (0)	(Ca-Na)SO <sub>4</sub> (3)	(Ca-Na)SO <sub>4</sub> (2)	(Ca-Na)SO <sub>4</sub> (2)		
		Major elements (mg/L)				
Calcium	14.00	20.00	28.00	19.00		
Magnesium	1.90	4.00	4.60	9.95		
Potassium	1.40	2.00	2.30	7.45		
Sodium	6.30	19.00	27.00	97.00		
Chloride	3.40	15.00	34.00	36.50		
Sulfate	7.80	16.00	11.00	8.00		
Bicarbonate	42.00	90.00	110.00	195.00		
		Minor elements (mg/L)				
Barium	0.014	0.021	0.042	0.230		
Manganese	0.008	0.010	0.004	0.016		
Silicium	5.60	5.90	5.40	7.20		
Strontium	0.083	0.22	0.17	0.52		
Fluoride	0.05	0.18	0.08	0.69		
Ammonium	0.02	0.05	0.04	0.59		
Iron	0.06	0.06	0.06	0.06		
Sulfides	0.02	0.02	0.02	0.04		
		In situ parameters				
Dissolved O <sub>2</sub> (%)	1.33	0.00	0.00	0.00		
pH	6.25	7.73	6.47	8.10		
TDS (mg/L)	32.00	103.00	161.00	204.00		

The bedrock aquifer samples for B-GP are mainly Ca–HCO<sub>3</sub> (53% of samples) with some Na–HCO<sub>3</sub> (29%), (Na–Ca)–Cl (15%), and (Na–Ca)SO<sub>4</sub> (3%). Bedrock groundwater samples from B-SLP are mainly Ca–HCO<sub>3</sub> (38%) and Na–HCO<sub>3</sub> (31%) with some (Na–Ca)–Cl (28%) and (Na–Ca)SO<sub>4</sub> (3%). The B-SLP samples are characterized by the highest median concentrations for all major and minor elements, calcium and sulfate excepted, as well as the highest levels of TDS and the most alkaline waters (median pH = 8.1). The median of TDS is under 300 mg/L, which ranks as excellent according to the guidelines for Canadian drinking water quality [38].

#### 3.2. *Correlation Matrix*

The correlation matrix of each data set is presented in Figure 3, and significance was evaluated using Equation (2). For the G-SLP groundwater data set,  $Ca^{2+}$  is significantly correlated (>0.38) with almost all other elements, except  $SO_4^{2-}$ ,  $SiO_2$ , and  $Sr^{2+}$ , whereas within the G-GP groundwater data set,  $Ca^{2+}$  is not significantly correlated (>0.29) with the other elements. Ba<sup>2+</sup> in the granular aquifers of the St. Lawrence Platform (G-SLP) is strongly correlated to multiple elements.



**Figure 3.** Correlation matrix of elements for the four data sets. Coefficient correlations in bold characters are considered as significant correlations according to Equation (2).

In the G-GP data set,  $Mg^{2+}$ ,  $Na^+$ ,  $K^+$ ,  $Cl^-$ ,  $HCO_3^-$ , and  $Sr^{2+}$  present multiple significant correlations with other elements.  $Sr^{2+}$  is often significantly correlated with other elements in the B-SLP, G-SLP, and G-GP data sets, although it has almost no correlation with other elements in the B-GP data set. We note significant correlations (>0.25) for Ca<sup>2+</sup>, Na<sup>+</sup>, and Cl<sup>-</sup> in the B-GP data set, whereas the B-SLP data set shows a significant correlation (>0.57) for HCO<sub>3</sub><sup>-</sup> and Sr<sup>2+</sup>. The correlation matrix of groundwater chemistry illustrates differences between the four groups on the basis of the elements Ca<sup>2+</sup>, Mg<sup>2+</sup>, Na<sup>+</sup>, HCO<sub>3</sub><sup>-</sup>, Ba<sup>2+</sup>, and Sr<sup>2+</sup>.

## 3.3. Principal Component Analysis

Ten components for each data set were extracted from the PCA. The first four components for G-GP and B-GP and the first three components for G-SLP and B-SLP had eigenvalues greater than 1 (Table 3). Loading values correspond to the importance of variance and vary from -1 to 1. There is no literature on determining a minimum threshold for loading values to be meaningful. We considered loadings greater than 0.7 as significant (bold values in Table 3).

	B-GP				G-GP			B-SLP			G-SLP			
Components														
Elements	C1	C2	C3	C4	C1	C2	C3	C4	C1	C2	C3	C1	C2	C3
Ca <sup>2+</sup>	0.75	-0.03	0.10	0.42	-0.15	-0.08	0.80	0.29	-0.11	-0.22	0.88	0.27	0.79	0.13
Mg <sup>2+</sup>	0.21	0.00	0.19	0.75	0.47	0.47	0.10	0.53	0.44	0.62	0.18	0.28	0.33	0.81
Na <sup>+</sup>	0.76	0.44	0.17	-0.18	0.97	-0.01	0.04	-0.02	0.94	0.24	0.04	0.77	0.40	-0.20
$K^+$	0.31	0.73	-0.05	-0.13	0.70	0.44	-0.04	0.21	0.27	0.91	-0.4	0.18	0.67	0.21
Cl-	0.97	0.03	0.03	0.02	0.89	0.02	0.06	0.03	0.95	0.19	-0.01	0.93	0.20	0.11
$SO_4^{2-}$	-0.06	0.13	0.80	-0.05	0.22	0.16	0.70	-0.31	0.05	0.09	0.93	0.13	0.41	-0.49
HCO <sub>3</sub> -	0.02	0.87	-0.03	0.15	0.90	0.16	0.01	0.22	0.79	0.43	-0.23	0.04	0.82	-0.01
Ba <sup>2+</sup>	-0.11	0.50	0.07	0.39	-0.06	0.84	-0.14	-0.10	0.24	0.89	-0.08	0.90	0.14	0.30
SiO <sub>2</sub>	-0.07	0.14	-0.36	0.57	0.11	-0.06	0.01	0.80	-0.08	-0.15	-0.02	0.20	0.12	0.76
$Sr^{2+}$	0.25	-0.14	0.67	0.10	0.29	0.79	0.32	0.06	0.61	0.64	0.07	0.93	0.06	0.27
Eigenvalue	2.29	1.78	1.30	1.30	3.43	1.80	1.27	1.21	3.12	2.77	1.76	3.37	2.25	1.74
Explained variance %	22.9	17.8	13.0	13.0	34.3	18.0	12.7	12.1	31.2	27.7	17.6	33.7	22.5	17.4
Cumulative % of variance	22.9	40.7	53.7	66.7	34.3	52.3	65.0	77.1	31.2	58.9	76.5	33.7	56.2	73.6

**Table 3.** Principal component loadings and explained variance for the four groups separated by geological provinces and aquifer types. Bold values indicate a loading > 0.70; a threshold applied for establishing significance. GP, Grenville Province; SLP, St. Lawrence Platform; G: granular aquifer; B: bedrock aquifer.

For the Grenville Province data sets, the first four components of B-GP explain 22.9%, 17.8%, 13.0%, and 13.0%, respectively, of the variance in the data set and a cumulative variance of 66.7%. Component 1 is characterized by high positive loadings of Ca<sup>2+</sup>, Na<sup>+</sup>, and Cl<sup>-</sup>. Component 2 has significant positive loadings values for K<sup>+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas components 3 and 4 have significant positive loadings for SO<sub>4</sub><sup>2-</sup> and Mg<sup>2+</sup>, respectively. The first four components of G-GP explain 34.3%, 18.0%, 12.7%, and 12.1% of the variance (77.1% of the total variance) in the data set. Component 1 is characterized by significant positive loadings for Na<sup>+</sup>, K<sup>+</sup>, Cl<sup>-</sup>, and HCO<sub>3</sub><sup>-</sup>, and component 2 is characterized by significant positive loading for Ca<sup>2+</sup> and SO<sub>4</sub><sup>2-</sup>, and component 4 has a significant positive loading for SiO<sub>2</sub>.

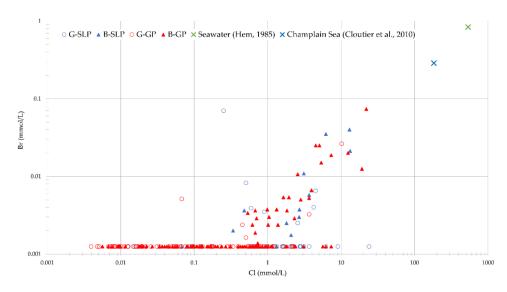
For the St. Lawrence Platform samples, the first three components of B-SLP explain, respectively, 31.2%, 27.7%, and 17.6% of the variance (76.5% in total) in the data set. Component 1 is characterized by significant positive loadings for Na<sup>+</sup>, Cl<sup>-</sup>, and HCO<sub>3</sub><sup>-</sup>, whereas component 2 has significant positive loadings for K<sup>+</sup> and Ba<sup>2+</sup>, and component 3 has significant positive loading values of Ca<sup>2+</sup> and SO<sub>4</sub><sup>2-</sup>. The first three components of G-SLP explain 33.7%, 22.5%, and 17.4% of the variance in the data set (73.6% in total). Component 1 is characterized by significant positive loadings for Na<sup>+</sup>, Cl<sup>-</sup>, Ba<sup>2+</sup>, and Sr<sup>2+</sup>. Component 2 has significant positive loadings for Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas component 3 has significant positive loadings for Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas component 3 has significant positive loadings for Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas component 3 has significant positive loadings for Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas component 3 has significant positive loadings for Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas component 3 has significant positive loadings for Ca<sup>2+</sup> and HCO<sub>3</sub><sup>-</sup>, whereas component 3 has significant positive loadings for Mg<sup>+</sup> and SiO<sub>2</sub>.

Overall, Na<sup>+</sup> and Cl<sup>-</sup> concentrations appear as the key elements for explaining the variance within the data sets because these elements have high positive loadings on the first component for all the four aquifer–bedrock groups. In contrast,  $Mg^{2+}$  and SiO<sub>2</sub> explain very little variance within the four data sets.

## 4. Interpretation and Discussion

## 4.1. Hydrogeological Processes Affecting Groundwater Chemistry

The postglacial Champlain Sea incursion onto the St. Lawrence Platform and Grenville Province at around ca. 11,000 to 13,000 year BP deposited a thick layer of marine clays [12,39] to create new hydrogeochemical conditions within the regional aquifers. Groundwater in some confined aquifers is characterized by higher concentrations of Na<sup>+</sup> and Cl<sup>-</sup> owing to stagnant saline water and marine clay pore waters [12,13]. On a Cl<sup>-</sup> vs. Br<sup>-</sup> plot with seawater [40] and Champlain Sea [25] samples (Figure 4), a number of samples from both the Grenville Province (n = 25) and St. Lawrence Platform (n = 11) geological provinces, plot along the seawater mixing line.



**Figure 4.** Binary diagram plotting Br<sup>-</sup> vs. Cl<sup>-</sup> vs. in mmol/L. Samples are represented according to their geological province and their type of aquifer. The green cross X represents the composition of average seawater [40]. The blue 'x' represents the composition of the Champlain Sea [25].

These samples along the seawater mixing line suggest a process of freshening the Champlain Sea groundwater. The freshwater samples are often dominated by  $Ca^{2+}$  and  $HCO_3^{-}$  ions, resulting from carbonate dissolution or/and silicate weathering, whereas the trapped seawater is dominated by Na<sup>+</sup> and Cl<sup>-</sup>. During this mixing, Na<sup>+</sup> is desorbed, resulting in an exchange of Ca<sup>2+</sup> with Na<sup>+</sup> [13,27] following Equation (4):

$$\frac{1}{2}Ca^{2+} + Na - X \leftrightarrow \frac{1}{2}Ca - X_2 + Na^+ \tag{4}$$

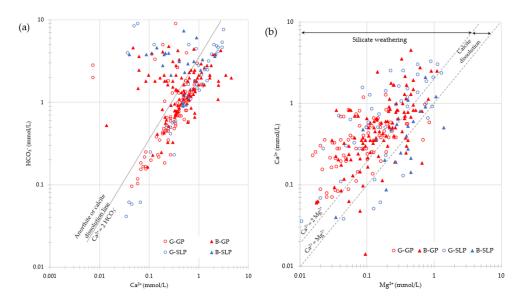
Cloutier et al. (2004) [12] demonstrated that the Champlain Sea incursion caused groundwater salinization through various stages of cation exchange; these exchanges produced Na–HCO<sub>3</sub> and Na–Cl groundwaters.

Carbonate and silicate dissolution in freshwater constitutes one of the main chemical processes in groundwater recharge areas [25]. The equilibrium equations for calcite (Equation (5)) and anorthite (Equation (6)) weathering by carbonic acid are also presented visually in a Ca<sup>2+</sup> vs.  $HCO_3^-$  plot (Figure 5a), and the equilibrium lines for the dissolution of dolomite based on Equation (7) are plotted in an Mg<sup>2+</sup> vs. Ca<sup>2+</sup> plot (Figure 5b) [40,41].

$$Calcite: CaCO_3 + H_2O + CO_2 \leftrightarrow Ca^{2+} + 2HCO_3^{-},$$
(5)

Anorthite : 
$$\operatorname{Ca}Al_2\operatorname{Si}_2O_8 + 3H_2O + 2\operatorname{CO}_2 \leftrightarrow \operatorname{Ca}^{2+} + 2\operatorname{HCO}_3^- + Al_2\operatorname{Si}_2O_5(OH)_4$$
, (6)

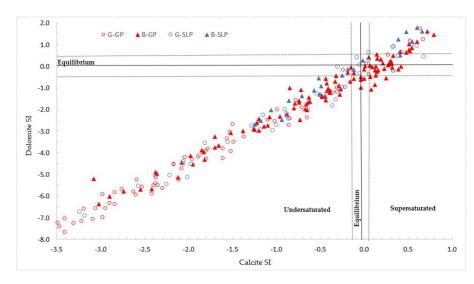
$$Dolomite: CaMg(CO_3)_2 + 2H_2O + 2CO_2 \leftrightarrow Ca^{2+} + Mg^{2+} + 4HCO_3^{-}.$$
(7)



**Figure 5.** Plots of (**a**)  $Ca^{2+}$  vs.  $HCO_3^{-}$  concentrations and (**b**)  $Mg^{2+}$  vs.  $Ca^{2+}$  concentrations. The groundwater samples are identified according to geological province and aquifer type. The equilibrium equations for calcite and anorthite weathering by carbonic acid are plotted in Figures 5a and 5b, and the equilibrium line for the dissolution of dolomite is plotted in Figure 5b. GP, Grenville Province; SLP, St. Lawrence Platform; G: granular aquifer; B: bedrock aquifer.

In Figure 5a, most groundwater samples plot along the equilibrium line of calcite or anorthite dissolution, except for those from the B-SLP data set. Carbonate dissolution depends on the dissolution of carbonic acid (Equations (5) and (6)). Some B-GP and B-SLP groundwater samples fall above the calcite or anorthite equilibrium line indicating that carbonate dissolution is not the dominant process (Figure 5a). The Ca<sup>2+</sup> concentrations of the Grenville Province samples originate mainly from silicate weathering (Figure 5b), corresponding to the presence of silicate-rich rocks and sediments. The Ca<sup>2+</sup> concentrations of the St. Lawrence Platform samples originate from silicate weathering and calcite dissolution (Figure 5b) owing to the presence of silicate-rich sediments and calcium-rich bedrock composed of sandstone conglomerate, dolomite, limestone, and shale [19–24].

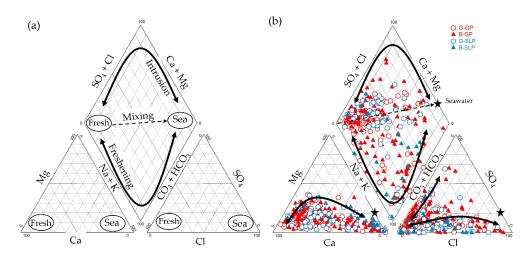
The rainwater generally has a short residence time and is undersaturated in calcite [27]. Our samples follow a linear relationship between dolomite and calcite SI (Figure 6), for which equilibrium is reached at a calcite SI and dolomite SI equal to zero with an uncertainty of  $\pm 0.5$  units for dolomite and  $\pm 0.1$  units for calcite [42,43]. Samples of all four data sets show an evolution from an undersaturation to an oversaturation state and confirm that dissolution and precipitation processes occur for dolomite and calcite. Generally, groundwater in the bedrock aquifers approaches the equilibrium state of calcite and dolomite, whereas the groundwater in granular aquifers is undersaturated in calcite and dolomite. The dissolution of calcite and dolomite is active in the sedimentary rock aquifer systems, and, in our study, it is mainly the result of a water–rock interaction with the dolostone and limestone formations of the St. Lawrence Platform [23,24].



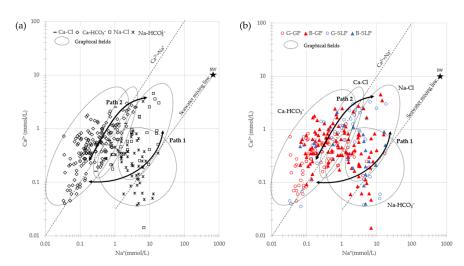
**Figure 6.** The calcite saturation index (SI) vs. dolomite SI. GP, Grenville Province; SLP, St. Lawrence Platform; G: granular aquifer; B: bedrock aquifer.

## 4.2. Chemical Evolution of Groundwater

Groundwater chemical evolution in the Grenville Province and St. Lawrence Platform can be discerned using  $Ca^{2+}$  vs.  $Na^+ log-log$  plots and Piper diagrams [14]. The salinization of groundwater because of the dissolution of salt minerals or mixing with seawater causes groundwater to evolve along a salinization path [13]. The salinization pathways observed for our samples in the Piper diagram (Figure 7) and  $Ca^{2+}$  vs.  $Na^+ log-log$  plot (Figure 8) indicate (i) a gradual evolution from Ca–HCO<sub>3</sub> groundwater toward a more concentrated  $Na-HCO_3$  type and a further evolution toward a Na–Cl groundwater via freshening and/or mixing with seawater, and (ii) evolution of Ca–HCO<sub>3</sub> groundwater toward a Ca–Cl type through intrusion. Most samples from granular aquifers are freshwater in the Piper diagrams (Figure 7), corresponding to Ca–HCO<sub>3</sub> clusters in the  $Ca^{2+}$  vs.  $Na^+ log-log$  plots (Figure 8). The latter plots suggest that the groundwater samples belonging to the four groups were subjected to a  $Ca^{2+}/Na^+$  cation exchange process, whereas the Piper diagrams suggest that cation exchange evolved through freshening, intrusion, and mixing with the fossil water of the Champlain Sea.

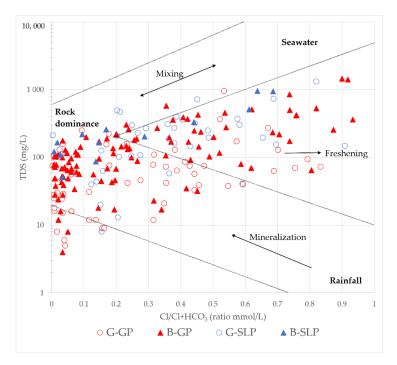


**Figure 7.** Piper diagrams show (**a**) the average composition of freshwater and seawater in relation to freshening, mixing, and intrusion [27,44]. The arrows correspond to the salinization paths presenting the Ca–HCO<sub>3</sub> type in recharge areas toward Na–HCO<sub>3</sub> and Na–Cl waters. The dashed line represents the evolution of freshwater toward seawater via a mixing process [27]. (**b**) Piper diagram with our groundwater samples reported according to their respective aquifer/geological province data sets. GP, Grenville Province; SLP, St. Lawrence Platform; G: granular aquifer; B: bedrock aquifer.



**Figure 8.** Plots of Na<sup>+</sup> vs. Ca<sup>2+</sup> concentrations in mmol/l. (a) The samples are grouped by ellipses representing the dominant facies (graphical fields); (b) the samples are grouped by ellipses representing their respective geological province–aquifer type data sets. The seawater mixing line and seawater concentration (SW) were defined using seawater ratios from Goldberg et al. [45]. GP, Grenville Province; SLP, St. Lawrence Platform; G: granular aquifer; B: bedrock aquifer.

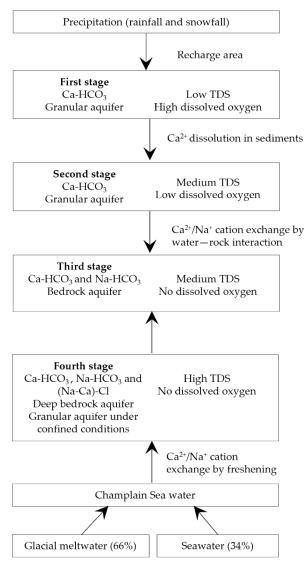
The chemical evolution of groundwater can also be interpreted using Gibbs's diagram (1970) [46]. This diagram of TDS (mg/L) versus the Cl<sup>-</sup>: Cl<sup>-</sup> + HCO<sub>3</sub><sup>-</sup> (mmol/L) is divided into three poles (rainfall, rock dominance, and seawater) and three hydrogeochemical mechanisms (mineralization, freshening, and mixing). Water–rock interactions and freshening appear to be the main processes affecting the Ca<sup>2+</sup>/Na<sup>+</sup> cationic exchange within both geological provinces (Figure 9). Water–rock interactions occur mainly for bedrock aquifers, whereas freshening occurs in both aquifers. Freshening occurs when freshwater recharge replaces saline groundwater originating from a seawater intrusion [47].



**Figure 9.** Gibbs's diagram of the TDS concentration vs. the Cl/Cl+HCO<sub>3</sub> ratio to determine the main process influencing groundwater evolution. Samples are classified according to their geological province and aquifer type. GP, Grenville Province; SLP, St. Lawrence Platform; G: granular aquifer; B: bedrock aquifer.

## 4.3. Regional Hydrogeochemical Conceptual Model

We can summarize the regional hydrogeochemical conceptual model of groundwater in the Grenville Province and St. Lawrence platform in four stages (Figure 10). We can then compare our model with the southeastern part of the Mauricie region of Quebec, a region that was the subject of a previous hydrogeochemical characterization by Lacasse [10]. The southeastern part of the Mauricie region is located between the Lanaudière and Eastern Mauricie region and is also located at the boundary of the Grenville and St. Lawrence Platform geological provinces.



**Figure 10.** A regional hydrogeochemical conceptual model of groundwater evolution and groundwater characteristics, including water type, aquifer type, TDS, dissolved oxygen, and the dominant geochemical processes.

The first stage of the conceptual model is the precipitation in the form of rainfall and snowfall. This precipitation infiltrates the recharge areas and initiates the dissolution reactions with the sediments. The recharging water, corresponding to Ca–HCO<sub>3</sub> groundwater, has highly dissolved oxygen and low TDS.

The second stage is the carbonate and silicate dissolution in freshwater, which constitutes one of the early chemical processes affecting regional groundwater recharge areas [25]. Acidic groundwater more easily dissolves calcite minerals, thereby increasing calcium concentrations in groundwater [27]. Carbonate dissolution is a major factor within the St. Lawrence Platform, with bedrock composed of calcium-rich sedimentary rocks. Silicate dissolution is effective in both the St. Lawrence Platform and Grenville Province, each with silicate-rich sediments. As for the southeastern part of the Mauricie region, the evolution of the Ca–HCO<sub>3</sub> water type toward Na–HCO<sub>3</sub> groundwater owing to cationic exchange is the most important process [10].

The third stage represents a  $Ca^{2+}/Na^+$  cation exchange during the interaction between the bedrock and groundwater. The process of silicate dissolution in bedrock aquifers with carbonic acid transforms Ca–HCO<sub>3</sub> groundwater into Na–HCO<sub>3</sub> waters. The final stage is the Ca<sup>2+</sup>/Na<sup>+</sup> cation exchange through freshening in bedrock and granular aquifers. Our data sets include some samples obtained from deep bedrock aquifers (60–114 m depth) and others from confined conditions under thick layers of clay, which are evolving from Ca– HCO<sub>3</sub> to Na–HCO<sub>3</sub> and finally to Na–Cl, as shown in Figure 9. The mixing of freshwater with seawater leads to the absorption of Ca<sup>2+</sup> and Mg<sup>2+</sup> by ion exchangers—usually clay and the release of Na<sup>+</sup> into the groundwater [27]. These observations suggest that fossil water of the Champlain Sea remains present within both geological provinces and, through Ca<sup>2+</sup> and Na<sup>+</sup> cation exchange, is the primary factor influencing groundwater evolution.

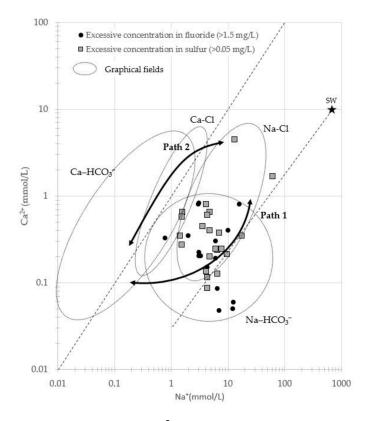
# 4.4. Trace Chemicals and Groundwater Quality

Thirty-three percent of sampled wells in the Lanaudière and Eastern Mauricie regions exceeded the Canadian drinking water aesthetic threshold for manganese (0.02 mg/l), and 10% of the sampled wells exceeded the maximum acceptable level (0.12 mg/L) for Mn. High levels of Mn represent an aesthetic concern for drinking water because Mn discolors water and can stain laundry [48]. Manganese is sensitive to variation in redox conditions [27]. High Mn groundwater samples from both the Grenville Province (n = 20) and St. Lawrence Platform (n = 12) originated from granular aquifers. Three hypotheses could explain the spatial variation in Mn concentrations within granular aquifers. First, the presence of organic matter decreases oxygen and nitrate concentrations in groundwater. This scenario creates a reducing environment that accelerates the solubilization of Mn [49]. Second, Mn concentrations increase in groundwater with the dissolution of Fe-oxyhydroxides under reducing conditions [50]. Third, groundwater is affected by the leakage of domestic sewage, characterized by low Eh and dissolved oxygen levels, in urban environments [49]. The Mn-rich groundwater samples recovered from wells on granular aquifers in the St. Lawrence Platform and groundwater samples collected from the transition zone between lowlands and highlands in Grenville Province were from sites proximal to (40–700 m) natural reducing environments, including bogs, fens, and backwater environments. These groundwater samples also have low dissolved oxygen and low Eh values (-457 to 237 mV). These environmental conditions suggest that the Mn contamination in granular aquifers reflects the introduction of organic matter to produce a natural reducing environment.

For the St. Lawrence Platform, 19% of the sampled wells exceeded the Canadian drinking water aesthetic threshold for sulfur (0.05 mg/L), expressed as hydrogen sulfide [48]. Sulfide in groundwater is produced through a variety of natural sources, including the dissolution of soluble  $SO_4^{2-}$  and  $S^{2-}$  minerals in sedimentary rocks. Sulfide can also occur naturally in coal, natural gas, and oil [51,52]. Furthermore, sulfide can originate from human activities, such as residential waste or atmospheric, industrial, and agricultural contamination [51]. The sampled wells with an excessive concentration in sulfide reach mainly deep, bedrock aquifers (15–107 m), characterized by no dissolved oxygen, basic pH (7.34–8.89), and low Eh rates (-368 to -225 mV). These physicochemical properties suggest that the sulfide contamination could be (i) generated from the anaerobic degradation of organic matter-containing sulfide in the sedimentary rocks of St. Lawrence Platform or (ii) introduced from natural gas; its presence within the St. Lawrence Platform was well documented by Pinet et al., Castonguay et al., Moritz et al., and Ladevèze et al. [53–56].

Although fluoride is added to the drinking water in many municipal water supplies to prevent tooth decay where fluoride is absent—high fluoride concentrations can lead to dental fluorosis and skeleton fluorosis [48]. Four percent of the sampled wells within the Grenville Province exceeded the Canadian drinking water and the World Health Organization threshold of 1.5 mg/l set to avoid fluorosis [48]. Fluoride concentration has primarily a geogenic origin for which the potential sources include fluorite, fluorapatite, amphibole, micas, and biotite [13,57]. Fluorite (CaF<sub>2</sub>) is the most common mineral containing fluorine. Fluoride concentrations increase when fluorite precipitation leads to calcium-poor groundwater or when  $Ca^{2+}$  is exchanged with  $Na^+$  leading to a decreased  $Ca^{2+}$  content in groundwater [14,58]. Fluoride concentrations above the 1.5 mg/L thresholds are mainly associated with wells in the Eastern Mauricie region within the Grenville Province, in locations characterized by a lack of dissolved oxygen, basic pH (7.4-8.8), deep bedrock aquifers (60–170 m depth), and Na–HCO<sub>3</sub> and Na–Cl water types. The concentration of fluoride remains constant in seawater at approximately 1.3 mg/L [58]. Given that the fluoride concentrations of the F-impacted wells range from 1.7 to 7.3 mg/L, seawater cannot be the sole source of fluoride in these samples. Fluorine can substitute with hydroxyl ions in amphibole, biotite, chlorite, and tourmaline because of the similarity between their ionic radius [59]. As the geology at the F-impacted sampling wells is tonalitic and trondhjemitic gneiss containing quartz, plagioclase, biotite, and amphibole, it is quite likely that these rocks are the main source of fluoride in groundwater [19–22].

The sampled wells with an excessive concentration in sulfide and fluoride are reported on  $Ca^{2+}$  vs. Na<sup>+</sup> log–log plots (Figure 11). These samples evolve along the salinization path, suggesting that the groundwater samples belonging to the excessive concentration in sulfide and fluoride were subjected to a  $Ca^{2+}/Na^+$  cation exchange process.

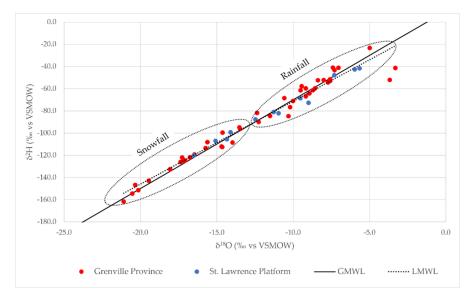


**Figure 11.** Plot of Na<sup>+</sup> vs. Ca<sup>2+</sup> concentrations in mmol/l. Samples with an excessive concentration in sulfide and fluoride evolve along the salinization path 1. The seawater mixing line and seawater concentration (SW) were defined using seawater ratios from Goldberg et al. [45].

Manganese, sulfur, and fluoride represent minor elements that could improve our understanding of the chemical composition of the aquifer, as these elements are spatially correlated to the specific geological context within the larger geological provinces, i.e., sulfur in the St. Lawrence Platform and fluoride in the Grenville Province.

#### 4.5. Isotopic Validation

We relied on the isotopic data of regional precipitation to establish the local meteoritic water line (LMWL) on a  $\delta^2$ H vs.  $\delta^{18}$ O binary diagram (Figure 12). The collected precipitation samples plot along the global meteoritic water line (GMWL), indicating no significant isotopic fractioning by evaporation. The global isotopic variation in rainwater is affected by kinetic fractionation during condensation and evapotranspiration [27]; rainfall is mainly isotopically depleted due to (i) altitude and (ii) the distance from the coast [60]. Two outlier points marked by low  $\delta^2$ H/ $\delta^{18}$ O ratios represent snowfall samples collected at high altitudes in St.-Michel-des-Saints (421 m) and La Tuque (167 m), Quebec.

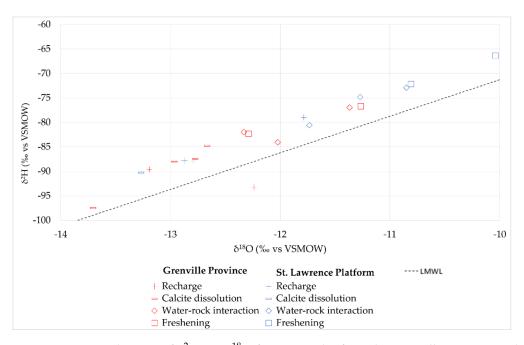


**Figure 12.** Binary diagram of  $\delta^2$ H vs.  $\delta^{18}$ O of precipitation data collected at the five precipitation gauges within the Grenville Province and St. Lawrence Platform. GMWL, global meteoritic water line; LMWL, local meteoritic water line.

In the present study, the isotope signatures of precipitation are interpreted according to 1 year times series data. However, long-term monitoring of the local precipitation stable isotope signatures at the investigated site would be valuable for future similar studies.

 $\delta^2$ H vs.  $\delta^{18}$ O isotopic ratios are determined for 20 samples to validate the regional hydrogeochemical conceptual model. Each hydrogeological process presented in Section 4.3 is represented by at least four samples. When we compare the  $\delta^2$ H vs.  $\delta^{18}$ O ratios of the 20 groundwater samples to the LMWL and include the respective geochemical evolution processes for each sample (Figure 13), we observe calcium dissolution for groundwater samples with lower isotopic ratios; the values correspond to cooler conditions. Under such conditions, the concentration of CO<sub>2(aq)</sub> increases to form carbonic acid, H<sub>2</sub>CO<sub>3</sub>, which decreases the pH of groundwater [27].

The groundwater samples dominated by freshening and water–rock interactions show relatively heavy isotopic ratios (Figure 13). The  $\delta^2$ H of samples undergoing freshening varies between -82.3 and -66.4, and the  $\delta^{18}$ O varies from -12.3 to -10.0. Samples strongly influenced by water–rock interactions have  $\delta^2$ H values varying between -84.1 and -72.9 and  $\delta^{18}$ O values ranging between -12.3 and -10.8. The volume of freshwater able to be mixed with seawater is greater during the early spring/late summer, a warmer period of the year characterized by heavier isotopes. Recharge area samples have  $\delta^2$ H values between -93.2 and -79.0 and  $\delta^{18}$ O values from -13.2 to -11.8. Finally, samples experiencing calcite dissolution present relatively lower isotopic ratios, i.e.,  $\delta^2$ H varying from -97.5 to -84.8 and  $\delta^{18}$ O varying from -13.7 to -12.7. These isotopic results suggest that the physicochemical parameters of groundwater, such as temperature, have a greater impact on hydrogeochemical processes affecting groundwater chemistry than the immediate geological environment, e.g., bedrock geology.



**Figure 13.** Binary diagram of  $\delta^2$ H vs.  $\delta^{18}$ O for 20 samples from the Grenville Province and St. Lawrence Platform. The dashed line represents the local meteoritic water line (LMWL).

### 5. Conclusions

In this study, the influence of the geological context on the chemical evolution of groundwater was investigated in the Lanaudière and Eastern Mauricie administrative regions of Quebec, two areas that cover portions of two geological provinces, i.e., the Grenville Province (crystalline rocks) and St. Lawrence Platform (sedimentary rocks), both covered by glacial and glaciomarine sediments (ca. 80,000 year BP to 11,000 year BP). We undertook chemical analyses of inorganic chemical constituents (major and trace elements) and stable water isotopes ( $\delta^2$ H and  $\delta^{18}$ O) on groundwater samples and then used bivariate plots, statistical analyses, and geochemical modeling methods to determine the key controls on groundwater evolution within each region.

The outcomes of this study are presented in the form of a conceptual model of groundwater geochemical evolution. In the study area, groundwater chemical evolution is mainly governed by four hydrogeochemical mechanisms regardless of the specific geology of the geological provinces: (i) precipitation infiltrating the recharge areas, (ii) carbonate and silicate dissolution by freshwater, (iii) Ca<sup>2+</sup>/Na<sup>+</sup> cation exchange through water–rock interactions, and (iv) Ca<sup>2+</sup>/Na<sup>+</sup> cation exchange through freshening. The fossil water of the Champlain Sea remains present in both regional aquifers (crystalline and sedimentary rocks) and is the main driver for groundwater chemical changes. The  $\delta^2$ H vs.  $\delta^{18}$ O isotopic ratios and the similar chemical composition of groundwater suggest that the physicochemical parameters of groundwater have a greater effect on hydrogeochemical mechanisms than the immediate geological environment.

However, we noted some differences in regard to minor elements at the local scale. We observed elevated concentrations of manganese in some groundwater samples in both the Grenville Province and St. Lawrence Platform. The presence of organic matter decreases the oxygen rate in groundwater, creating a reductive environment that accelerates the solubilization of manganese. Some groundwater samples collected from the sedimentary rock aquifers of the St. Lawrence Platform also revealed an excess of sulfide. This sulfide likely originated from either the anaerobic degradation of organic materials containing sulfide or by the incursion of natural gas. Some deeper wells in the Grenville Province had high fluoride concentrations. The local geology at these wells appears to be the main source of fluoride in groundwater as the local bedrock is composed of tonalitic and trondhjemitic

gneiss. Therefore, minor elements in addition to major elements make it possible to improve our understanding of how the geological context affects groundwater chemistry.

This study not only characterizes the natural groundwater quality of our study area but also contributes to a better understanding of groundwater chemical evolution within aquifers belonging to two different geological contexts. Our results can be used to elaborate effective groundwater resource management/protection strategies by identifying areas with high salinity contamination. Nonetheless, expanded knowledge of the lithology of the stratigraphic units within the Lanaudière and Eastern Mauricie aquifers would improve our understanding of the relationship between groundwater chemistry and aquifer geology using trace elements that appear to be better tracers of geological influence on groundwater chemistry than major elements.

**Author Contributions:** Conceptualization, R.T. and J.W.; Methodology, R.T. and J.W.; Software, R.T.; Validation, J.W. and R.C.; Formal Analysis, R.T., J.W. and R.C.; Investigation, R.T. and J.W.; Writing— Original Draft Preparation, R.T.; Writing—Review and Editing, J.W., R.C. and L.B.; Supervision, J.W. and R.C.; Project Administration, J.W. and R.C.; Funding Acquisition, J.W. and R.C. All authors have read and agreed to the published version of the manuscript.

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**Conflicts of Interest:** The authors declare that they have no conflict of interest.

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